

Distribution of lanthanides in amphibole and plagioclase from plagioclases of the Ray-Iz massif (Polar Urals)

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Relevance. This work shows the results of mineralogical, petrographic, and geochemical studies of plagioclases localized among chromite-bearing ultramafic rocks of the Ray-Iz massif (Polar Urals). The spatial confinement of plagioclases to the Tsentralnoye chrome ore deposit (corundum-bearing plagioclases) and the Yugo-Zapadnoye IV ore occurrence (plagioclases without corundum) determine the need for a detailed study of the rocks of the vein series of chromite-bearing sections.

The purpose of the work. The study of mineralogical and petrographic characteristics, as well as the geochemistry of lanthanides in plagioclase and amphibole from plagioclases of the Ray-Iz massif (Polar Urals).

Results. In view of the results of microprobe analysis, it was found that the compositions of amphibole in plagioclases vary considerably in accordance with the modern classification of calcium amphiboles, pargasite, edenite, magnesiohornblende and actinolite. Two types of zoning are determined in amphibole. Amphibole grains with zoning of the first type are characterized by an increase in the contents of Al_2O_3 , TiO_2 , FeO , Na_2O from the center toward the periphery of the grains; when zoning of the second type, on the contrary, there is a decrease in these components. The chemical composition of plagioclase corresponds to oligoclase $An_{10-30.1}$. The decrease in the content of the anorthite molecule from the center to the edge of the grain was noted. The study of nature of the distribution of REE in rock-forming minerals made it possible to determine that amphibole is the main mineral concentrator of lanthanides in plagioclases: the total REE content in the mineral is 250–450 ppm, while 16–18 ppm of lanthanides are concentrated in the plagioclase. It was established that the reason for the existence of two types of distribution spectra of lanthanides identified earlier is the ratio of the total amount of amphibole and plagioclase in the rock. The distribution of rare-earth elements in calcic amphiboles from plagioclases is compared with the distribution of REEs in the amphiboles of the Nyurundukansky mafic complex in the North-Western Baikal region, the studies of which were carried out by S. G. Skublov. Using the amphibole-plagioclase geothermometer by T. Holland, J. Blundy and geobarometer by M. W. Schmidt parameters of amphibole plagioclases formation were determined.

Conclusion. The nature of the distribution of lanthanides in the main rock-forming minerals – plagioclase and amphibole – is found. The comparison of parameters and compositional characteristics of amphiboles allowed us to conclude that there is a direct relationship between temperature, the sum of REE and TiO_2 in the mineral.

Keywords: amphibole, plagioclase, plagioclase, REE, Ray-Iz, Polar Urals.

Introduction

The Ray-Iz ultramafic massif is one of the largest in the Urals and is located at the base of the crust-mantle section of the Voykarskaya Paleozoic paleo-island arc zone overriding the main Ural deep fault to the edge of the Eastern European continent [1].

Plagioclase bodies are found among chromite-bearing ultramafites of the massif – plagioclases without corundum and corundum-bearing ones [1–3]. Both types of plagioclases form a vein series in the rocks of the dunite-harzburgite complex.

A brief history of the study of plagioclases

Corundum-bearing plagioclases are known in the southwestern part of the chrome ore deposit called Tsentral'noye. They compose the well-known occurrence of corundum mineralization called Rubinovy Log, and have been attracted the attention of researchers due to the unique ruby mineralization. Corundums and host rocks were first described by V. P. Sorokin and B. V. Perevozchikov in 1973 [4]. Single microprobe analyses of rock-forming minerals of phlogopite plagioclases from the Rubinovy Log occurrence were first published by N. I. Bryanchaninova and A. B. Makeev with co-authors [5]. Zoning in the structure of rocks with ruby mineralization was described by S. V. Shcherbakova. She has investigated the composition of minerals, and concluded about the genetic relationship of plagioclase rocks and their host ultramafites [6, 7].

The study of the composition and age of plagioclases, both corundum-bearing and corundum-free in recent years [1–3] allowed to establish the existence of two types of distribution spectra of rare earth elements in them. At the same time, the question of the content of rare earths in amphiboles and plagioclases, the main rock-forming minerals of plagioclases, remained open. In this paper, we studied the chemical composition of minerals and the nature of the distribution of lanthanides in plagioclase and amphibole; it is concluded that there are two REE distribution spectra in plagioclases.

Chemical composition and zoning of rock-forming minerals

Microprobe analysis of minerals was performed in the laboratory of physical and chemical research methods at the Institute of Geology and Geochemistry, Ural Branch of the Russian Academy of Sciences (Ekaterinburg, the analyst is V. V. Khiller) using the electronic probe microanalyzer CAMECA SX 100.

In view of the results of microprobe analysis, it was found that the compositions of amphibole vary considerably in accordance with the modern classification of calcic amphiboles [8, 9], pargasite, edenite, magnesiohornblende and actinolite.

Amphibole grains exhibit chemical zoning. Central parts in amphibole are from plagioclase without corundum (sample Y-12/1) of the Yugo-Zapadnoye IV ore occurrence of chrome ores. As a rule, they are represented by hornblende and edenite, and the peripheral ones are represented by a low-alumina, actinolite – a kind of alkalies – poor amphibole. Its Na_2O content is 1.63–2.97 wt.%, and the amount of alumina is in the range of 4.17–9.59 wt.% (Table 1). In addition, there are minor impurities of MnO 0.11–0.26

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wt.%, Cr₂O₃ 0.01–0.21 wt.% and TiO₂ 0.38–0.76 wt.% in the mineral. Amphibole iron content $f = \text{Fe}^{2+} \times 100\% / (\text{Fe}^{2+} + \text{Mg})$ varies from 25.33 to 31.67%. Ferruginosity and the content of Al₂O₃, TiO₂, FeO, Na₂O decrease from the center toward the periphery of the grains (analysis points are shown in Fig. 1; Table 1). The chemical composition of plagioclase corresponds to oligoclase An_{10–24.9}. The amount of anorthite molecule decreases from the center to the edge of the grain (the analysis points are shown in Fig. 1; Table 2).

In amphibole from amphibole plagioclase without corundum (sample Y-200/1), the inverse zoning is shown. The central parts of the grains are actinolite, and the peripheral parts are hornblendite and edenite. The Na₂O content in the mineral is 1.28–2.66 wt.%, and the amount of alumina varies between 3.46–10.51 wt.%. The mineral contains significant impurities of FeO (6.01–8.35 wt.%), as well as small impurities of TiO₂ (0.2–0.54 wt.%) and Cr₂O₃ (0–0.04 wt.%, that is almost at the detection limit of the device). Ferruginosity of amphibole varies from 15.11 to 23.29%. Ferruginosity and the content of Al₂O₃, TiO₂, FeO, Na₂O increase from the center toward the periphery of the grain. The chemical composition of plagioclase corresponds to oligoclase An_{14.8–17.7}.

In amphibolic plagioclase (sample Y-292/2) from the occurrence of corundum mineralization called Rubinovy Log (Tsentralnoye chromium ore deposit), amphibole is in intergrowth with plagioclase (Fig. 2). The central parts of amphibole grains correspond to edenite, and the edge ones correspond to pargasite; some individual actinolite grains are also present. The alumina content in amphibole varies from 3.54 to 15.45 wt.%; Na₂O – from 0.79 to 3.48 wt.%; FeO – 4.73–5.36 wt.%; the amount of Cr₂O₃ is not large and amounts to 0.06–0.18 wt.%. The mineral contains impurities of TiO₂ (0.11–0.15 wt.%). Amphibole ferruginosity in the most aluminiferous varieties (edenite, pargasite) varies from 13.44 to 15.02% and from 11.46 to 12.13% in actinolite. Ferruginosity and Al₂O₃ content, TiO₂, FeO, Na₂O increase toward the periphery. Pargasite is characterized by the highest content of Na₂O and Al₂O₃. The chemical composition of plagioclase corresponds to oligoclase An_{13.6–30.1}.

Plagioclase contains BaO impurities in an amount of from 0 to 0.11 wt.%; K₂O (0.02–0.40 wt.%) and SrO, the amount of which in amphibole plagioclases without corundum (sample Y-12/1, sample Y-200/1) varies from 0 to 0.11 wt.%, and in corundum-bearing amphibole plagioclase (sample Y-292/2) – from 0.98 to 1.26 wt.% (Table 2). Among the regularities in the distribution of chemical elements in plagioclase, it is worth noting an increase in the amount of Na₂O from the central parts of the grains toward the peripheral ones in all the samples studied. The edge parts of the grains are more acidic.

Features of the trace element composition of rock-forming minerals. Geothermobarometry

The study of the trace element composition of the main rock-forming minerals of plagioclase and amphibole in rocks made it possible to establish that amphibole is the main lanthanide-concentrating mineral: the total REE content in the mineral is 250–450 ppm, while 16–18 ppm of lanthanides are concentrated in the plagioclase.

According to the results of the ICP-MS analysis, chondrite (C1) normalization was performed [10] and the behavior of REEs

**Table 1. The chemical composition of amphibole from amphibole plagioclases, wt.%.
Таблица 1. Химический состав амфибола из амфиболовых плагиоклазитов, мас.%.
Sample number**

Sample number	Sample Y-12/1						Sample Y-200/1		Sample Y-292/2	
	1_37 e	1_38 c	2_39 c	2_40 e	3_41 c	3_42 e	5_3 e	5_4 c	13_3 c	13_3 e
SiO ₂	50.86	48.54	52.04	51.98	51.25	52.09	50.92	53.53	47.96	46.12
TiO ₂	0.59	0.59	0.46	0.60	0.67	0.50	0.43	0.30	0.15	0.12
Al ₂ O ₃	7.98	7.93	5.08	5.25	5.41	4.79	6.44	4.89	13.1	15.45
Cr ₂ O ₃	0.04	0.09	0.04	0.02	0.21	0.01	0.01	0.00	0.09	0.06
FeO	9.92	10.30	10.43	10.35	10.73	10.41	7.53	6.50	4.97	5.36
MgO	14.84	15.62	16.47	16.39	16.11	16.49	16.88	18.08	17.96	17.01
MnO	0.19	0.23	0.18	0.28	0.26	0.22	0.16	0.17	0.11	0.11
CaO	10.46	10.44	10.47	10.76	10.62	10.26	11.72	12.51	11.26	11.51
Na ₂ O	2.97	2.38	1.77	1.75	1.76	1.71	2.23	1.57	3.11	3.48
K ₂ O	0.08	0.12	0.06	0.07	0.07	0.07	0.07	0.08	0.16	0.19
Total	97.93	96.24	97.00	97.45	97.09	96.55	96.39	97.63	98.87	99.41
<i>Formula coefficients (rate per 23 oxygen atoms)</i>										
Si	7.24	7.08	7.47	7.44	7.38	7.51	7.31	7.52	6.66	6.41
Ti	0.06	0.07	0.05	0.07	0.07	0.05	0.05	0.03	0.02	0.01
Al	1.34	1.36	0.86	0.89	0.92	0.81	1.09	0.81	2.14	2.53
Al ^{IV}	0.76	0.92	0.53	0.56	0.62	0.49	0.69	0.48	1.34	1.59
Al ^{VI}	0.58	0.44	0.33	0.32	0.30	0.32	0.40	0.33	0.81	0.94
Cr	0.01	0.01	0.01	0.00	0.02	0.00	0.00	0.00	0.01	0.01
Fe	1.18	1.26	1.25	1.24	1.29	1.26	0.90	0.76	0.58	0.62
Mg	3.15	3.40	3.53	3.50	3.46	3.54	3.61	3.79	3.72	3.52
Mn	0.02	0.03	0.02	0.03	0.03	0.03	0.02	0.02	0.01	0.01
Ca	1.60	1.63	1.61	1.65	1.64	1.58	1.80	1.88	1.68	1.71
Na	0.82	0.67	0.49	0.49	0.49	0.48	0.62	0.43	0.84	0.94
K	0.02	0.02	0.01	0.01	0.01	0.01	0.01	0.01	0.03	0.03

Notes: e – grain edge; c – grain centre.

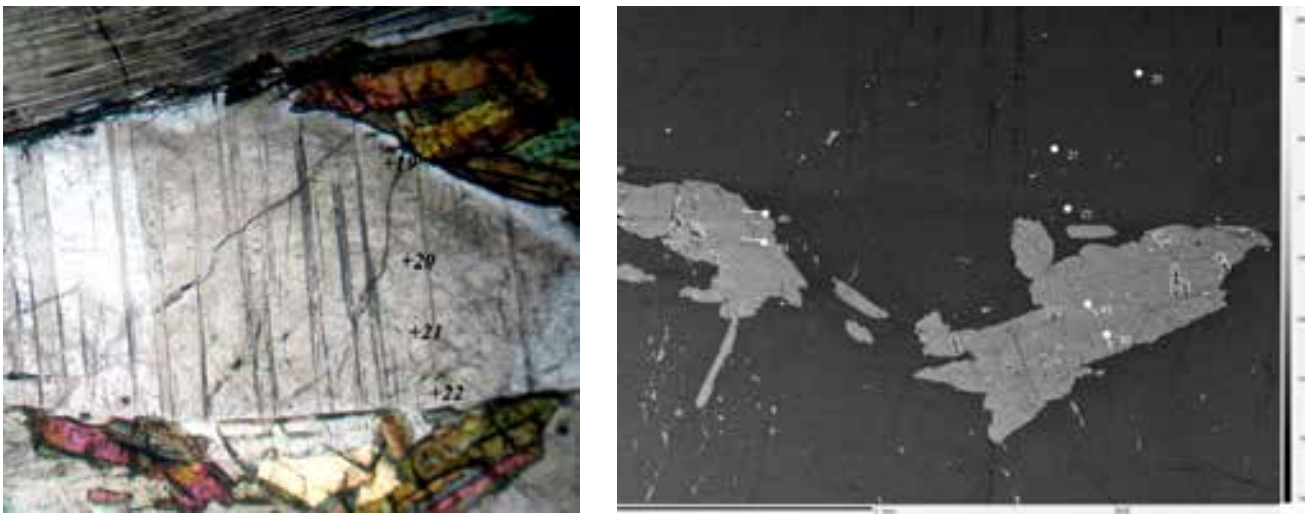


Figure 1. Plagioclase in the intergrowths with amphibole. Sample. Y-12/1. The photo of a thin section with an analyzer is on the left. BSE image is on the right. Analysis points – see Table 1, 2.
Рисунок 1. Плагиоклаз в сростании с амфиболом. Обр. Y-12/1. Слева – фото шлифа с анализатором. Справа – BSE-изображение. Точки анализов – см. табл. 1, 2.

in monomineralic fractions of amphibole and plagioclase from plagioclasites was analyzed. The REE content in amphibole, plagioclase, and in the rock as a whole was determined by ICP-MS analysis at the Institute of Geology, Ural Branch of the Russian Academy of Sciences (Ekaterinburg) and Institute of Minerology of the Ural Branch of the Russian Academy of Sciences (Miass).

Amphiboles and plagioclases of corundum-bearing amphibole plagioclasites (sample Y-292/2) with a distribution spectrum of lanthanides of the first type (Fig. 3, a) are characterized by enrichment with light lanthanides and depletion by heavy ones – the curve has a negative slope for both the main minerals and the rock as a whole.

Amphibole plagioclasites without corundum of the Yugo-Zapadnoye IV ore occurrence of chromite (Fig. 3, b, c) with the distribution spectra of REEs of the second type are characterized by a slight depletion of light lanthanides – the distribution curve for the rock has a sub-horizontal position. If plagioclase, which constitutes more than 90% of the rock thickness, is characterized by enrichment with light lanthanides, then for amphibole contained in plagioclase in an amount of 3 to 10%, the type of REE distribution spectrum is different. In amphibole from plagioclasites without corundum, a significantly lower content of light REEs is observed compared to heavy ones, and a negative europium anomaly is also determined [15] (Fig. 3, b, c).

The analysis of REE distribution in amphiboles depending on the faces of metamorphism was in [11]. When comparing the results of this work with those obtained by S. G. Skublov, a similarity of the REE distribution spectra in the amphiboles of corundum-bearing amphibole plagioclasites is observed (sample Y-292/2), as well as REE spectra of amphiboles formed under granulite facies of metamorphism. Amphiboles from amphibole plagioclasites without corundum are close to the mineral from the rocks of the upper amphibolite facies.

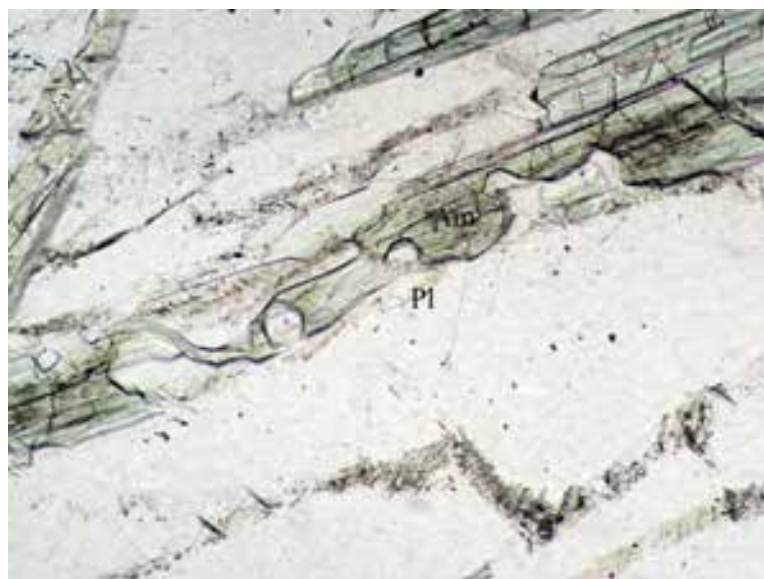


Figure 2. Intergrowth of plagioclase and amphibole (pargasite). Sample. Y-292/2. The photo of a thin section without an analyzer. 100x magnification (1 mm in vertical frame dimension). Am – amphibole, Pl – plagioclase.
Рисунок 2. Сростание плагиоклаза и амфиболом (паргасита). Обр. Y-292/2. Фото шлифа без анализатора. Увеличение 100x (1 мм по высоте кадра). Ам – амфибол, ПИ – плагиоклаз.

**Table 2. The chemical composition of plagioclase from amphibole plagioclases, wt.%.
Таблица 2. Химический состав плагиоклаза из амфиболовых плагиоклазитов, мас. %.**

Sample number	Sample Y-12/1					Sample. Y-200/1				
	19 e	20 c	21 c	22 e	1_1 c	1_2 e	2_15 e	2_16 e	2_17 c	2_18 c
SiO ₂	66.47	61.53	62.31	64.28	64.33	65.01	64.27	64.49	64.50	64.21
TiO ₂	0.00	0.00	0.01	0.00	0.03	0.02	0.00	0.00	0.01	0.03
Cr ₂ O ₃	0.03	0.48	0.01	0.04	0.01	0.03	0.05	0.01	0.01	0.00
Al ₂ O ₃	20.86	22.97	23.02	22.47	22.34	22.10	21.95	21.99	21.98	22.32
FeO	0.09	0.03	0.07	0.01	0.00	0.04	0.02	0.05	0.05	0.04
CaO	2.09	5.13	4.97	3.87	3.51	3.18	3.44	3.43	3.57	3.69
Na ₂ O	10.64	8.44	8.84	9.47	9.60	9.61	9.80	9.68	9.54	9.43
K ₂ O	0.05	0.07	0.07	0.06	0.05	0.40	0.03	0.04	0.05	0.06
SrO	0.06	0.07	0.07	0.00	0.09	0.04	0.04	0.11	0.08	0.09
BaO	0.00	0.00	0.00	0.03	0.00	0.00	0.03	0.00	0.01	0.00
Total	100.34	98.75	99.42	100.26	100.00	100.44	99.63	99.82	99.80	99.88
<i>Formula coefficients (per 5 cations)</i>										
Si	2.91	2.77	2.78	2.83	2.84	2.86	2.84	2.85	2.86	2.84
Ti	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
Cr	0.00	0.02	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
Al	1.08	1.22	1.21	1.17	1.16	1.15	1.15	1.15	1.15	1.16
Fe	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
Ca	0.10	0.25	0.24	0.18	0.17	0.15	0.16	0.16	0.17	0.18
Na	0.90	0.74	0.76	0.81	0.82	0.82	0.84	0.83	0.82	0.81
K	0.00	0.00	0.00	0.00	0.00	0.02	0.00	0.00	0.00	0.00
Sr	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
Ba	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
An., %	10.0	24.9	23.9	18.3	16.8	15.1	16.5	16.5	17.2	17.7

Notes: e – grain edge; c – grain centre.

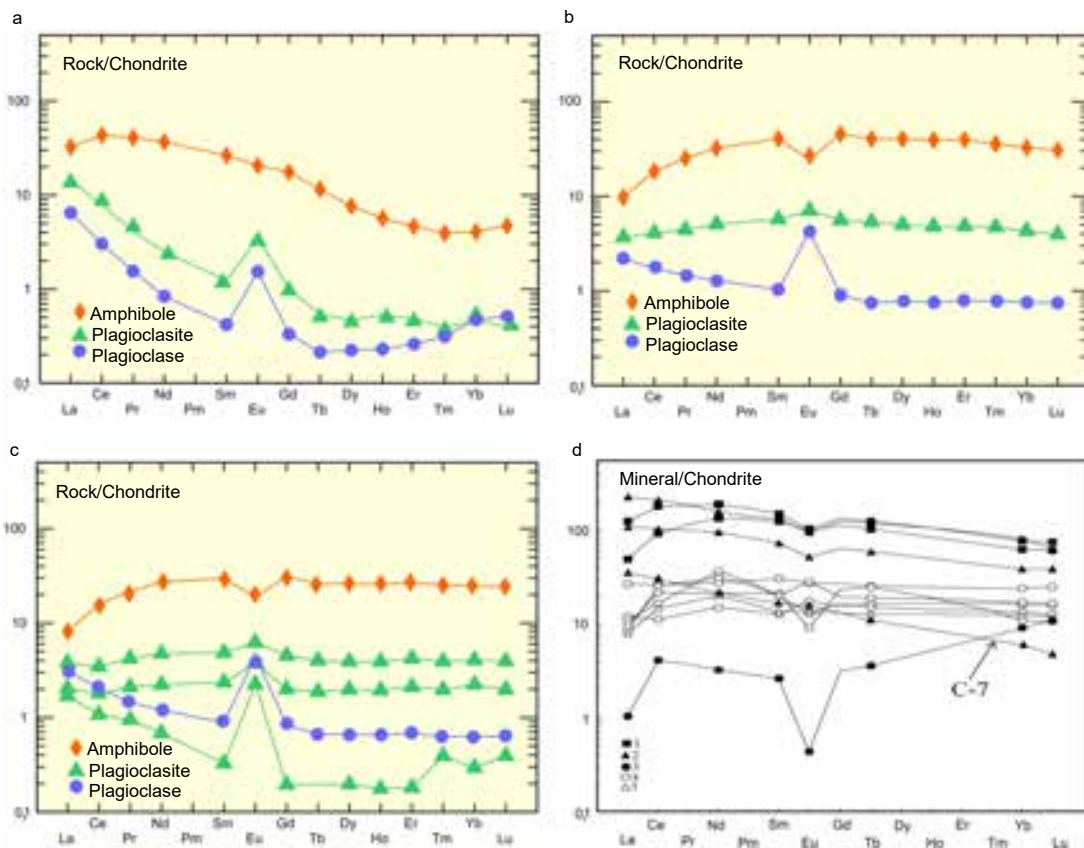


Figure 3. REE distribution spectra in plagioclases and main rock-forming minerals. a – sample Y-292/2 – amphibole plagioclase (Tsentrallye deposit); b – sample Y-12/1– amphibole plagioclase (Yugo-Zapadnoe IV); c – sample Y-200/1– amphibole plagioclase (Yugo-Zapadnoe IV); d – mafic complex of the northwestern Baikal region [11]. Amphibole compositions are shown. 1 – kaersutites, 2 – amphibole granulite facies (C-7 is close to the amphibolite facies of metamorphism), 3 – metasomatic amphiboles, 4 – high-temperature amphibolite facies, 5 – low-temperature amphibolite facies.

Рисунок 3. Спектры распределения РЗЭ в плагиоклазитах и главных породообразующих минералах. а – Y-292/2 – амфиболовый плагиоклазит (месторождение Центральное); б – Y-12/1 – амфиболовый плагиоклазит (рудопоявление Юго-Западное IV); в – Y-200/1 – амфиболовый плагиоклазит (рудопоявление Юго-Западное IV); г – мафический комплекс Северо-Западного Прибайкалья [11]. Показаны составы амфиболов. 1 – керсутитов, 2 – амфиболов гранулитовой фации (С-7 близок к амфиболитовой фации метаморфизма), 3 – метасоматических амфиболов, 4 – высокотемпературной амфиболитовой фации, 5 – низкотемпературной амфиболитовой фации.

Table 3. P–T parameters of plagioclase formation and amphibole compositional features.**Таблица 3. P–T-параметры образования плагиоклазитов и особенности состава амфибола.**

Sample number	T, °C	P, kbar	Amount of REE in amphibole, ppm	TiO ₂ , wt. %
Y-12/1	622–674	3,3–4,2	453,90	0,38–0,76
Y-200/1	591–625	2,2–4,1	330,40	0,20–0,54
Y-292/2	591–612	7,2–9,0	255,62	0,11–0,15

In amphiboles, a direct proportional dependence of TiO₂ and the sum of REEs was determined. The highest amount of REE is 454 ppm, with a TiO₂ content of 0.38–0.76 wt.% is observed in amphibole from plagioclase without corundum (sample Y-12/1); in the amphibole from the sample Y-200/1, the REE amount is 330 ppm and the TiO₂ content is 0.2–0.54 wt.%. The smallest amount of REE among the studied amphiboles (256 ppm) and TiO₂ content of 0.11–0.15 wt.% is observed in amphibole from corundum-bearing plagioclase (sample Y-292/2).

The temperature of formation of amphibole plagioclases was determined by Holland and Blundy amphibole-plagioclase geothermometer [12], and the pressure by Schmidt amphibole geobarometer [13]. For amphibole plagioclase (sample Y-12/1), the calculated pressure was 3.3–4.2 kbar, temperature – 622–674 °C. The pressure for the amphibole plagioclase Y-200/1 is 2.2–4.1 kbar, the temperature is 591–625 °C. For corundum-bearing amphibole plagioclase (sample Y-292/2) from the Tsentralnoye deposit, the parameters are the following: T = 591–612 °C, P = 7,2–9,0 kbar (Table 3) [14].

Conclusion

The comparison of parameters and compositional characteristics of amphiboles allow us to assume a direct relationship between temperature, the sum of REE and TiO₂, as well as the inverse correlation between pressure and total REE content. According to S. G. Skublov, with an increase in the temperature of metamorphism, the total REE concentration in amphiboles increases [11], which is also reflected in our research results.

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REFERENCES

- Vakhrusheva N. V., Ivanov K. S., Stepanov A. E., Shokalsky S. P., Azanov A. N., Khiller V. V., Shiryayev P. B. 2016, Plagioclases from chromite-bearing ultramafic rocks of the Ray-Iz massif. *Litosfera* [Lithosphere], no. 5, pp. 134–145. (In Russ.)
- Vakhrusheva N. V., Ivanov K. S. 2018, *Priroda i vozrast plagioklazitov ul'traosnovnogo massiva Ray-Iz* (Polyarnyy Ural) [The nature and age of plagioclase of the ultrabasic massif Ray-Iz (Polar Ural)]. *Doklady Akademii nauk* [Doklady Earth Sciences], vol. 480, no. 1. pp. 80–84. <https://doi.org/10.7868/S0869565218130169>
- Vakhrusheva N. V., Shiryayev P. B., Stepanov A. E., Bogdanova A. R. 2017, *Petrologiya i khromitnosnost' ul'traosnovnogo massiva Ray-Iz* (Polyarnyy Ural) [Petrology and chromite content of the ultrabasic massif Ray-Iz (Polar Ural)]. Ekaterinburg, 265 p.
- Sorokin Yu. P., Perevozchikov B. V. 1973, Ruby (diamond spar) from micaceous rocks of the hyperbasite massif Ray-Iz (Polar Ural). *Zapiski RMO* [Proceedings of the Russian Mineralogical Society], vol. 102, no. 6, pp. 692–696. (In Russ.)
- Bryanchaninova N. I., Makeev A. B., Zubkova N. V., Filippov V. N. 2004, Sodium-strontium mica – Na_{0.50}Sr_{0.25}Al₂(Na_{0.25–0.75}) [Al_{1.25}Si_{2.75}O₁₀](OH)₂ from the Rubinovoy Log. *Doklady Akademii nauk* [Doklady Earth Sciences], vol. 395, no. 1, pp. 101–107.
- Shcherbakova S. V. 1976, *O dvukh tipakh rubinovoy mineralizatsii v ul'traosnovnogo massiva Ray-Iz na Polyarnom Urale* [About two types of ruby mineralization in the ultrabasic massif Ray-Iz in the Polar Urals]. *Proceedings of the Geological Institute*, vol. 210, pp. 111–119.
- Shcherbakova S. V., Suturin A. N. 1990, *Geokhimiya i mineralogiya metasomatitov s rubinom (massiv Ray-Iz, Polyarnyy Ural)* [Geochemistry and mineralogy of metasomatites with ruby (Ray-Iz massif, Polar Ural)]. *Geochemical searches for gems*. Novosibirsk, pp. 167–198.
- Leake B. E., Woolley A. R., Arps C. E. S., Birch W. D., Gilbert M. C., Grice J. D., Hawthorne F. C., Kato A., Kisch H. J., Krivovichev V. G., Linthout K., Laird J., Mandarino J. A., Maresch W. V., Nickel E. H., Rock N. M. S., Schumacher J. C., Smith D. C., Stephenson N. C. N., Ungaretti L., Whittaker E. J. W., Guo Y. 1997, Nomenclature of Amphiboles: Report of the Subcommittee on Amphiboles of the International Mineralogical Association Commission on New Minerals and Mineral Names. *Canadian Mineralogist*, vol. 35, pp. 219–246. <https://doi.org/10.1180/minmag.1997.061.405.13>
- Hawthorne F. C., Oberti R., Harlow G. E., Maresch W. V., Martin R. F., Schumacher J. C., Welch M. D. 2012, Nomenclature of the amphibole supergroup. *American Mineralogist*, vol. 97, pp. 2031–2048. <https://doi.org/10.2138/am.2012.4276>
- Sun S.-S., McDonough W. F. 1989, Chemical and isotopic systematics of oceanic basalts: implication for mantle composition and processes. *Magmatism in the ocean basin*. *Geol. Soc. Sp. Publ.*, pp. 313–346. <https://doi.org/10.1144/GSL.SP.1989.042.01.19>
- Skublov S. G. 2005, *Geokhimiya redkozemel'nykh elementov v porodooobrazuyushchikh metamorficheskikh mineralakh* [Geochemistry of rare earth elements in rock-forming metamorphic minerals]. Saint Petersburg, 147 p.
- Holland T., Blundy J. 1994, Non-ideal interactions in calcic amphiboles and their bearing on amphibole-plagioclase thermometry. *Contrib. Mineral. Petrol.*, vol. 116, pp. 433–447. <https://doi.org/10.1007/bf00310910>
- Schmidt M. W. 1992, Amphibole composition in tonalite as a function of pressure: an experimental calibration of the Al-in-hornblende barometer. *Contrib. Mineral. Petrol.*, vol. 110, pp. 304–310. <https://doi.org/10.1007/BF00310745>
- Bogdanova A. R., Vakhrusheva N. V. 2018, *Geokhimiya RZE v amfibole iz plagioklazitov massiva Ray-Iz* (Polyarnyy Ural) [Geochemistry of REE in amphibole from plagioclases of the Ray-Iz massif (Polar Ural)]. *Materials of the All-Russian Scientific Conference to mark the 70th anniversary of the founding of the Ural branch of the Russian Mineralogical Society, VII Readings in memory of corresponding member of RAS, S. N. Ivanov*. Ekaterinburg, pp. 30–31.
- Bogdanova A.R., Vakhrusheva N.V. 2019, *Raspredeleniye lantanoidov v granatovykh amfibolitakh i amfibolovykh plagioklazitakh massiva Ray-Iz* (Polyarnyy Ural) [Distribution of lanthanides in garnet amphibolites and amphibole plagioclases of the Ray-Iz massif (Polar Urals)]. *Ultramafic-mafic complexes: Geology, structure, ore potential*. Conference proceedings. Irkutsk, pp. 46–52.

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Распределение лантаноидов в амфиболе и плагиоклазе из плагиоклазитов массива Рай-Из (Полярный Урал)

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Актуальность. В настоящей работе приведены результаты минералого-петрографических и геохимических исследований плагиоклазитов, локализованных среди хромитоносных ультрамафитов массива Рай-Из (Полярный Урал). Пространственная приуроченность плагиоклазитов к месторождению хромовых руд Центральному (корундсодержащие плагиоклазиты) и рудопроявлению Юго-Западному IV (бескорундовые плагиоклазиты) определяет необходимость детального исследования пород жильной серии хромитоносных разрезов.

Цель работы. Исследование минералого-петрографических характеристик, а также геохимии лантаноидов в плагиоклазе и амфиболе из плагиоклазитов массива Рай-Из (Полярный Урал).

Результаты. По результатам микрозондового анализа было установлено, что составы амфиболов в плагиоклазитах заметно варьируют, соответствуя по современной классификации кальциевых амфиболов паргаситу, эдениту, магнезиогорнблендиту и актинолиту. В амфиболе установлена зональность двух типов. Зерна амфиболов с зональностью первого типа характеризуются увеличением содержания Al_2O_3 , TiO_2 , FeO , Na_2O от центра к периферии зерен, с зональностью второго типа – напротив, снижением этих компонентов. Плагиоклаз по химическому составу соответствует олигоклазу $An_{10-30,1}$. Отмечено снижение содержания анортитовой молекулы от центра к краю зерна. Изучение характера распределения РЗЭ в породообразующих минералах позволило установить, что главным минералом-концентратором лантаноидов в плагиоклазитах является амфибол: суммарное содержание РЗЭ в минерале составляет 250–450 г/т, тогда как в плагиоклазе концентрируется 16–18 г/т лантаноидов. Установлено, что причиной существования двух типов спектров распределения лантаноидов, выявленных ранее, является соотношение суммарного количества амфиболов и плагиоклаза в породе. Проведено сравнение распределения редкоземельных элементов в кальциевых амфиболов из плагиоклазитов с распределением РЗЭ в амфиболов нюрнундуканского мафического комплекса в Северо-Западном Прибайкалье, исследования которых выполнены С. Г. Скубловым. С помощью амфибол-плагиоклазового геотермометра (Т. Holland, J. Blundy) и геобарометра (М. W. Schmidt) были определены параметры образования амфиболовых плагиоклазитов.

Выводы. Выявлен характер распределения лантаноидов в главных породообразующих минералах – плагиоклазе и амфиболе. Сравнение параметров и особенностей состава амфиболов позволило сделать вывод о прямой зависимости между температурой, суммой РЗЭ и TiO_2 в минерале.

Ключевые слова: амфибол, плагиоклаз, плагиоклазит, геохимия РЗЭ, Рай-Из, Полярный Урал.

ЛИТЕРАТУРА

- Вахрушева Н. В., Иванов К. С., Степанов А. Е., Шокальский С. П., Азанов А. Н., Хиллер В. В., Ширяев П. Б. Плагиоклазиты из хромитоносных ультрамафитов массива Рай-Из // Литосфера. 2016. № 5. С. 134–145.
- Вахрушева Н. В., Иванов К. С. Природа и возраст плагиоклазитов ультраосновного массива Рай-Из (Полярный Урал) // ДАН. 2018. Т. 480, № 1. С. 80–84. <https://doi.org/10.7868/S0869565218130169>
- Вахрушева Н. В., Ширяев П. Б., Степанов А. Е., Богданова А. Р. Петрология и хромитоносность ультраосновного массива Рай-Из (Полярный Урал). Екатеринбург: ИГГ УрО РАН, 2017. 265 с.
- Сорокин Ю. П., Перевозчиков Б. В. Рубин (алмазный шпат) из слюдитов гипербазитового массива Рай-Из (Полярный Урал) // Записки ВМО. 1973. Т. 102, № 6. С. 692–696.
- Брянчинова Н. И., Макеев А. Б., Зубкова Н. В., Филиппов В. Н. Натрий-стронциевая слюда – $Na_{0,50}Sr_{0,25}Al_2(Na_{0,25-0,75})[Al_{1,25}Si_{2,75}O_{10}](OH)_2$ из Рубинового Лога // ДАН. 2004. Т. 395, № 1. С. 101–107.
- Щербакова С. В. О двух типах рубиновой минерализации в ультраосновном массиве Рай-Из на Полярном Урале // Тр. ВСЕГЕИ. 1976. Т. 210. С. 111–119.
- Щербакова С. В., Сутурин А. Н. Геохимия и минералогия метасоматитов с рубином (массив Рай-Из, Полярный Урал) // Геохимические поиски самоцветов. Новосибирск: Наука, 1990. С. 167–198.
- Leake B. E., Woolley A. R., Arps C. E. S., Birch W. D., Gilbert M. C., Grice J. D., Hawthorne F. C., Kato A., Kisch H. J., Krivovichev V. G., Linthout K., Laird J., Mandarino J. A., Maresch W. V., Nickel E. H., Rock N. M. S., Schumacher J. C., Smith D. C., Stephenson N. C. N., Ungaretti L., Whittaker E. J. W., Guo Y. Nomenclature of Amphiboles: Report of the Subcommission of the International Mineralogical Association Commission on New Minerals and Mineral Names // Canadian Mineralogist. 1997. Vol. 35. P. 219–246. <https://doi.org/10.1180/minmag.1997.061.405.13>
- Hawthorne F. C., Obert R., Harlow G. E., Maresch W. V., Martin R. F., Schumacher J. C., Welch M. D. Nomenclature of the amphibole supergroup // American Mineralogist. 2012. Vol. 97. P. 2031–2048. <https://doi.org/10.2138/am.2012.4276>
- Sun S.-S., McDonough W. F. Chemical and isotopic systematics of oceanic basalts: implication for mantle composition and processes // Magmatism in the ocean basin // Geol. Soc. Sp. Publ., 1989. P. 313–346. <https://doi.org/10.1144/GSL.SP.1989.042.01.19>
- Скублов С. Г. Геохимия редкоземельных элементов в породообразующих метаморфических минералах. СПб.: Наука, 2005. 147 с.
- Holland T., Blundy J. Non-ideal interactions in calcic amphiboles and their bearing on amphibole-plagioclase thermometry // Contrib. Mineral. Petrol. 1994. Vol. 116. P. 433–447. <https://doi.org/10.1007/bf00310910>
- Schmidt M. W. Amphibole composition in tonalite as a function of pressure: an experimental calibration of the Al-in-hornblende barometer // Contrib. Mineral. Petrol. 1992. Vol. 110. P. 304–310. <https://doi.org/10.1007/BF00310745>
- Богданова А. Р., Вахрушева Н. В. Геохимия РЗЭ в амфиболе из плагиоклазитов массива Рай-Из (Полярный Урал) // Материалы Всерос. науч. конф., посвящ. 70-летию основания Урал. отд-ния Рос. минералог. о-ва. VII Чтения памяти чл.-корр. РАН С. Н. Иванова. Екатеринбург: ИГГ УрО РАН, 2018. С. 30–31.
- Богданова А. Р., Вахрушева Н. В. Распределение лантаноидов в гранатовых амфиболитах и амфиболовых плагиоклазитах массива Рай-Из (Полярный Урал) // Ультрамафит-мафитовые комплексы: геология, строение, рудный потенциал: материалы конф. Иркутск: Изд-во «Оттиск», 2019. С. 46–52.

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