

Biostratigraphic Zonation of the Upper Cretaceous in the East European Platform

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Abstract—Two biostratigraphic zonations, one based on macrofauna (primarily on inocerams and belemnites) and the other one based on microfossils, mostly on benthic foraminifers, are distinguished in the regional scale of the Upper Cretaceous of the East European platform. Both zonations are correlated to each other and with the Global Upper Cretaceous Scale (Olfer'ev and Alekseev, 2002).

Key words: Upper Cretaceous, zonation of macrofauna and benthic foraminifers, correlation, East European platform.

INTRODUCTION

In 2001, the Interdepartmental Stratigraphic Committee of Russia authorized the Regional Stratigraphic Scale of the Upper Cretaceous in the East European platform after its preliminary consideration at the regional conference held January 12, 2001. The scale has not been characterized so far in publications and cannot be properly used in practice of geological survey and relevant works. Accordingly, the main objective of this communication is to present in a brief manner the biostratigraphic zonations included in the scale. Characterization of superhorizons and horizons will be published elsewhere.

Ammonites are extremely rare in the Upper Cretaceous of the Russian plate, where concurrent planktonic foraminifers are of provincial character and rock sequences reveal stratigraphic ranges of calcareous nannoplankton index species, which differ from those in the Mediterranean region. Owing to these peculiarities, it is impossible to use directly the Global Upper Cretaceous Scale suggested recently (Olfer'ev and Alekseev, 2002) in practice of dating and correlation of regional stratigraphic subdivisions. That is why the regional biostratigraphic scheme appears to be necessary.

The first regional biozonation of the Upper Cretaceous in the East European platform was elaborated in 1954 (*Resheniya...*, 1955). To a considerable extent, it was based on macrofauna zonation established long ago by Arkhangel'skii (1912) in Upper Cretaceous sequences of the Volga River basin (Table 1). In the Cenomanian Stage, Arkhangel'skii distinguished the lower *Exogyra conica* and *Actinocamax primus* zones, and the *Lingula krausei* Beds above them. He corre-

lated the *Inoceramus bringniarti* Zone with the lower Turonian Substage and attributed barren Cretaceous strata to the upper Turonian. The *Inoceramus involutus* Zone was regarded as corresponding to the Emscherian (Coniacian) Stage, whereas the Senonian interval was divided into the *Inoceramus pachti*, *Pteria tenuicostata*, *Belemnitella mucronata*, and *Belemnitella lanceolata* zones.

The scheme of 1954 was based on paleontological characterization of Upper Cretaceous sequences in eastern, western, and southern areas of the Russian plate. In addition to inocerams and belemnites, it accounted for distribution of ammonites found shortly before by Mikhailov (1951) in Campanian and Maastrichtian strata of the Donetsk basin and western Ukraine. The *Neohibolites ultimus*, *Pecten asper*, *Acanthoceras rhomagense*, and *Actinocamax plenus* zones were established in the Cenomanian Stage. Distribution of inocerams was used to subdivide the Turonian-Coniacian-lower Santonian interval into the *Inoceramus labiatus*, *I. lamarcki*, *I. involutus*, and *I. cardisoides* zones (from the base upward). The *Oxytoma tenuicostata* Zone divided into the lower *Stensioeina exculta*–*Anomalina stelligera* and upper *Ataxophragmium orbignynaformis* foraminiferal subzones (?) was attributed to the upper Santonian. The lower Campanian Substage was divided into the *Gonioteuthis quadrata* and *G. mammillata* zones, and the upper one included the *Hoplitoplacenticeras coesfeldiense* and *Belemnitella mucronata* zones. The *Bostrychoceras polyplolum* and *Belemnitella langei* zones were distinguished in the lower Maastrichtian Substage, while the *Discoscaphites constrictus* corresponded to the upper Maastrichtian. The last zone was additionally divided into the *Acanthoscaphites tridens*, *Belemnitella lan-*

ceolate, *Pachydiscus neubergicus*, and *Belemnitella americana* subzones. Remarkable in the scheme was discrimination of foraminiferal subzones in the *Oxytoma tenuicostata* Zone (Table 1).

The scale of 1954 was modernized in 1959 (Naidin, 1960; *Resheniya...*, 1962), though its essence remained unchanged. In the renewed version, subdivision of Santonian, Campanian, and Maastrichtian stages was based on belemnites studied by Jeletzky (1948, 1951, 1955, 1958) and Naidin (1951, 1952, 1954, 1956). Versions of the scale published later on were supplemented with data on faunal assemblages from the Crimea, which turned out to be here more diverse within some intervals than in concurrent strata of the platform (Gerasimov *et al.*, 1962; Naidin, 1969).

At present, the former scales based on zonation of macrofauna find a limited application, because the main distribution area of Upper Cretaceous deposits is concealed, as drilling results showed, under younger sediments. Correlation and subdivision of borehole sections can be substantiated by microfossils only, and zonation of benthic foraminifers gained the prime importance for the situation we have now in the East European platform.

After insignificant corrections, biostratigraphic schemes elaborated by Naidin *et al.*, (1984a, 1984b, 1986) have been used in this work as a basis for the suggested regional zonal scale (Table 2).

GEOGRAPHIC LIMITATIONS

The zonal scale described below is applicable to southern areas of the Russian plate, where the widespread Upper Cretaceous deposits accumulated in a vast, relatively shallow basin, one in a system of sea basins, which represented the northern periphery of the Tethys. We mean primarily those areas, which lie inside the territory of the Russian Federation.

In terms of structural geology, the study region comprises a southern part of the Moscow syneclyse, the Voronezh anteclyse, the southwestern sector of the Volga-Ural anteclyse with an internal Ul'yanovsk-Saratov depression, and the northwestern margin of the circum-Caspian depression (Fig. 1).

In terminology of Naidin *et al.*, (1986), the region under consideration represents a part of the European paleobiogeographic province (EPP) corresponding to southern and temperate (partially) latitudinal zones of Europe. In other words, it may be termed as a part of low- or sub-Boreal zone. Since the Late Cretaceous climate was warmer than now, environments in the concurrent sea basin of the East European platform corresponded to those of either the subtropical or, most likely, temperate zone.

On the east, the EPP southern boundary was situated to the south of the Crimea and northern Caucasus. Inside the EPP, there can be distinguished several subprovinces separated by sublatitudinal boundaries,

which controlled distribution of belemnites, ammonites, and other invertebrates. For instance, Christensen (1990) distinguished the northerly Central Russian and southerly Central European subprovinces of the EPP. The analyzed region corresponded almost entirely to the northern part of the EPP, i.e., to the Central Russian subprovince. Marine biota that populated the region was of a rather monotonous composition in all the areas from the west to the east, and this situation is favorable for elaboration of a uniform biozonation. Despite its remoteness, the Kaliningrad area was also a part of the subprovince.

The Moscow syneclyse is a particular case however, because benthic fauna from this structure is impoverished, mostly represented by inoceramids. Foraminifers and single ammonite and belemnite remains occur here at discrete levels, while radiolarians represent the main group of nannoplankton. Accordingly, our zonal scale is valid only for the Cenomanian, Turonian, and basal Coniacian intervals of the Upper Cretaceous sequence in the Moscow syneclyse. It is clear that the Late Cretaceous biota of this syneclyse included many taxa in common with biota of the Timan-Pechora province (Marinov *et al.*, 2002), where an independent biostratigraphic scheme should be elaborated.

ZONATION OF MACROFAUNA

According to Naidin *et al.* (1986, p. 14) "it is reasonable to distinguish chrono-zones in the European paleobiogeographic province, which can be based on ammonites and inocerams for the Cenomanian, on inocerams for the Turonian and Coniacian, on inocerams and belemnites for the Santonian, and on belemnites and ammonites for the Campanian and Maastrichtian." This alternation of guide fossil groups reflects their relative importance in separate intervals of the sequence that was controlled by the global biota evolution, on the one hand, and by changes in concrete paleogeographic environments, on the other. In other words, it is historical-geological in essence. Unfortunately, the Upper Cretaceous deposits of the Russian platform are almost lacking ammonites, single finds of which are only fragmentary described in publications. Thus, the ammonite scale proper cannot be used in the study region.

Cenomanian. The *Schloenbachia varians-Turrilites costatus* Zone corresponds in range to the lower-basal middle Cenomanian and is correlative with two lower Cenomanian zones of the western Kazakhstan (*Stratigraficheskie...*, 1996). *Turrilites costatus* Lam., the last of two ammonite index species, is unknown in the platform, whereas *Schloenbachia varians* (J. Sow.) occurs quite frequently in platform deposits (Arkhangel'skii, 1912, 1922, 1926; Milanovskii, 1940; Gerasimov *et al.*, 1962; Lavrova, 1971; Glazunova, 1972; Naidin, 1974). Exactly the distribution interval of *Schloenbachia varians* defines stratigraphic range of the zone, because this form appears simultaneously with *Mantelliceras man-*

Table 1. Correlated stratigraphic schemes of Upper Cretaceous deposits in central and southeastern areas of the East European platform and in some adjacent regions

Table 1. (Contd.)

Global Upper Cretaceous Scale (Olfer'ev and Alekseev, 2002)							Biozonation of the Upper Cretaceous in the Saratov and Penza areas of the Volga River basin (Arkhangel'skii, 1912)			Unified stratigraphic schemes of Upper Cretaceous deposits in the Russian platform of 1954 and 1958						Stratigraphic scheme of Upper Cretaceous deposits in the Russian platform (Najdin, 1960; Gerasimov <i>et al.</i> , 1962)				Upper Cretaceous biozonation in the Russian platform and Crimea (Najdin, 1969)		
System	Series	Stage	Substage	Zone or Subzone	Stage	Substage	Zone	Stage	Substage	Zone	Stage	Substage	Zone	Stage	Substage	Zone	Stage	Substage	Zone or Subzone			
Cretaceous	upper	Turonian	Coniacian	Paratexanites serratomarginatus	Emscherian	?	?	?	?	?	?	?	?	?	?	?	?	?	?			
			middle	Gauthiericeras margae		Inoceramus involutus		Coniacian	upper	Inoceramus involutus	Coniacian	Inoceramus involutus		Coniacian	Inoceramus involutus	upper (Inoceramus involutus Beds)						
			lower	Peroniceras tridorsatum				lower	upper	Inoceramus wandereri [1958]	lower	Inoceramus wandereri [1958]	?	Coniacian	Inoceramus wandereri	lower (Inoceramus wandereri Beds)						
			upper	Forresteria (Harleites) petroceniensis				upper	upper													
			middle	Subprionocyclus neptuni				lower	upper													
			lower	Romaniceras deve-rianum				upper	upper													
			upper	Romaniceras orna-tissimum				lower	upper													
			middle	Romaniceras kallesi				upper	upper													
			lower	Kame-runoceras turoni-ense				lower	upper													
			upper	Mammites nodosoides				upper	upper													
			middle	Watinoceras coloradoense				lower	upper													
			lower	Pseudaspidoco- ras flexuosum				lower	upper													
			upper	Watinoceras devonense				upper	upper													
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Upper Cretaceous biozonation in the East European platform (Naidin and Kopaevich, 1977)				Upper Cretaceous biozonation in the East European paleobiogeographic province (Naidin <i>et al.</i> , 1984b, 1986)				Boreal standard of the Upper Cretaceous (Zakharov <i>et al.</i> , 1987)				Regional stratigraphic scheme of western Kazakhstan, Santonian–Maastrichtian (<i>Stratigraficheskie...</i> , 1996); Stratigraphy of the Mangyshlak Mts., Cenomanian–Coniacian (Marcinowski <i>et al.</i> , 1996)				Regional stratigraphic scheme of Upper Cretaceous deposits in the Russian platform (approved by the ISC Bureau, February 2, 2001)													
Stage		Substage		Zone or Subzone		Stage		Substage		Zone or Subzone		Zone or Subzone		Stage		Substage		Zone or Subzone		Zone or Subzone		Stage							
Coniacian		Coniacian		?		Coniacian		Coniacian		?		Volviceramus involutus		Coniacian		Coniacian		Volviceramus involutus/Volviceramus koeneni		Volviceramus involutus		Coniacian							
Turonian	lower	Inoceramus involutus		Coniacian	upper	Inoceramus involutus		Coniacian	lower	Inoceramus involutus		Volviceramus involutus	Coniacian	lower	Mytiloides incertus	Magadiceramus subquadratus		Volviceramus involutus	Coniacian	upper	Magadiceramus subquadratus	Coniacian							
		Inoceramus wandereri				Inoceramus schloenbachi				Inoceramus schloenbachi					middle	Volviceramus involutus/Volviceramus koeneni													
?		?		?		?		?		?		Inoceramus costellatus		Coniacian		Cremnoceramus crassus/Cremneceramus deformis		Volviceramus koeneni		Cremnoceramus crassus–Cremnoceramus deformis		Coniacian							
?		?		?		?		?		?		Inoceramus costellatus		lower	Mytiloides incertus	Cremnoceramus bronniarti		Volviceramus koeneni		Cremnoceramus bronniarti		Coniacian							
?		?		?		?		?		?		Inoceramus costellatus		middle		Mytiloides rotundatus		Cremnoceramus rotundatus		lower	Mytiloides scupini–Mytiloides incertus	Coniacian							
?		?		?		?		?		?		Inoceramus costellatus		upper	Mytiloides incertus	Inoceramus costellatus		Mytiloides scupini–Mytiloides incertus		Mytiloides scupini–Mytiloides incertus									
?		?		?		?		?		?		Inoceramus costellatus		lower		Inoceramus costellatus		Mytiloides scupini–Mytiloides incertus		Mytiloides scupini–Mytiloides incertus									
?		?		?		?		?		?		Inoceramus costellatus		middle	Mytiloides incertus	Inoceramus lamarcki		Inoceramus lamarcki		Mytiloides lamarcki		Coniacian							
?		?		?		?		?		?		Inoceramus costellatus		upper		Inoceramus apicalis		Inoceramus apicalis		Mytiloides lamarcki		Coniacian							
?		?		?		?		?		?		Inoceramus costellatus		lower	Mytiloides incertus	Mytiloides hercynicus		Mytiloides subhercynicus–Mytiloides hercynicus		Mytiloides subhercynicus–Mytiloides hercynicus		Coniacian							
?		?		?		?		?		?		Inoceramus costellatus		middle		Mytiloides labiatus		Mytiloides labiatus		Mytiloides labiatus		Coniacian							
?		?		?		?		?		?		Inoceramus costellatus		upper	Mytiloides incertus	Mytiloides kossmati		Mytiloides kossmati		Mytiloides kossmati		Coniacian							
?		?		?		?		?		?		Inoceramus costellatus		lower		Watinoceras amudariense/Mytiloides hattini		Watinoceras amudariense/Mytiloides hattini		Watinoceras amudariense/Mytiloides hattini		Coniacian							
?		?		?		?		?		?		Inoceramus costellatus		middle	Mytiloides incertus	Mytiloides kossmati		Mytiloides kossmati		Mytiloides kossmati		Coniacian							
?		?		?		?		?		?		Inoceramus costellatus		upper		Preactinocamax plenus triangulus		Preactinocamax plenus triangulus		Preactinocamax plenus triangulus		Coniacian							

Table 1. (Contd.)

Global Upper Cretaceous Scale (Olfer'ev and Alekseev, 2002)				Biozonation of the Upper Cretaceous in the Saratov and Penza areas of the Volga River basin (Arkhangel'skii, 1912)			Unified stratigraphic schemes of Upper Cretaceous deposits in the Russian platform of 1954 and 1958			Stratigraphic scheme of Upper Cretaceous deposits in the Russian platform (Naidin, 1960; Gerasimov <i>et al.</i> , 1962)			Upper Cretaceous biozonation in the Russian platform and Crimea (Naidin, 1969)										
System	Series	Stage	Substage	Zone or Subzone	Stage	Substage	Zone	Stage	Substage	Zone	Stage	Substage	Zone	Subzone	Stage	Substage	Zone or Subzone						
Cretaceous	upper	Cenomanian	upper	Nigericeras scotti	Cenomanian	?	?	Cenomanian	upper	Acanthoceras rhomagense Scaphites aequalis [1958] Actinocamax plenus [1954]	Cenomanian	upper	Scaphites aequalis (S) [1958]	Actinocamax plenus	Cenomanian	upper	Scaphites aequalis						
				Neocardioceras juddii																			
				Metoicoceras geslinianum																			
				Calycoceras guerangeri																			
				Alternacanthoceras jukesbrowni																			
			middle	Acanthoceras rhomagense	Turrilites acutus	Lingula krausei Beds	Cenomanian	lower	Exogyra conica	Neohibolites ultimus Pecten asper [1954] Actinocamax plenus [1958]	Cenomanian	lower	Exogyra conica	Actinocamax primus	Cenomanian	upper	Scaphites aequalis						
				Turrilites costatus																			
			lower	Mantelliceras dixoni	Mantelliceras saxbii	Exogyra conica	Cenomanian	lower							lower								
				Man-telli- ceras man- telli																			
				Sharpe- iceras schlüteri																			
				Neostlin- goceras carica- nense																			

Note: (1) In the third column, years in square brackets after the taxon name mean one of two references given in the column heading. (2) Empty cells with question marks correspond to intervals undistinguished in corresponding schemes as individual stratigraphic subdivisions. Some of them have been implicitly regarded as continuations of adjacent zones.

telli (J. Sow.) and has its evolutionary termination in the *Turrilites costatus* Subzone.

The zone is an equivalent of *Neohibolites ultimus* and *Praeactinocamax primus* beds of the belemnite scale and corresponds in range to the lower Cenomanian Substage in the Upper Cretaceous stratigraphic scale for the Russian platform of 1958. The same correlation between the indicated subdivisions is established as well in northwestern Germany (Christensen *et al.*, 1992). Except for Ostreidae, belemnites represent fossils most widespread in Cenomanian deposits of the platform, which enable the stage identification. It should be mentioned, however, that *Neohibolites ultimus* (Orb.) has not been confidently detected in the platform, whereas in the Crimea Mountains it is characteristic of the terminal lower–middle Cenomanian strata (Naidin *et al.*, 1975; Alekseev, 1989). *Neohibolites* forms represent the belemnite group that is more thermophilic than *Belemnitella* species, the Cen-

omanian *Actinocamax* forms of the Central Russian subprovince included.

Inoceramus crippsi Mant. is an inoceramid form characteristic of the zone in question, but this taxon is known only beyond limits of the region under consideration.

Acanthoceras rhomagense–*Inoceramus crippsi* Zone. Extremely rare ammonite specimens have been found in the higher interval of the middle Cenomanian Substage. Single finds of *Acanthoceras rhomagense* (Brgt.) from the Voronezh area (Semenov, 1986) imply that sedimentary sections of the Russian plate may include the synonymous zone established in western Europe. Tröger (1981) pointed out that *Inoceramus crippsi* Mant. still exists at this level. Stratigraphic range of this taxon extends up to the middle–upper Cenomanian boundary (Atabekyan, 1994).

Inoceramus pictus pictus–*Praeactinocamax plenus longus* Zone. Macrofossils are very rare in the upper

Upper Cretaceous biozonation in the East European platform (Naidin and Kopaevich, 1977)			Upper Cretaceous biozonation in the East European paleobiogeographic province (Naidin et al., 1984b, 1986)			Boreal standard of the Upper Cretaceous (Zakharov et al., 1987)			Regional stratigraphic scheme of western Kazakhstan, Santonian–Maastrichtian (<i>Stratigraficheskie...</i> , 1996); Stratigraphy of the Mangyshlak Mts., Cenomanian–Coniacian (Marcinowski et al., 1996)			Regional stratigraphic scheme of Upper Cretaceous deposits in the Russian platform (approved by the ISC Bureau, February 2, 2001)		
Stage	Substage	Zone or Subzone	Stage	Substage	Zone or Subzone	Zone or Subzone	Stage	Substage	Zone or Subzone	Zone or Subzone	Stage	Substage	Stage	
		?			?						Mytiloides hattini			
Cenomanian	upper	Scaphites aequalis	Cenomanian	upper	Sciponoceras gracile; <i>Inoceramus pictus bohemicus</i> – <i>Praeactinocamax plenus</i> Beds in Mangyshlak	<i>Praeactinocamax plenus triangulus</i>	Cenomanian	upper	<i>Neocardioceras juddii</i>	<i>Inoceramus pictus bohemicus</i>		upper		
lower	Neohibolites ultimus Parahibolites tourtiae		lower	<i>Inoceramus crippsi</i>	<i>Alternacanthooceras jukesbrownii</i>	<i>Turrilites acutus</i>	<i>Alternacanthooceras jukesbrownii</i>	middle	<i>Actinocamax plenus</i>	<i>Inoceramus pictus</i> <i>Praeactinocamax plenus longus</i>		middle	Cenomanian	
					<i>Eucalycoceras pentagonum</i>	<i>Turrilites costatus</i>	<i>Turrilites acutus</i>	middle	<i>Acanthoceras jukesbrownii</i>	<i>Acanthoceras rhomagense</i> / <i>Inoceramus crippsi</i>		lower		
					<i>Mantelliceras mantelli</i> and <i>Schloenbachia varians</i>	<i>Inoceramus crippsi</i>	<i>Turrilites costatus</i>	lower	<i>Mantelliceras dixoni</i>	<i>Turrilites costatus</i> – <i>Schloenbachia varians</i> / <i>Praeactinocamax primus primus</i> – <i>Neohibolites ultimus</i> / <i>Inoceramus crippsi</i>		lower		
							<i>Inoceramus crippsi</i>		<i>Mantelliceras mantelli</i>					

Cenomanian Substage. Single specimens of *Praeactinocamax plenus longus* Nadj. found in the East European platform imply that we may have here a subdivision analogous to the *Praeactinocamax plenus* Zone of western Kazakhstan and correlative with the *Inoceramus pictus* Zone distinguished by Wright et al. (1984) in western Europe. According to Marcinowski et al. (1996), the *plenus* Beds of Mangyshlak correspond to the *Calycoceras guerangeri* and *Metoicoceras geslinianum* zones of the ammonite scale. Consequently, the subdivision under consideration may be tentatively correlated with two lower zones of the upper Cenomanian accepted in the general scale. We should mention as well that *Praeactinocamax plenus* (Blainv.) has been found only in the *Metoicoceras geslinianum* Zone of western Europe. According to Marcinowski and his colleagues, the upper boundary of *Praeactinocamax plenus* Zone coincides with the first occurrence level of ammonites *Neocardioceras juddii* (Barr. et Guerne) in the Mangyshlak (Table 3).

Inoceramus pictus boemicus Zone. This zone corresponds to the higher upper Cenomanian interval representing terminal portion of this substage (Tröger, 1981). In the regional scheme elaborated for the East European platform, we correlated this subdivision with the *Neocardioceras juddii* Zone of the general ammonite scale, although Nadin (*Stratigraficheskie...*, 1996) assumed that *Inoceramus pictus boemicus* Leonhard represents subspecies characteristic of the entire upper Cenomanian of western Kazakhstan. Christensen et al. (1992) placed the lower boundary of the same inoceramid zone at the base of the *Metoicoceras geslinianum* Zones of the ammonite scale accepted in northwestern Germany.

Mytiloides hattini Zone. The regional scale of the Cenomanian Stage is crowned by the *Mytiloides hattini* Zone of inoceramids that has been established in the Pueblo section of the United States (Walaszczyk, 1992; Kennedy et al., 2000), where it corresponds to the *Nigericeras scotti* Zone of the ammonite standard. The *Mytiloides hattini* interval zone comprises beds

Table 2. Stratigraphic scheme of Upper Cretaceous deposits in the East European platform

General stratigraphic scale						Regional stratigraphic subdivisions						
System	Series	Stage	Substage	Lower	Upper	Substage	Stage	Substage	Lower	Upper	Substage	
Cretaceous	Campanian						Shilov					
Santonian	Upper						Severksi					
Upper	Lower	Upper	Upper	Upper	Upper	Upper	Lower	Upper	Lower	Upper	Upper	Upper
Abathomphalus mayarensis	Anapachydiscus terminus	Nannoplankton zonation (Perch-Nielsen, 1985)	Zone of planktonic foraminifers	Superhorizon, horizon	Molluscan and echinoderm regional zone or subzone	Zone of benthic foraminifers	Subzone or beds with benthic foraminifers					
Arkhangel'skiella cymbiformis	Pachydiscus epiplectus	CC26 Nephrolithus frequens	CC25 Arkhangelskiella cymbiformis	Sudzha	Beleniella junior-Neobeleniella kazimirovensis	Brotzenella praecutica-Han-zawata ekblomi	Brotzenella complanata	Bolivinoides draco draco				
Gansserina gansseri	Pseudokossmaticeras terrene	CC24 Reinhardites levius	CC23		Beleniella sumensis	Brotzenella complanata	Brotzenella complanata	Brotzenella complanata				
Globotruncana aegyptiaca		Tranolithus phacelosus	a		Beleniella lanciolata	Neoflaellina reticulata						
Globotruncanita havanensis	Nostaceras hyatti				Acanthoscaphites costatus							
Globotruncana calcarea	Didymoceras doneziatum	CC22 Quadrum trifidum	CC21 Quadrum nitidum	Tereshkino	Beleniella licheni/Micraster grimmensis	Beleniella licheni/Micraster najdini/Micraster grimmensis	Beleniella licheni/Micraster najdini/Micraster grimmensis	Angulogavelinella gracilis				
Globotruncana venicosa	Bostrioceras polyplocum	CC20 Ceratolithoides aculeus			Beleniella langei langei/Didymoceras doneziatum	Beleniella langei langei/Didymoceras doneziatum	Beleniella langei langei/Didymoceras doneziatum	Globorotalites emayensis	Brotzenella taylorensis			
Hopliplacenticeras marroni		CC19 Calculus ovalis		Maslovo	Belenioplacenticeras coesfeldiensis/Beleniella mucronata mucronata	Beleniella mucronata mucronata	Beleniella mucronata mucronata	Brotzenella montereensis	Brotzenella montereensis			
Delavarella campanensis	Placenticeras bidensatum	CC18 Aspidolithus parcus		Alekseeva	Belemnella mammillatus	Belemnella mammillatus	Belemnella mammillatus	Cibicidoides akulagayensis	Cibicidoides akulagayensis			
Dicarinella asymmetrica	Placenticeras paraplanum	CC17		Dubenk	Belemnella praecursor	Belemnella praecursor	Belemnella praecursor	Cibicidoides temirensis	Cibicidoides temirensis			
Lucianorhabdus cayeyii	Texanites gallicus	CC16	Calculus obscurum		Sphenocerasmus patootensis/Belemnella praepraecursor	Sphenocerasmus patootensis/Belemnella praepraecursor	Sphenocerasmus patootensis/Belemnella praepraecursor	Gavelinella stelligera	Gavelinella stelligera			
Reinhardites antiphorus	Reinhardites antiphorus	CC15			Tereshka	Texanites texanus/Sphenocerasmus cardiosoides/Belemnella propinqua	Texanites texanus/Sphenocerasmus cardiosoides/Belemnella propinqua	Gavelinella infrastriatonica				

Table 2. (Contd.)

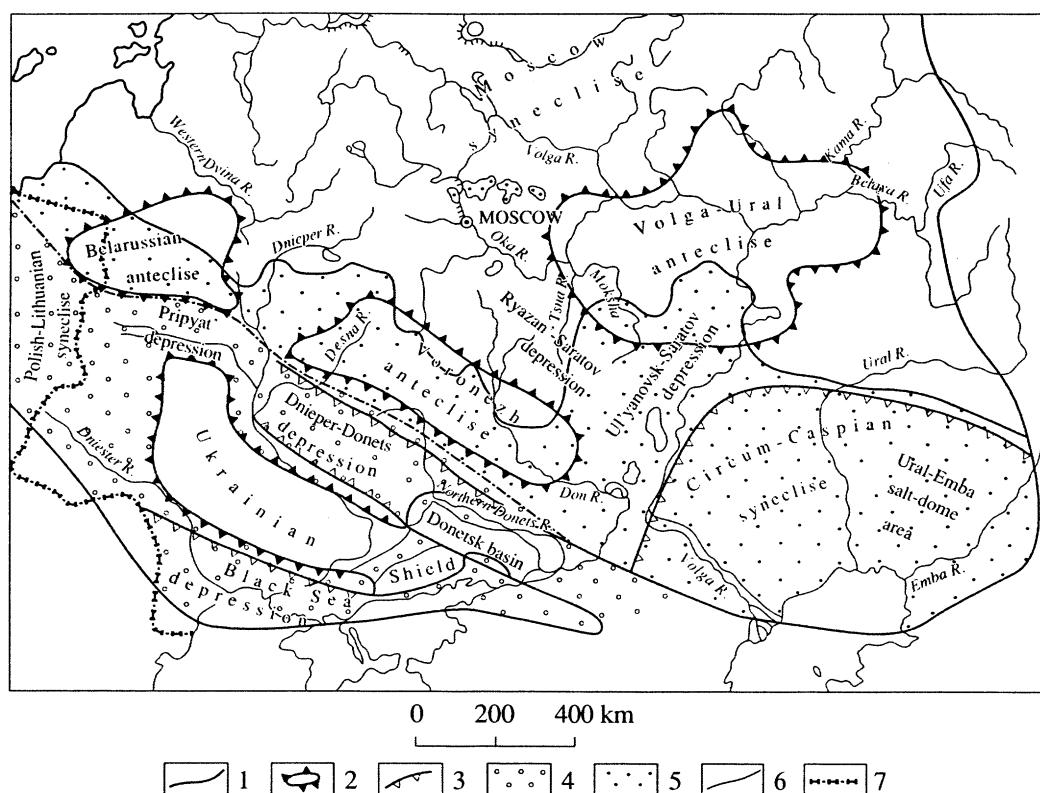


Fig. 1. Distribution areas of Upper Cretaceous deposits in different structures of the Russian platform (after Naidin *et al.*, 1986) and paleogeographic zoning of the Late Cretaceous time (after Christensen, 1990): (1) the Russian platform boundary; (2) shield or antecline; (3) synecline or depression; (4) Central European and (5) Central Russian paleobiogeographic subprovinces; (6) present-day distribution limits of Upper Cretaceous deposits; (7) state boundary of the former USSR.

between the first occurrence levels of *Mytiloides hattini* Elder and *Mytiloides puebloensis* Wal., the last of which marks the base of the overlying zone. Biozone of *Mytiloides hattini* Elder is actually much wider, because this species existed in Europe and America even during the Turonian time. This situation engendered the erroneous idea that the *Mytiloides hattini* Zone corresponds in Mangyshlak to the Turonian *Watrinoceras coloradoense* Zone of the ammonite scale (Marcinowski *et al.*, 1996) or to the basal Turonian of Poland (Walaszczuk, 1992) (Table 3). Strictly speaking, this particular zone is most likely unrecorded over the greater part of the Russian plate, where hiatus between Cenomanian and Turonian stages is usually very distinct: various horizons of the lower and middle Cenomanian are overlain here by terminal lower or middle Turonian beds. We included this subdivision in our scheme in order to retain the necessary succession. Moreover, it seems to be recognizable in some sections. For instance, the middle–upper Cenomanian Melovatka Formation of the Saratov area yields nannoplankton of the 10 zone (unpublished data of M.N. Ovechkina), and Kopaevich detected planktonic foraminifers of the *Whiteinella archaeocretacea* Zone in upper beds of that formation.

Turonian Stage. *Praeactinocamax plenus triangulus* Zone. According to Naidin *et al.* (1986), the Turonian Stage begins in the platform with the *Praeactinocamax plenus triangulus* Zone. This taxon of belemnites is rather frequent near the Turonian base in the East European platform (Tuskar' Formation of the Voronezh antecline, Oktyabr'sk and Bannovo formations of the Khoper monocline, Murom–Lomovka and Ul'yanovsk–Saratov depressions), but its stratigraphic range is imprecisely calibrated with the ammonite standard. According to stratigraphic position, the zone may correspond to the *Watrinoceras devonense* Zone of the ammonite scale.

***Mytiloides labiatus*–*Mytiloides kossmati* Zone.** In recent years, stratigraphic range of the *Mytiloides labiatus*–*Mytiloides kossmati* Zone that is terminal one in the lower Turonian Substage changed considerably. Naidin *et al.* (1984a, 1984b, 1986) regarded the upper part of the lower Turonian (provided that the stage is divided in two subunits) as corresponding to the wider *Mytiloides labiatus*–*Mytiloides subhercynicus*–*Mytiloides hercynicus* Zone. As is established at present, the *Mytiloides subhercynicus* and *M. hercynicus* zones correspond to the middle–upper Turonian *Collignoniceras woollgari* Zone of the ammonite standard (Kennedy *et al.*, 2000). Marcinowski *et al.* (1996) correlated the

Table 3. Correlation between principal stratigraphic schemes of Cenomanian–Turonian boundary deposits

Mytiloides kossmati and *M. labiatus* zones, which have been distinguished by Walaszczyk (1992) in the terminal lower Turonian, with upper strata of the *Watinoeras coloradoense* Zone coupled with *Mammites nodosoides* Zone of the general ammonite scale. As was established afterward (Kennedy *et al.*, 2000), the base of *Mytiloides kossmati* Zone is situated outside the *Watinoeras coloradoense* Zone and coincides with the base of underlying *Pseudaspidoceras flexuosum* Zone of the ammonite scale. Therefore, the *Mytiloides labiatus*–*Mytiloides kossmati* Zone of the East European platform is correlative with three upper ammonite zones of the lower Turonian in the general scale. This subdivision is established in Mordovia, Ul'yanovsk and Kuibyshev areas (Glazunova, 1972), in southwestern Lithuania and Belarus, in the southern periphery of the Donetsk basin, and in sections of the L'viv centrocline (Naidin and Morozov, 1986).

The inoceramid zonation established by Walaszczyk (1992) in central Poland and traceable in the Mangyshlak sections (Marcinowski *et al.*, 1996) has been used to determine regional zones of the middle and upper Turonian in our scale.

Mytiloides hercynicus–*Mytiloides subhercynicus* Zone. In the East European platform, the middle Turonian substage spans a succession of *Mytiloides hercynicus*–*Mytiloides subhercynicus*, *Inoceramus apicalis*, and *Inoceramus lamarcki* zones. Kennedy *et al.* (2000) established that the base of *Mytiloides subhercynicus* Zone corresponds in the Pueblo section of the United States to the first occurrence level of ammonite species *Collignoneras woollgari* (Mant.) whose appearance is suggested to be an indicator of the middle Turonian base (Bengtson, 1996). The ammonite form *Collignoneras woollgari* (Mant.) also known from Turonian deposits of the Russian plate (Ovechkina *et al.*, 2002) is extremely rare here, whereas zonal index species of inoceramids are more widespread. The latter have been encountered in sections of the L'viv centrocline and in eastern areas of the Dnieper–Donets depression (Naidin and Morozov, 1986).

Inoceramus apicalis Zone. This zone spans the middle interval of the middle Turonian. It has been established with confidence in the Volga River basin (Penza, Ul'yanovsk, and Saratov areas; Glazunova, 1972), where its thickness is insignificant. The zone has been also distinguished in sections of the L'viv centrocline and in the circum-Caspian depression (Naidin and Morozov, 1986) (Sobetskii *et al.*, 1982).

Inoceramus lamarcki Zone. This zone is terminal one in the middle Turonian Substage. Corresponding deposits are widespread and readily recognizable almost everywhere in the East European platform. Inoceramids characteristic of the zone, the index species included, have been listed by description of Turonian deposits in a series of explanatory notes to edited medium-scale geological maps.

Inoceramus costellatus Zone has been established in Germany (Ernst *et al.*, 1963), Poland (Walaszczyk, 1992), and western Kazakhstan (Marcinowski *et al.*, 1996). Its presence in the Russian plate has not been substantiated yet by paleontological data. In the East European paleobiogeographic province, Naidin *et al.* (1986) recognized the wider *Inoceramus costellatus*–*Inoceramus striatoconcentricus* Zone of the terminal Turonian interval.

Mytiloides striatoconcentricus Zone. In the East European platform, this subdivision is traceable in the Volgograd area and Ilovlya River basin (Glazunova, 1972), and also in the Saratov area and circum-Caspian depression (Sobetskii *et al.*, 1982).

Mytiloides scupini–*Mytiloides incertus* Zone crowns the upper Turonian. It is established in the Salzgitter-Salder section near Hannover in Germany above the *Mytiloides striatoconcentricus* Zone (Walaszczyk and Tröger, 1996; Kauffman *et al.*, 1996). The lower boundary of the Coniacian Stage is defined in the same section at the top of zone under consideration (Kauffman *et al.*, 1996). The *Mytiloides scupini*–*Mytiloides incertus* Zone has not been recognized yet in the Russian plate, but its presence is proved in adjacent western and eastern regions (Kauffman *et al.*, 1996; Walaszczyk, 1996; Marcinowski *et al.*, 1996; Walaszczyk, Wood, 1998).

Coniacian Stage. Like in the case of Turonian deposits, we used the inoceramid zonation to outline biostratigraphy of the Coniacian Stage in the Russian plate. In distinction from the scheme suggested by Naidin *et al.* (1986), we managed to considerably detail the stage zonation, using new data from a series of works (Walaszczyk, 1992; Marcinowski *et al.*, 1996; Kauffman *et al.*, 1996; Walaszczyk and Wood, 1998), and to verify stratigraphic range of the stage. According to the former concept, the Coniacian Stage base coincided "with the base of the *Inoceramus schloenbachi* (=*Inoceramus deformis*) Zone" (Naidin *et al.*, 1986, p. 15). Recent data (Kauffman *et al.*, 1994, 1996) showed, however, that biozone of aforementioned species corresponds in the North American Interior basin to the lower part of the *Volviceramus koeneni* Zone, thus being located at the base of the middle Coniacian Substage.

As the International Working Group recommended to authorize the first occurrence level of bivalve mollusk *Cremnoceramus rotundatus* (sensu Tröger) for the lower boundary of the Coniacian Stage, we discriminated near the stage base the synonymous zone that was originally distinguished in Germany (Ernst *et al.*, 1983; Wood *et al.*, 1984).

The *Cremnoceramus rotundatus* Zone has been established in the Salzgitter-Salder section (Kauffman *et al.*, 1996). In the East European platform, it is known in the Bryansk area of the Desna River basin and further eastward in the Mangyshlak (Marcinowski *et al.*, 1996). Recent reexamination of classical Coniacian

sequences in the Lower Saxony, central Poland, and the Western Interior basin (Walaszczyk, Wood, 1998; Walaszczyk, Cobban, 2000) showed that the species *Cremnoceramus rotundatus* (sensu Tröger non Fiege) is a junior synonym of North American *Cremnoceramus erectus* (Meek). Accordingly, it was suggested to regard this taxon as a subspecies of the *Cremnoceramus deformis* (Meek) phylogenetic lineage and to rename the *Cremnoceramus rotundatus* Zone as the *Cremnoceramus deformis erectus* Zone.

The *Cremnoceramus brongniarti* Zone established by Walaszczyk (Walaszczyk, 1992; Marcinowski *et al.*, 1996) approximately corresponds in range to the former *Cremnoceramus erectus* Zone (Ernst *et al.*, 1983; Wood *et al.*, 1984). According to recent data (Walaszczyk and Wood, 1998), it can be divided into the *Cremnoceramus waltersdorffensis hannovrensis* and *Cremnoceramus inconstans* zones. The unit under consideration has been recognized in the Desna River basin, southern Donetsk basin (Kotsyubinskii *et al.*, 1974), and Ural-Emba region (Marcinowski *et al.*, 1996).

The *Cremnoceramus crassus-Cremnoceramus deformis* Zone is terminal one in the lower Coniacian Substage. It is known in sections of the L'viv centrocline, southern Donetsk basin, and also in the west of the Voronezh anteclide, in the east of the Dnieper-Donets depression (Naidin and Morozov, 1986), and in the Mangyshlak (Marcinowski *et al.*, 1996). In opinion of Walaszczyk and Cobban (2000), the zone is an equivalent of the *Cremnoceramus crassus crassus* Zone of the terminal lower Coniacian in the United States. Walaszczyk (2000) also assumed that *Cremnoceramus* Beds do not crown the lower Coniacian Substage. After revision of fossils collected by Olfer'ev in the Bryansk area, Walaszczyk concluded that the terminal lower Coniacian subdivision corresponds to the *Inoceramus gibbosus* Zone.

After recommendations of the International Working Group, the lower-middle Coniacian boundary is defined at the appearance level of inoceramid form *Volviceramus koeneni* (Möll.) representing the index species of lower zone in the middle Coniacian Substage. The last zone completely corresponds in range to the *Peroniceras tridorsatum* Zone of the general ammonite scale (Kauffman *et al.*, 1994, 1996).

The *Volviceramus koeneni* Zone is recognizable above the appearance level of its index species and subspecies *Cremnoceramus schloenbachi schloenbachi* (Böhm) (Kauffman *et al.*, 1994). In the Russian plate, this zone is known only the L'viv centrocline (Naidin and Morozov, 1986).

The *Volviceramus involutus* Zone was widely acknowledged some time ago under the name "*Inoceramus involutus* Zone," because this index species with distinctive morphological features is easily recognizable. This subdivision was regarded formerly as corresponding in stratigraphic range either to the entire

Coniacian Stage, or to its upper half, if the stage is divided in two substages (Table 1). With its new, much narrower range approximately corresponding to the *Cremnoceramus wandereri* (Andern) Biozone (Kauffman *et al.*, 1994), the zone crowns the Coniacian Stage and can be traced in many areas of the East European platform. Hardenbol *et al.* (1998) correlated this subdivision with the *Gauthiericeras margae* Zone of the ammonite scale.

Magadiceramus subquadratus Zone. In accord with the same recommendations of the International Working Group (Kauffman *et al.*, 1996), the middle-upper Coniacian boundary is defined at the first occurrence level of inoceramid forms *Magadiceramus subquadratus* (Schlüter) representing index species of the indicated zone. In two well-known stratigraphic schemes (Marcinowski *et al.*, 1996; Hardenbol *et al.*, 1998), the *Magadiceramus subquadratus* Zone is correlated with the upper Coniacian *Paratexanites serratomarginatus* Zone of the ammonite scale. In the Russian plate, the zone under consideration has been established in western Ukraine and in the Bryansk area.

Santonian Stage. The International Working Group on Santonian Stage recommended to accept the first occurrence level of inoceramid form *Cladoceramus undulatoplicatus* (Roem.) for the Coniacian-Santonian boundary. This form is however unknown in the East European platform, where we have instead the abundant inoceramids of *cardissoides* group, such as *Sphenoceramus pachti* (Ark.) and *S. cardissoides* (Goldf.), which are, in opinion of Arkhangel'skii (1912, 1922), indicators of the Santonian base. As we already mentioned considering general biozonation of the Santonian Stage (Olfer'ev and Alekseev, 2002), the coincidence of basal levels of the *Cladoceramus undulatoplicatus* and *Sphenoceramus cardissoides* zones is doubtful. Recent investigations in northwestern Germany (Kaplan and Kennedy, 2000) and Austria (Lamolda and Hancock, 1996) showed that *cardissoides* forms appear earlier than *Cladoceramus undulatoplicatus* (Roemer). The inverted order of occurrence levels is recorded, however, in southern England (Bailey *et al.*, 1983) and in France, on the opposite side of the English Channel (Amedro and Robaszynski, 2000), and Naidin noted the same situation in western Kazakhstan (*Stratigraficheskie...*, 1996). In the Western Interior basin of North America (Kauffman *et al.*, 1994) and also in the Boreal province (Dhondt, 1992; Lopez *et al.*, 1992), Caucasus and Central Asia (Atabekyan, 1986; Moskvin, 1986), both taxa have concurrent ranges.

Sphenoceramus cardissoides-Texanites texanus-Belemnitella propinqua propinqua Zone. Following Arkhangel'skii (1912), Naidin (*Stratigraficheskie...*, 1996), Hardenbol *et al.* (1998), Tröger and Summesberger (1994), Walaszczyk (1992), and Marcinowski *et al.* (1996), we discriminate a basal Santonian interval that yields *Sphenoceramus cardissoides* (Goldf.). Other

characteristic taxa occurring in this interval are *Texanites texanus* (Roemer) known from the Terepsha Formation of the Voronezh antecline and more widespread *Belemnitella propinqua propinqua* (Mob.) (Naidin and Kopaevich, 1977; Christensen, 1990). Exactly this triad of fossils defines stratigraphic range of the lower Santonian Substage in the East European platform.

Sphenoceramus patootensis–Belemnitella precursor praepraecursor Zone. Naidin *et al.* (1986) regarded the complex *Sphenoceramus patootensis–Gonioteuthis granulata* Zone as an equivalent of the upper Santonian. We accept the first taxon for the index species of the substage, although the concurrent interval in inoceramid zonations (Tröger, 1989; Walaszczuk, 1992) is termed as the *Sphenoceramus patootensisformis* Zone or the unit 29 in nomenclature of Tröger. The fact is that the last form has stratigraphic range spanning the coupled upper Santonian and lower Campanian interval and is therefore inappropriate index species for the upper Santonian proper.

As for the second index species included by Naidin and his colleagues in the complex zone name, it seems to be inappropriate as well. Belemnites of the genus *Gonioteuthis* are characteristic of southern areas of the Central European subprovince distinguished by Christensen (1990), being extremely rare in northern areas of the Russian plate, where they have been found in single sections located on the southern slope of the Voronezh antecline. Findings of subspecies *Gonioteuthis granulata granulata* (Blainv.) are unknown so far from the study region of the Central Russian subprovince, where the Santonian–Campanian sequences yield instead the abundant representatives of the genus *Belemnitella*. According to Naidin (1964), subspecies *Belemnitella precursor praepraecursor* Najd. is most characteristic of the upper Santonian, and we used it to designate the second index taxon of the corresponding zone.

Campanian Stage. Belemnitella precursor mucronatiformis Zone. Naidin *et al.* (1986) who distinguished the *Actinocamax laevigatus–Belemnitella precursor mucronatiformis* Zone of the East European paleobiogeographic province as the lower one in the Campanian Stage regarded it as a stratigraphic equivalent of the *Gonioteuthis granulata quadrata–Placenticeras bidorsatum* Zone in Germany (Naidin, 1979). In the name of this belemnite zone, we left only the last taxon, because *Actinocamax laevigatus* Arkh. is widespread, as established now, not only in the so-called “*Pteria* Beds” of a disputable Campanian age, but also in the underlying upper Santonian deposits of the Voronezh antecline, their basal strata included. We do not exclude as well that the zone under discussion includes the topmost portion of the upper Santonian (*marsupites* Beds).

Belemnitella mucronata alpha Zone. Naidin *et al.* (1986) attributed to this zone the beds with *Gonioteuthis quadrata quadrata* Blainv. and *Belemnitella mucronata alpha* Schatsky. In the zone name, we retain

only the second taxon, because the first one is extremely rare and has been established in the *Pteria* Beds of the Boguchar and Belya Gorka localities (Naidin *et al.*, 1980, p. 34), where they belong to underlying zones of the lower Campanian and upper Santonian that is incompatible with the currently authorized stratigraphic range of *Gonioteuthis quadrata quadrata* Blainv.

Belemnellocamax mammillatus Zone. Because of the same reasons, we excluded subspecies *Gonioteuthis quadrata gracilis* (Stoll.) from the former name of the terminal lower Campanian zone and left only the index species *Belemnellocamax mammillatus* (Nilss.) whose subspecies *B. mammillatus mammillatus* (Nilss.) and especially *B. mammillatus volgensis* Najd. are widespread in eastern areas of the Russian plate. Subspecies *Gonioteuthis quadrata gracilis* (Stoll.) typical of the uppermost lower Campanian in northwestern Europe (Christensen, 1990) is known in the study region from the *Pteria* beds of the Boguchar and Belya Gorka sections (Naidin *et al.*, 1980), i.e., from the lower level of the *Belemnitella praecursor mucronatiformis* Zone.

Hoplitoplacenticeras coesfeldiense–Belemnitella mucronata mucronata Zone. For the lower zone of the upper Campanian, we accepted the name *Hoplitoplacenticeras coesfeldiense–Belemnitella mucronata mucronata* Zone suggested by Naidin *et al.* (1986), although these ammonite taxa were formerly detected only southward of the study region, in the L'viv centrocline, southern Donetsk basin, and western Kazakhstan (Naidin and Morozov, 1986; *Stratigraficheskie skhemy*, 1996). Recently, V.B. Sel'tser discovered the ammonite assemblage of *Hoplitoplacenticeras coesfeldiense coesfeldiense* (Schlüt.), *H. coesfeldiense costulosum* (Schlüt.), *H. cf. vari* (Schlüt.), and *Trachyscapites gibbus* (Schlüt.) in the Mezino-Lapshinovka section located to the northwest of Saratov. Belemnite rostra of *Belemnitella mucronata mucronata* (Schlüt.) are known from the basal upper Campanian almost everywhere in the Russian plate. The indicated zone corresponds to the lower zone of the general upper Campanian ammonite zonation.

The *Belemnitella langei* Zone originally distinguished by Naidin *et al.* (1986) represents the overlying zone of the upper Campanian Substage. Its deposits are widespread in the Voronezh antecline and Volga River basin. The unit base is defined at the first occurrence level of its index species. According to changes in phylogenetic lineage of the late Campanian *Belemnitella langei* Schatsky, Naidin distinguished the succession (from the base upward) of *Belemnitella langei minor*, *Belemnitella langei langei*, and *Belemnitella langei najdini* subzones. It was also stated that “stratigraphic ranges of subzones represent epiboles of corresponding subspecies” (Naidin and Kopaevich, 1977, p. 98).

Characteristic late Campanian ammonite taxa included in the nomenclature of above subzones enable correlation between these and concurrent West Euro-

pean biostratigraphic subdivisions. In general, the upper Campanian *Belemnitella langei* Zone of the Russian plate is subdivided (from the base upward) into the *Belemnitella langei minor*–*Bostrychoceras polyplcum*, *Belemnitella langei langei*–*Didymoceras donzianum*, and *Belemnitella langei najdini*–*Micraster grimmensis* subzones. The last index species represents the echinoderm assemblage of the Vol'sk section.

The *Belemnitella langei najdini*–*Micraster grimmensis* Subzone yields foraminifers of the *Brotzenella taylorensis* Zone (*Prakticheskoe...*, 1991) that is correlative with the *Osangularia navarroana* Zone of northwestern Germany. Consequently, the aforementioned regional subdivision of macrofauna corresponds to the lower part of the upper Campanian *Micraster grimmensis* Zone in that country (Schönfeld and Burnett, 1991; Christensen, 1993; Beniamovskii and Kopaevich, 1998).

The *Belemnella licharewi*–*Micraster grimmensis* Zone is terminal one in the upper Campanian Substage. Jeletzky (1948) originally attributed the *Belemnella licharewi* Zone to the basal lower Maastrichtian, but he denied his initial opinion later on (Jeletzky, 1958). Nevertheless, the Maastrichtian age of *licharewi* Beds was argued for since the 1950s in a series of works (Naidin, 1960, 1996; Naidin *et al.*, 1984a, 1984b, 1986; Kopaevich *et al.*, 1987). According to additional data, the *Belemnella licharewi* Zone bears foraminifers of the *Angulogavelinella gracilis* Zone (Kopaevich *et al.*, 1987) that is an analogue of the West European *Bolivinoides peterssoni* Zone (Beniamovskii and Kopaevich, 1998). Consequently, the *Belemnella licharewi* Zone should be correlated with the upper part of the *Micraster grimmensis* Zone of Germany (Schönfeld and Burnett, 1991). With due account for all these data, we discriminated the described zone under the double name in our scheme.

Maastrichtian Stage. *Acanthoscaphites tridens* Zone. The regional scale of the Upper Cretaceous of the East European platform includes two zones attributed to the Maastrichtian Stage, which correspond, according to Naidin (1978), to the *Hoploscaphites constrictus* (J. Sow.) Biozone. The lower *Acanthoscaphites tridens* Zone is divided into the lower *Belemnella lanceolate* and upper *Belemnella sumensis* subzones. The base of the former coincides with the base of synonymous belemnite zone of northwestern Europe and Baltoscandia (Christensen, 1990). Simultaneously, it approximately corresponds to the lower boundary of the ammonoid *Pseudokossmaticeras tersense* Zone and lies inside the subzone CC23a of the nannoplankton scale (Hancock *et al.*, 1993).

The *Belemnella sumensis* Subzone is an equivalent of the West European *Belemnella occidentalis* Zone (Kopaevich *et al.*, 1987). Its base is situated in the upper portion of the nannoplankton subzone CC23b (Schönfeld and Burnett, 1991; Jagt and Felder, 1999), thus being approximately at the level of lower boundary of

the *Pachydiscus epiplectus* Zone of the general ammonite scale (Hancock *et al.*, 1993).

The *Belemnitella junior*–*Neobelemnella kazimiroviensis* Zone terminates the Maastrichtian Stage of the Russian plate and corresponds to synonymous zone in the stage stratotype (Jagt, 1999). It is not easy to divide this zone in two parts, as it is evident from a thorough stratigraphic analysis by Kopaevich *et al.* (1987). The zone base approximately corresponds to the lower boundary of the nannoplankton zone NC22 or subzone CC25b (McLaughlin *et al.*, 1995) and to the base of *Anapachydiscus fresvillensis* Zone in the ammonite scale, which is believed to be the boundary between the lower and upper Maastrichtian.

ZONATION OF BENTHIC FORAMINIFERS

Benthic foraminifers are widely used in practice of subdivision of Upper Cretaceous deposits in the East European platform. Scientists who contributed much to development of the regional foraminiferal scale are V.S. Akimets, V.T. Balakhmatova, V.I. Baryshnikova, V.P. Vasilenko, A.M. Voloshina, A.A. Grigyalis, L.G. Dain, E.L. Lipnik, and E.V. Myatlyuk. The first zonal scheme for Upper Cretaceous benthic foraminifers from the Russian plate has been suggested by Grigyalis *et al.*, (1974, 1980). In fact, it was a modification of scheme elaborated by Vasilenko (1961) for the Mangyshlak. In recent years, the scheme was considerably refined by Kopaevich and Beniamovskii. The regional zonation of benthic foraminifers presented in this work was approved at the workshops in Kiev (1988) and Kamenets-Podol'skii (1989). With minor corrections, it was included afterward in "Practical Guide to Microfauna of the USSR, vol. 5: Mesozoic Foraminifers" published in 1991. In 2000, Bureau of the Regional Interdepartmental Stratigraphic Commission on Central and Southern Russian Platform organized a special Working Group that somewhat amended the scheme. We took account of these corrections in our scheme though with some modifications, bearing in mind new data on distribution ranges of index species, which have been obtained in the last two years of comprehensive study on the Upper Cretaceous biostratigraphy in the Russian plate.

Cenomanian Stage. The lower *Gavelinella ceno-*
matica Zone of the stage that has been originally dis-
tinguished by Vasilenko (1961) corresponds to the
Schlöenbachia varians–*Turrilites costatus* Zone of
macrofauna. In addition to index species, its foramin-
iferal assemblage includes *Hagenowella chapmani*
(Cushm.), *Tritaxia crenomana* Gorb., *Arenobulimina*
conoidea (Pern.), *Eggerellina crenomana* Akim., *Mars-*
sonella gomelina Akim., *Marginulina jonesi* (Reuss),
Gavelinella baltica Brotz., *G. mirtchinkii* (Akim.), *G.*
minutissima (Akim.), *Lingulogavelinella formosa*
(Brotz.), *L. orbiculata* (Kusn.), *L. ornatissima* (Lipn.),
Cibicidoides gorbenkoi (Akim.), and *Guembelitria*

cenomana (Kell.). The appearance level of these benthic foraminifers defines the base of subdivision.

The *Lingulogavelinella globosa* Zone of the upper Cenomanian has been distinguished by Akimets (1970). Its lower boundary corresponds to the appearance level of index species *Lingulogavelinella globosa* (Brotz.) in association with *Spiroplectammina cuneata* Vass., *Gaudryina arenosa* Akim., *G. angustata* *angustata* Akim., *G. folium* Akim., *Gyroidinoides concica* (Vass.), *G. nitidus* (Reuss), *Gavelinella vesca* (N. Byk.), *G. loevi* (Akim.), *Brotzenella belorussica* (Akim.), *B. berthelini* (Kell.), *Cibicides polyrraphes* *polyrraphes* (Reuss), *C. minusculus* Akim., and *Bolivina spectabilis* (Akim.). Concurrent planktonic foraminifers are represented by *Hedbergella caspia* (Vass.), *H. delrioensis* (Carsey), *Whiteinella archeocretacea* (Pess.), *W. brittonensis* (Loeb. et Tapp.), and *W. paradubia* (Sigal.). This zone spans the middle and upper Cenomanian substages, except for the *Turrlites costatus* Subzone of the general scale.

Turonian Stage. The *Gavelinella nana* Zone was distinguished by Akimets *et al.*, (1991) at the appearance level of benthic foraminifers *Arenobulimina presli* (Reuss), *Globorotalites hangensis* Vass., *G. turonicus* Kaev., *Eponides turonicus* Lipn., *Gavelinella nana* (Akim.), *G. ammonoides* (Reuss), *G. kelleri* *dorsococonvexa* Akim., and *Reussella turonica* Akim. It is correlated with three lower Turonian zones of molluscan fauna (*Praeactinocamax plenus triangulus*, *Mytiloides labiatus*—*Mytiloides kossmati*, and *Mytiloides hercynicus*—*Mytiloides subhercynicus*). Northward of Moscow, the aforementioned assemblage is known from the Chernevo Formation that yields ammonite species *Collignoniceras woollgari* (Mant.) identified by Atabekyan and nannoplankton characteristic of the zone CC11 (Ovechkina *et al.*, 2002). In the southeast of the East European platform (Bragina *et al.*, 1999), larger forms of *Hedbergella* group, such as *Hedbergella holzli* (Hagn et Zeil.), *Whiteinella paradubia* (Sigal), and *W. bornholmensis* (Dougl.) prevail among foraminifers of the zone.

The *Gavelinella moniliformis* Zone has been originally established by Akimets (1974) as corresponding to that part of the upper Turonian of the Russian plate, which begins with the molluscan *Inoceramus apicalis* Zone. The index species first appearing in the unit is associated with *Spiroplectammina praelonga* (Reuss), *Verneuilina muensteri* Reuss, *Gaudryina laevigata* Franke, *G. variabilis* Mjatl., *Arenobulimina minima* Vass., *Globorotalites multiseptus* Brotz., *Stensioeina praeeexsculpta* Kell., *S. laevigata* Akim., *Eponides concinnus* Brotz., *Osangularia whitei* *praeceps* (Brotz.), *Gavelinella ammonoides* (Reuss), *Reussella carinata* Vass., and *R. kelleri* Vass. *Gavelinella praefrasantonica* (Mjatl.) and *Ataxophragmum nautiloids* Brotz. appear in terminal strata of the zone. The listed benthic taxa may occur in association with keeled planktonic foraminifers *Marginotruncana lapparenti* (Brotz.), *M.*

marginata (Reuss), *M. pseudolinneiana* Pess., and *M. renzi* (Gand.). The indicated stratigraphic interval yields nannoplankton of the zone CC12.

Coniacian Stage. The *Gavelinella kelleri* Zone in original nomenclature of Grigyalis *et al.*, (1974) corresponds to epiboles of species *Ataxophragmum nautiloids* Brotz., *Gavelinella kelleri* Mjatl., and *G. praefrasantonica* Mjatl. Characteristic of this zone are first appearing *Gaudryina coniacica* Akim., *Ataxophragmum compactum* Brotz., *Valvularia praebiconvexa* Lipn., *Stensioeina emscherica* Baryshn. [the known in Russia is subspecies *S. granulata granulata* (Orb.)], *Eponides biconvexus* Marie, *Gavelinella giedroyci* Grig., and *Cibicidoides praeeriksdalensis* (Vass.). Other taxa typical of the zone are *Heterostomella carinata* (Franke), *Marssonella oxycone* Reuss, *Globorotalites multiseptus* Brotz., *Osangularia whitei* *whitei* (Brotz.), and *Gavelinella moniliformis* (Reuss). Rather abundant *Hedbergella agalarovae* (Vass.), *Whiteinella brittonensis* (Loeb. et Tapp.), *Marginotruncana marginata* (Reuss), *M. renzi* (Gand.), *M. coronata* (Bolli), and *Heterohelix moremani* (Cushm.) represent planktonic foraminifers of the unit. The assemblage under consideration also includes *Spiroplectammina praelonga* (Reuss), *Stensioeina praeeexsculpta* *praeeexsculpta* Kell., *Gavelinella ammonoides* (Reuss), and *Cibicides polyrraphes* *polyrraphes* (Reuss), which disappear in the overlying zone. The *Gavelinella kelleri* Zone is correlative with three lower inoceramid zones of the Coniacian Stage and correspond to the lower substage of the latter. The concurrent nannoplankton characteristic of zone CC13 suggests the same stratigraphic level.

The *Gavelinella thalmanni* Zone has been adopted at the workshop aimed to elaborate the Upper Cretaceous foraminiferal scale for the East European platform and held in Kamenets-Podol'skii (1989). The zone is named after the index species *Gavelinella thalmanni* (Brotz.) in distinction from the former *Gavelinella costulata* Zone accepted in the Practical Guide to Microfauna of the USSR (*Prakticheskoe...*, 1991). The last name appears to be inappropriate because of two reasons. First, species *Gavelinella costulata* described by Myatlyuk considerably differs from *Gavelinella costulata* (Marie) described much earlier (Marie, 1941) as a form appearing in the Campanian Stage. As a result, one can meet in publications two synonymous taxa—*Gavelinella costulata* (Mjatl. non Marie) and *G. costulata* (Marie non Mjatl.), stratigraphic ranges of which are different. Second, *Gavelinella costulata* (Mjatl. non Marie) appears, in distinction from *Gavelinella thalmanni* (Brotz.), already in the lower *Gavelinella kelleri* Zone. Characteristic benthic foraminifers appearing in the *Gavelinella thalmanni* Zone in association with its index species are *Spiroplectammina embaensis* Mjatl., *Gavelinella infrafasantonica* (Balakhm.) [*G. vombensis* (Brotz.) in nomenclature of West European paleontologists], *Cibicidoides eriksdalensis* (Brotz.), and *Bolivinitella eleyi*

(Cushm.). *Gaudryina laevigata* Franke, *Ataxophragmium compactum* Brotz., *Stensioeina emscherica* Barysch., “*Gavelinella costulata* (Mjatl.)”, *Cibicidoides praeriksdalensis* (Vass.), and *Reussella kelleri* Vass. continue their existence at this level, while evolution of *Gavelinella kelleri* (Mjatl.) and *Gavelinella moniliformis* (Reuss) is terminated here. The zone under consideration corresponds to three upper inoceramid zones of the Coniacian Stage, i.e., to the middle and upper substages of the latter. Coexisting nannoplankton of zone CC14 is another evidence in favor of this conclusion.

Santonian Stage. The *Gavelinella infrasantonica* Zone Vasilenko (1961) who originally distinguished this zone used a lame index species to term the unit despite the mass abundance of *Gavelinella infrasantonica* (Balakhm.) in the corresponding interval, because this form is known in the East European platform and northwestern Germany (Schönfeld, 1990) beginning from the middle Coniacian Substage. Other diagnostic taxa of the zone are the appearing here *Spiroplectammina rosula* (Ehr.), *Arenobulimina courta* (Marie), *Martinottiella communis* (Orb.), *Ataxophragmium crassum* (Orb.), *Neoflabellina rugosa* (Orb.), *Valvularia laevis* Brotz., *V. marie* Vass., *Stensioeina exsculpta exsculpta* (Reuss), *Eponides concinnus planus* Vass., *Osangularia whitei crassa* Vass., *O. w. polycamerata* Vass., *Cibicides ribbingi* Brotz., *Praebulimina ventricosa* (Brotz.), and *Cuneus buliminoides* (Brotz.). The zonal foraminiferal assemblage also includes *Gaudryina laevigata* Franke, *Globorotalites multiseptus* Brotz., *Osangularia whitei whitei* (Brotz.), *Cibicides excavatus* Brotz., *Cibicidoides eriksdalensis* (Brotz.), and *Bolivinitella eleyi* (Cushm.), all known from underlying strata as well. Taxa terminating their evolution in the zone are *Spiroplectammina embaensis* Mjatl., *Arenobulimina orbignyi* (Reuss), *Gavelinella costulata* (Mjatl. non Marie), *G. giedroyci* Grig., and *G. thalmanni* (Brotz.). The *Gavelinella infrasantonica* Zone is identical in range to the *Texanites texanus–Sphenoceramus cardissoides–Belemnitella propinqua propinqua* Zone of molluscan fauna.

The *Gavelinella stelligera* Zone is included in our scale in the range suggested by Akimets *et al.*, (1979). This upper zone of the Santonian Stage reveals predominance of first appearing species *Gaudryina rugosa* Orb., *Arenobulimina senonica* Mjatl., *Ataxophragmium orbignyaeformis* Mjatl., *Orbignyina variabilis* (Orb.), *Stensioeina granulata incondite* Koch, *S. granulata perfecta* Koch, *S. mursataiensis* Vass., *Gavelinella stelligera* (Marie), *G. santonica* Akim., and *Sitella carseyae* (Plumm.). This assemblage is well correlative with those characterizing the terminal middle (beginning from the *Stensioeina granulata incondite* Zone after Schönfeld, 1990) and upper Santonian of northwestern Germany (Table 4).

Foraminifers from the upper part of *Gavelinella stelligera* Zone represent an assemblage transitional in

composition to the Campanian one. All characteristic taxa listed above coexist at this level with newcomers *Heterostomella praevoeolata* Mjatl., *Arenobulimina vialovi* Wolosch., *A. obliqua* (Orb.), *Egerrellina brevis* (Orb.), *E. ovoidea* Marie, *Orbignyina convexocamerata* Wolosch., *O. inflata* (Reuss), *O. irreperata* Wolosch., *Voloschinovella aequigranensis* (Beis.), *Valvularia biconvexa* Lipn., *Stensioeina pommerana* Brotz., *Eponides aff. grodnoensis* Akim., *Osangularia cordieriana* (Orb.), *Gavelinella bistellata* (Gorb.), *G. daina* (Mjatl.), *G. costulata* (Marie), *Brotzenella insignis* (Lipn.), *Pullenia dampelae* Dain, *Cuneus triangulus* (Cushm. et Park.), *Bolivina finalis* Wolosch., and *Bolivinoides strigillatus* (Chapm.). Concurrent planktonic forms are *Rugoglobigerina ordinata* (Subb.) and *R. kelleri* (Subb.). Owing to abundance of *Stensioeina pommerana* (Brotz.) and absence of *Gavelinella clementiana* (Orb.), we defined the interval of indicated assemblage as the *Stensioeina pommerana* Beds, which are correlative with the synonymous subzone of the European paleobiogeographic province (Beniamovskii and Kopaevich, 1998) or with the *Stensioeina pommerana–Eponides frankei* Zone of northwestern Germany (Schönfeld, 1990). The last subdivision corresponds to the upper *marsupites* Beds of the terminal Santonian in Germany. In areas framing the Donetsk basin on the north, the distribution interval of the assemblage under consideration was defined as the *Gavelinella daina–Orbignyina inflata* Beds of the lower Campanian (Naidin *et al.*, 1980), which were erroneously placed above the *Gavelinella clementiana clementiana* Beds. Correlation between various zonal schemes of the Santonian Stage in the European paleobiogeographic province is illustrated in Table 4. The *Gavelinella stelligera* Zone of benthic foraminifers corresponds to the *Sphenoceramus patootensis–Belemnitella praecursor praecursor* Zone of molluscan fauna.

Campanian Stage. The *Gavelinella clementiana clementiana* Zone distinguished by Akimets (1980) can be established by appearance of its index species and is traceable throughout the territory of the European paleobiogeographic province, although its deposits are preserved after erosion in a limited number of localities. It corresponds to the molluscan *Belemnitella praecursor mucronatiformis* Zone of the East European platform, and its assemblage is also known from the lower Campanian *Placenticeras bidorsatum* Zone of Germany.

The *Cibicidoides temirensis* Zone was established for the first time by Vasilenko (1961). This upper zone of the lower Campanian is divided in two subzones. The lower *Cibicidoides temirensis* Subzone corresponds to the interval, where its index species show mass abundance. Index species of the upper *Cibicidoides aktulagayensis* Zone coexist with concurrently appearing *Neoflabellina rugosa leptodisca* (Wedek.), *Gavelinella clementiana laevigata* Marie, *Cibicidoides beaumontianus* (Orb.), *Bolivinoides decoratus* (Jones), and *B. laevigatus* Marie. These subzones are equiva-

Table 4. Correlated Santonian zonations of benthic foraminifers in the European paleobiogeographic province (EPP)

lents of two upper subzones of the *Bolivinoides decoratus* Zone that has been originally established in western Germany (Koch, 1977) and is recognizable in the European paleobiogeographic province (Beniamovskii and Kopaevich, 1998). It is impossible to trace in the Russian plate the lower *Bolivinoides decoratus* Subzone s. str. (Prakticheskoe..., 1991) or *Bolivinoides decoratus*–*Bolivinoides granulatus* Subzone (Beniamovskii and Kopaevich, 1998) originally recognized in western Kazakhstan (Naidin et al., 1984; Prakticheskoe..., 1991; Stratigraficheskie..., 1996). This situation can be explained either by a cryptic hiatus, or by ecological factors, which controlled provincialism of faunal assemblages. The *Cibicidoides temirensis* Subzone corresponds to the *Belemnitella mucronata alpha* Zone of cephalopods, whereas the *Cibicidoides aktulagayensis* Subzone is correlative with the *Belemnelloamax mammillatus* Zone.

Brotzenella montereiensis Zone. The upper Campanian Substage is divided into three zones of benthic foraminifers. The lower *Brotzenella montereiensis* was initially established by Dolitskaya (1961) in the interval, where its index species appears in association with *Arenobulimina puschi* (Reuss), *Orbignyina simplex* (Reuss), *Eponides grodnoensis* Akim., *Brotzenella menneri* (Kell.), *Cibicidoides voltzianus* (Orb.), *Sitella laevis* (Beiss.), *Eouvigerina campanica* Dain, and *Pseudouvigerina cretacea* Cushm. Taxa terminating their evolution in that interval are *Spiroplectammina lingua* Akim., *Brotzenella insignis* Lipn., and *Praebulimina ventricosa* (Brotz.). In its stratigraphic range, the *Brotzenella montereiensis* Zone corresponds to the *Hoplitoplacenticeras coesfeldiense*–*Belemnitella mucronata mucronata* of cephalopods and to the lower ammonite zone of the upper Campanian in the general scale.

The *Globorotalites emdyensis* Zone was recognized by Akimets (1974) and named after the index species described by Vasilenko in 1961 that is indistinguishable from simultaneously described *Globorotalites hiltermanni* Kaever, 1961. The zone is divided (Fig. 2) into the *Globorotalites emdyensis* s. str., *Bolivinoides draco miliaris*, and *Brotzenella taylorensis* subzones, which correspond to zones BF5, BF6, and BF7 in the scheme of Beniamovskii and Kopaevich, (1998). The lower zonal boundary is defined at the first occurrence level of the index species in association with *Heterostomella foveolata* (Marss.) and *Orbignyina pinguis* Wolosch. The base of *Bolivinoides draco miliaris* Subzone corresponds to the appearance level of this subspecies and associated taxa *Eponides frankei* Brotz., *Bolivina incrassata incrassata* (Reuss), and *B. kalinini* Vass. Indicative forms of the upper subzone are *Neoflabellina praereticulata* Hilt., *Osangularia navarroana* (Cushm.), *Stensioeina gracilis stellaria* Vass., *Brotzenella taylorensis* (Carsey), *Bolivinoides decoratus giganteus* Hilt. et Koch, and *Silicosigmoilina volganica* (Kuzn.).

The *Globorotalites emdyensis* Subzone corresponds to the *Belemnitella langei minor*–*Bostrychoceras polypicum* Subzone of cephalopods. Equivalents of *Bolivinoides draco miliaris* and *Brotzenella taylorensis* subzones are the *Belemnitella langei langei*–*Didymoceras donezianum* Subzone and lower strata of the *Belemnitella langei najdini*–*Micraster grimmensis* Subzone, respectively.

The *Angulogavelinella gracilis* Zone initially recognized by Beniamovskii et al., (1973) is terminal one in the Campanian Stage of the East European platform, though earlier (Naidin et al., 1984a, 1984b; Beniamovskii et al., 1988) its range was believed to be wider, spanning in addition the *Neoflabellina reticulata* Zone distinguished afterward (Beniamovskii and Kopaevich, 1998) and attributed to the lower Maastrichtian. In our zonation, the *Angulogavelinella gracilis* Zone is accepted as subdivision signifying appearance of its index species in association with *Spiroplectammina suturalis* Kalin., *Cuneus minutus* (Marss.), *Bolivinoides delicatulus* Cushm., and *B. peterssoni* Brotz. Its assemblage also includes *Silicosigmoilina volganica* (Kuzn.) and *S. epigona* (Rzeh.).

The *Angulogavelinella gracilis* Zone is confidently correlated (Beniamovskii and Kopaevich, 1998) with the *Bolivinoides peterssoni*–*Globorotalites hiltermanni* Zone of northwestern Germany (Schönfeld, 1990), where the latter is attributed to the terminal part of the *Micraster grimmensis*–*Cardiaster granulosus* Zone that is the uppermost one in the Campanian Stage of western Europe. In the Russian plate, the *Angulogavelinella gracilis* Zone corresponds to the upper portion of the *Belemnitella langei najdini*–*Micraster grimmensis* Subzone and to the *Belemnella licharewi*–*Micraster grimmensis* Zone (Prakticheskoe..., 1991).

Maastrichtian Stage. The *Neoflabellina reticulata* Zone established by Koch (1977) is the lowermost one in the Maastrichtian Stage. It is defined by appearance of index species in association with *Heterostomella bullata* Akim., *Cibicidoides bembix* (Marss.), *Bolivina decurrens* (Ehr.), and *Pseudouvigerina cristata* (Marss.). Other characteristic forms of zonal assemblage are *Spiroplectammina suturalis* Kal., *Orbignyina pinguis* Wolosch., *O. sacheri* (Reuss), *O. inflata* (Reuss), *Stensioeina pommerana* Brotz., *Eponides frankei* Brotz., *Osangularia navarroana* (Cushm.), *Brotzenella menneri* (Kell.), *Cibicidoides beaumontianus* (Orb.), *C. aktulagayensis* (Vass.), *C. voltzianus* (Orb.), *Cuneus minutus* (Marss.), *Sitella laevis* (Vass.), *Bolivina incrassata incrassata* Reuss, *Bolivina kalinini* (Vass.), and *Bolivinoides delicatulus* Cushm. Taxa terminating their evolution at this level are *Globorotalites emdyensis* Vass., *Gavelinella clementiana laevigata* (Marie), and *Gemellides orcinus* (Vass.). The described zone corresponds to the lower part of the *Belemnella lanceolata* Subzone and is correlative with the subzone CC23a of calcareous nannoplankton.

Fig. 2. Detailed zonation of the late Santonian–Maastrichtian benthic foraminifers (after Beniamovskii and Kopaevich, 1988, with modifications).

Table 5. Foraminiferal biozonations in the Mangyshlak–eastern circum-Caspian region (after Beniamovskii and Kopaevich, 1998, with modifications) and northwestern Germany correlated with the scheme suggested for the Russian platform

Mangyshlak and eastern circum-Caspian region		North-western Germany		Zonation of benthic foraminifers in the European paleobiogeographic province (Beniamovskii and Kopaevich, 1998)		Regional zones of benthic foraminifers		Zonation of benthic foraminifers in the Russian platform as suggested in this work	
Substages and their subordinate subdivisions				East European paleobiogeographic province		West European paleobiogeographic province			
m ₂	m ₂	m ₂ ²	m ₂ ²	Ilanzawai ekblomi/Pseudotextularia elegans	BF13	Pseudotextularia elegans	XXVI	Pseudotextularia elegans – I. Zone	(Schönenfeld, 1977)
						Gavelinella danica	XXV		(Koch, 1977)
m ₁	m ₁ ³	m ₁ ²	m ₁ ²	Bolivinoides draco draco	BF11	Bolivinoides draco draco	XXIV		
	m ₁ ²	m ₁ ¹	m ₁ ¹	Bolivinoides paleocenicus/Neoflabellina reticulata	BF12	Bronzenella complanata	XXII		
m ₁	m ₁ ¹	m ₁ ¹	m ₁ ¹	Neoflabellina reticulata/Bolivina decurrens	BF10b	Bronzenella complanata	XXII	Neoflabellina reticulata	
				Angulogavelinella gracilis/ Bolivinoides peterssoni	BF9	Angulogavelinella gracilis	XXII	peterssonihiltermanni – C.R. Zone	
cp ₂	cp ₂	cp ₂ ²	cp ₂ ²	Neoflabellina praereticulata/Bronzenella taylorensis	BF8a	Osangularia navarroana	BF8b	Bronzenella taylorensis	
				Bolivinoides decurrens/gigantus	BF7	Bolivinoides draco milianis	XXI	Bolivinoides draco milianis	
cp ₂	cp ₂	cp ₂ ³	cp ₂ ³	Bolivinoides draco milianis	BF6	Cibicidoides volzianus	XIX	Neoflabellina numismalis	
				Globorotalites hiltermanni (= G. emydensis)	BF5	Bronzenella montereensis	XVIII	leopolitana – P.R. Zone	
cp ₁	cp ₂	cp ₁ ¹	cp ₂ ¹	Bronzenella montereensis/Leopolianna leopolitana	BF4c	Cibicidoides akulagayensis	XVII	laevigatus – P.R. Zone	
				Bolivinoides decoratus	BF4b	Cibicidoides temirensis	XVI	voltaianus – P.R. Zone	
cp ₁	cp ₁ ³	cp ₁ ²	cp ₁ ²	Cibicidoides akulagayensis/C. volzianus	BF4a	Bolivinoides decoratus	XV	granulatus – P.R. Zone	
				Bolivinoides decoratus				wedekindi – P.R. Zone	
cp ₁	cp ₁ ¹	cp ₁ ¹	cp ₁ ¹	Gavelinella clementiana clementiana	BF3	Gavelinella clementiana clementiana	XVI	clementiana – P.R. Zone	
								concinna – I Zone	
		st ₃	st ₃	Stensioeina pommerana	BF2b	Bolivinoides strigillatus	XIII	ponneriana/frankei P.R. Zone	
				Gavelinella stelligera	BF2a			strigillatus – P.R. Zone	
st ₂				Stensioeina granulata perfecta	BF1	Stensioeina granulata perfecta	XII	Stensioeina granulata perfecta	

The *Brotzenella complanata* Zone was originally recognized by Grigyalis *et al.*, (1974), and its stratigraphic range was somewhat refined afterward (Akimets *et al.*, 1983). It includes the *Brotzenella complanata* s. str. and *Bolivinoides draco draco* subzones. Foraminiferal assemblage of the former is similar in general to that of the underlying zone, but newcomers *Brotzenella complanata* (Reuss) and *Bolivina incrassata crassa* Vass. represent its distinctive taxa. The lower subzone corresponds to the upper part of *Belemnella lanceolate* Subzone or to the subzone CC23b of calcareous nannoplankton. The upper subzone is established above the extinction level of subspecies *Bolivinoides draco miliaris* Hilt. et Koch that is replaced by subspecies *Bolivinoides draco draco* (Marss.) associated with *Gavelinella midwayensis* (Plum.). This subzone corresponds to the *Belemnella sumensis* Subzone and bears nannoplankton of the zone CC24.

The *Brotzenella praeacuta-Hanzawaia ekblomi* Zone is terminal one in the Maastrichtian Stage. In sections of northwestern Germany (Koch, 1977) and western Kazakhstan (Naidin *et al.*, 1984a, 1984b), there were distinguished, though under different names, two subzones of the unit (Table 5): the lower *Gavelinella danica* or *Brotzenella praeacuta* subzones and the upper *Pseudotextularia elegans* or *Hanzawaia ekblomi* subzones (in western or eastern regions, respectively). It is likely that in the eastern Russian plate (Ul'yanovsk and Saratov areas), the *Brotzenella praeacuta* and *Hanzawaia ekblomi* subzones can be recognized as separate subdivisions (Beniamovskii *et al.*, 1988; Dmitrienko *et al.*, 1988; Alekseev *et al.*, 1999), but they cannot be discriminated in the west so far.

The following taxa of benthic foraminifers are characteristic of the *Brotzenella praeacuta-Hanzawaia ekblomi* Zone: *Spiroplectammina suturalis* Kal., *S. kelleri* Dain, *S. kasanzevi* Dain, *Neoflabellina reticulata* (Reuss), *Stensioeina pommerana* Brotz., *S. caucasica* (Subb.), *Osangularia navarroana* (Cushm.), *Gavelinella affinis* (Hant.), *G. danica* (Brotz.), *G. midwayensis* (Plum.), *G. pertusa* (Marss.), *G. welleri* (Plum.), *Brotzenella complanata* (Reuss), *B. praeacuta* (Vass.), *Hanzawaia ekblomi* (Brotz.), *Cibicidoides beaumontianus* (Orb.), *C. bembix* (Marss.), *C. spiropunctatus* (Gall. et Marss.), *Anomaliodoides pinguis* (Jenn.), *A. ukrainicus* (Wolosch.), *Karreria fallax* Rzeh., *Cuneus minutus* (Marss.), *Bulimina quadrata* Plum., *B. inflata* Seg., *Bolivina decurrens* Ehr., *B. plaita* Carsey, *Bolivinoides draco draco* (Marss.), *B. decoratus giganteus* Hilt. et Koch, *B. incrassata incrassata* Reuss, and *B. peterssoni* Brotz. It is remarkable that among concurrent planktonic forms there are present *Abathomphalus mayaroensis* (Bolli) and *Globotruncanita stuarti* (Lappar.), both identified by T.E. Ulanovskaya in the Maastrichtian sequence of the Pervomaiskaya-Chir monocline, the Rostov region. The described zone yields fossils characteristic of nannoplankton zones CC25 and CC26 (Alekseev *et al.*, 1999) and corresponds to the

Belemnitella junior-*Neobelemnella kazimiroviensis* Zone of cephalopods.

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