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How Cr³⁺ and Fe³⁺ affect Mg–AI order–disorder transformation at high temperature in natural spinels

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Abstract Three natural $Mg(Al_{2-v}Cr_v)O_4$ spinels (y \sim 0.07–0.16), highly ordered in terms of Mg–Al, and one $Mg(Al_{2-\nu}Fe^{3+}_{\nu})O_4$ spinel ($\nu \sim 0.08$), highly ordered also in terms of Fe³⁺, were studied by means of X-ray singlecrystal diffraction. All samples were heated in situ from 25 to 1000 °C in order to follow both thermal expansion and evolution of the structural state of spinel with temperature. Thermal expansion was monitored by means of the variation of cell edge a with temperature, and found to be well represented throughout the temperature range by a regression line $a = a_0 (1 + \alpha \Delta T)$, slightly different at lower and higher temperatures. Thermal expansion coefficient α_l , referring to the lower temperature range (i.e. during pure thermal expansion), was slightly lower than α_2 , calculated only over the highest temperatures. The trend showed different slopes for individual crystals.

Structural evolution with temperature was studied by means of the variation of oxygen positional parameter u, which is strongly influenced by intersite cation exchange and thus closely correlated with inversion parameter x. In particular, in the three Cr samples, in which Cr resides only in the octahedral site, u parameter variations and hence the order–disorder process, started at about 700 °C. Instead, in the Fe³⁺ sample, this process was triggered at lower temperatures, starting at 550 °C with Fe³⁺–Mg exchange followed at higher temperatures by that of Mg–Al. Cr contents in the Cr samples affected the occupancy of Al in the tetrahedral site at the highest temperatures.

In both Mg–Al–Cr and Mg–Al–Fe³⁺ compositions, if $Cr\sim Fe^{3+}$, parameter *u* reached the same value only when the Mg–Al exchange was dominant, i.e. at the highest temperatures, but not before. Cation distribution at each temperature was obtained by the bond-length model, applying thermal expansion to pure bond lengths. This method is applied here to complex compositions for the first time.

Keywords In situ heating \cdot Cr³⁺ spinel \cdot Fe³⁺ spinel \cdot X-ray diffraction \cdot Order–disorder

Introduction

The temperature dependence of the cation distribution between T and M sites in spinels has been extensively studied by means of several analytical techniques and heating methods (in situ and quench) such as: NMR (Wood et al. 1986; Millard et al. 1992; Maekawa et al. 1997), ESR (Schmocker and Waldner 1976), powder neutron diffraction (Peterson et al. 1991; Harrison et al. 1999; Redfern et al. 1999), magnetic susceptibility (Harrison and Putnis 1999), Mössbauer spectroscopy (O'Neill et al. 1992; Larsson 1995; Andreozzi et al. 2001) and powder and single-crystal X-ray diffraction (Yamanaka and Takeuchi 1983; O'Neill and Dollase 1994; O'Neill et al. 1992; Della Giusta et al. 1996; Andreozzi et al. 2000; Carbonin et al. 2002). However, most of these studies were performed on synthetic samples (end members or binary joins) obtained from material quenched from high annealing temperatures, and thus characterized by high structural disorder. Comparatively, only a few of the above works regard natural spinels, which show a much lower inversion degree at room temperature, due to ordering during very slow continuous cooling on a geological time scale. Moreover, studies based on in situ heating are scarce, although such experiments are really the best way to obtain information on high-temperature intracrystalline cation distribution, not influenced by effects due to quenching processes. For these reasons, highly ordered natural spinels belonging to the series spinel-magnesiochromite and spinel-magnesioferrite, characterized by Cr^{3+} or Fe^{3+} cations substituting for

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 Al^{3+} , were studied by means of single-crystal X-ray diffraction and in situ heating. The aim of this work was to understand how Cr^{3+} and Fe^{3+} cations influence the Mg–Al order–disorder reaction with temperature.

Selected specimens were three pink Mg–Al–Cr spinels from the Olkhon metamorphic complex, Lake Baikal, Russia (Makrygina and Petrova 1998; Lavina et al. 2003), characterized by increasing chromium contents from 0.07 (L–Cr) to 0.12 (M–Cr) and 0.16 (H–Cr) atoms per four oxygens. The fourth sample (L–Fe³⁺) is a green ferrian variety (obsolete mineral name: chlorospinel) from the Shishimsk Mountains, Urals, Russia, occurring in a chlorite schist with magnetite (Bothwell and Hey 1958).

It is well known in the literature that the Cr^{3+} cation is completely ordered in the M site (Navrotsky and Kleppa 1967; Burns 1975; O'Neill and Dollase 1994), as well as that the Fe³⁺ cation often shows an essentially ordered distribution in the octahedral site in natural samples (Osborne et al. 1981; Carbonin et al. 1996; Della Giusta et al. 1996; Lucchesi and Della Giusta 1997).

On the basis of similar contents of Fe^{3+} (0.08 atoms per four oxygens) and Cr^{3+} (0.07), and the ordering of both cations in the M site, L–Fe³⁺ and L–Cr samples were compared.

Oxygen positional parameter u was chosen to monitor the evolution of the spinel structural state, as it may be measured routinely with extreme accuracy and precision. Cation distribution at each temperature was obtained by the bond-length model, applying thermal expansion to pure bond lengths. This method is applied here to complex compositions for the first time.

Experimental

Data collection

X-ray single-crystal data were collected up to $2\theta = 75^{\circ}$ (Moka radiation monochromatized by a flat graphite crystal) using a Siemens Aed II 4-circle diffractometer. For each crystal, the same set of 87 independent reflections were measured using the ω -2 θ scan mode, with profile recording for each reflection, and then used for all heating runs. Intensity data were collected from room temperature to 1000 °C in a controlled atmosphere (Ar) using a microfurnace installed on the diffractometer (Molin et al. 2001; Carbonin et al. 2002). The time for most data collections was about 2 h, including the time required for orientation matrix and cell parameter measurement (50 min). At 700 °C, on the basis of the kinetic data from the literature (Andreozzi and Princivalle 2002), the crystal was maintained at the same temperature for about 15 h before data collection, in order to approach equilibrium as closely as possible. However, it is our opinion, that until 650 °C the equilibrium distribution was obviously not achieved considering the time used for the experiments. Moreover, in order to define at which temperature intracrystalline reactions begin, several backand-forth runs were carried out, continuously monitoring at each temperature oxygen coordinate u, which is order-disorder transformation-dependent at a given temperature.

For all heating runs, intensities were corrected for spherical absorption by taking into account the mean radius of the crystal.

Unit-cell parameters were determined at each temperature by centering 24 reflections in the range $25^{\circ} < \theta < 34^{\circ}$. Each reflection was centered at both positive and negative values of 2θ and ω angles using the double-step scan routine. The mean of the ω centres was taken as the true value.

Refinements

Structural refinements were carried out in the $Fd\bar{3}m$ space group (with origin at $\bar{3}m$) with the Shelxl-97 program (Sheldrick 1997) without chemical constraints. No violations of this symmetry were detected. Refined parameters were: scale factor, secondary extinction coefficient, oxygen positional parameter u, anisotropic displacement parameters U(O), U(M) and U(T). For L–Cr³⁺ and L–Fe³⁺ samples, one scattering curve, Mg²⁺ and Al^{1.5+}, was assigned to T and M sites, respectively, not constrained to full site occupancy because of the Mg–Al inversion. For M–Cr³⁺ and H– Cr³⁺ samples, Mg¹⁺ and two scattering curves, Al^{1.5+} and Cr^{1.5+}, were assigned to T and M sites, respectively.

In spinels, the mean atomic number (m.a.n.) determined by least-squares refinement of site occupancies is quite sensitive to the ionization level of oxygen (Della Giusta et al. 1986) so, for the O-scattering curve, an ionization level between $O^{1.5-}$ and O^{2-} was chosen to obtain the best values for all conventional agreement factors.

Structural refinement results are listed in Tables 1 and 2.

Sample characterization

After data collection, chemical analyses were performed on the polished surface of the same single crystal used for X-ray study, using the Cameca/Camebax Microbeam electron microprobe at the Istituto di Geoscienze e Georisorse, CNR Padova. Analyses were performed at 15 kV and 15 nA sample current using the wavelength-dispersive method (WDS). X-ray counts were converted into oxide weight percentages using the PAP correction program supplied by CAMECA. Synthetic spinel (MgAl₂O₄) and synthetic oxide standards were used. The results of the microprobe analyses are reported in Table 3.

Inversion parameter x (Al in T) at room temperature was calculated according to the bond-length method, following Carbonin et al. (1996) and using ionic radii at room temperature from Lavina et al. (2002). This method determines cation distribution by using a soft chemical constraint. At the temperature at which cation exchange occurs, the x variation was determined with the same method as before, by applying thermal expansion coefficient α_1 (see next section) to pure bond lengths (Carbonin et al. 2002).

Results and discussion

The variation of cell edge *a* with temperature is clearly shown by two regression lines $a = a_0 (1 + \alpha \Delta T)$, slightly different at lower and higher temperatures and with different slopes for individual crystals. Thermal expansion coefficient α_1 , referring to the lower temperature range (i.e. during pure thermal expansion, preceding the variation in parameter *u*) is slightly lower than α_2 , calculated only over the highest temperatures (Fig. 1). The best-fit values of the coefficients of these equations are listed in Table 4.

Oxygen positional parameter u is greatly influenced by intersite cation exchange, which is temperaturedependent, and thus closely correlated with inversion parameter x; for this reason u is considered as a good

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T (°C)	a (Å)	п	x^{a}	$U_{eq}(O)$ (Å ²)	$U_{eq}(M) (\dot{A}^2)$	$U_{eq}(T)$ (Å ²)	e ⁻ (M)	e ⁻ (T)	e ⁻ tot	MO (Å)	TO (Å)	$R_{all}{}^{b}~\%$	wR_2^b %
L-Cr 1st run 25 200 400 450 400 400	8.0914(11) 8.1018(11) 8.1169(10) 8.1207(10) 8.1168(10) 8.1168(10) 8.0917(10)	$\begin{array}{c} 0.26344(8)\\ 0.26344(8)\\ 0.26355(8)\\ 0.26337(9)\\ 0.26348(8)\\ 0.26327(8)\\ 0.26346(7)\\ \end{array}$	0.121 0.113 0.126 0.118 0.118 0.134	$\begin{array}{c} 0.0046(2)\\ 0.0061(3)\\ 0.0082(3)\\ 0.0091(3)\\ 0.0083(3)\\ 0.0043(2)\\ \end{array}$	$\begin{array}{c} 0.0042(2)\\ 0.0058(2)\\ 0.0083(2)\\ 0.0083(2)\\ 0.0088(2)\\ 0.0083(2)\\ 0.0043(2)\\ \end{array}$	0.0050(3) 0.0070(3) 0.0099(4) 0.0109(3) 0.0096(3) 0.0052(3)	13.51(5) 13.58(6) 13.67(7) 13.62(5) 13.62(5) 13.62(5)	12.29(7) 12.36(8) 12.36(8) 12.36(8) 12.36(8) 12.34(7) 12.33(7) 12.38(6)	39.30(12) 39.51(14) 39.71(16) 39.65(13) 39.65(13) 39.61(12)	1.920(1) 1.922(1) 1.927(1) 1.927(1) 1.927(1) 1.927(1)	1.940(1) 1.944(1) 1.945(1) 1.945(1) 1.948(1) 1.944(1) 1.941(1)	2.01 2.47 2.72 2.72 2.01	2.96 2.93 2.59 2.59 2.38 2.38
2nd run 500 450	8.1256(8) 8.1217(8)	0.26329(10) 0.26337(9)	0.132 0.126	0.0102(4) 0.0093(3)	$\begin{array}{c} 0.0094(2) \\ 0.0090(2) \end{array}$	0.0113(4) 0.0105(4)	13.54(7) 13.57(6)	12.26(9) 12.24(8)	39.34(16) 39.39(15)	1.929(1) 1.928(1)	1.946(1) 1.946(1)	3.42 2.71	3.04 2.70
3rd run 550 600 650 700 800 900 1000	8.1298(9) 8.1335(10) 8.1335(10) 8.1381(7) 8.1423(10) 8.1503(11) 8.1577(11) 8.1667(14)	0.26346(8) 0.26342(9) 0.26333(9) 0.26165(11) 0.26132(12) 0.26118(15)	$\begin{array}{c} 0.119\\ 0.122\\ 0.122\\ 0.253\\ 0.283\\ 0.293\end{array}$	0.0103(3) 0.0111(3) 0.0124(3) 0.0164(4) 0.0183(4) 0.0180(4) 0.0207(5)	0.0101(2) 0.0106(2) 0.0113(2) 0.0113(2) 0.0128(3) 0.0147(3) 0.0159(3) 0.0174(4)	0.0121(3) 0.0135(4) 0.0135(4) 0.0137(4) 0.0147(4) 0.0168(5) 0.0184(6)	13.59(6) 13.68(6) 13.68(6) 13.52(5) 13.43(6) 13.47(7) 13.65(6) 13.67(7)	12.31(7) 12.46(7) 12.25(8) 12.25(8) 12.38(8) 12.33(8) 12.57(12)	39.48(13) 39.82(14) 39.29(14) 39.23(15) 39.26(15) 39.79(16) 39.92(19)	1.929(1) 1.930(1) 1.932(1) 1.942(1) 1.947(1) 1.951(1) 1.955(1)	1.950(1) 1.950(1) 1.950(1) 1.950(1) 1.933(2) 1.929(2) 1.926(2) 1.926(2)	3.16 3.24 3.34 4.28 5.08	2.55 3.16 3.308 3.59 4.58
M-Cr 1st run 28 200 400 450 400 28	8.1019(10) 8.1116(10) 8.1263(10) 8.1305(7) 8.1265(9) 8.1018(10)	$\begin{array}{c} 0.26346(10)\\ 0.26343(11)\\ 0.26344(10)\\ 0.26336(14)\\ 0.26330(10)\\ 0.26330(10)\\ 0.26348(10)\\ \end{array}$	$\begin{array}{c} 0.114\\ 0.117\\ 0.116\\ 0.122\\ 0.127\\ 0.113\end{array}$	0.0057(5) 0.0076(5) 0.0097(5) 0.0101(6) 0.0101(5) 0.0052(4)	0.0055(3) 0.0068(3) 0.0088(2) 0.0088(2) 0.0087(3) 0.0089(3)	0.0051(4) 0.0073(4) 0.0098(4) 0.0111(6) 0.0099(4) 0.0054(4)	13.69(58) 13.62(62) 13.72(54) 13.81(78) 13.81(78) 13.83(56)	$\begin{array}{c} 12.13(10)\\ 12.13(11)\\ 12.23(9)\\ 12.23(9)\\ 12.36(13)\\ 12.14(9)\\ 12.31(10)\end{array}$	39.51(1.16) 39.37(1.24) 39.66(1.08) 39.97(1.56) 39.38(1.15) 39.98(1.12)	1.923(1) 1.925(1) 1.929(1) 1.930(1) 1.930(1) 1.930(1)	1.943(1) 1.945(2) 1.949(1) 1.948(2) 1.948(2) 1.947(1) 1.943(1)	4.19 4.21 4.54 3.71 3.87	3.02 3.15 3.75 3.75 3.00
2nd run 500 450	8.1338(9) 8.1302(8)	0.26343(10) 0.26334(12)	0.117 0.124	$\begin{array}{c} 0.0108(5) \\ 0.0108(6) \end{array}$	$\begin{array}{c} 0.0100(3) \\ 0.0094(3) \end{array}$	0.0113(4) 0.0104(5)	13.61(58) 13.71(68)	12.15(9) 12.16(11)	39.36(1.17) 39.58(1.36)	1.930(1) 1.930(1)	1.950(1) 1.948(2)	3.97 4.43	2.82 3.26
3rd run 550 600 650 700 800 900 1000	8.1379(9) 8.1419(11) 8.1461(9) 8.1502(9) 8.1502(9) 8.1577(8) 8.1538(12) 8.1733(4)	0.26339(12) 0.26331(14) 0.26308(12) 0.26187(15) 0.26187(16) 0.26187(16) 0.26144(12)	$\begin{array}{c} 0.120\\ 0.126\\ 0.143\\ 0.202\\ 0.236\\ 0.236\\ 0.267\end{array}$	$\begin{array}{c} 0.0114(6)\\ 0.0135(6)\\ 0.0145(6)\\ 0.0179(7)\\ 0.0178(7)\\ 0.0178(7)\\ 0.0209(6)\\ 0.0212(7) \end{array}$	$\begin{array}{c} 0.0103(3)\\ 0.0112(3)\\ 0.0117(3)\\ 0.0117(3)\\ 0.0131(4)\\ 0.0146(4)\\ 0.0158(4)\\ 0.0170(4)\end{array}$	0.0121(5) 0.0126(6) 0.0135(5) 0.0135(6) 0.0135(6) 0.0144(7) 0.0167(6) 0.0182(6)	13.81(70) 13.58(78) 13.58(78) 13.43(71) 13.40(85) 13.40(85) 13.68(94) 13.42(77) 13.71(78)	$\begin{array}{c} 12.33(11)\\ 12.05(12)\\ 12.05(11)\\ 11.07(11)\\ 111.92(13)\\ 112.21(13)\\ 12.25(10)\\ 12.61(11)\end{array}$	39.94(1.40) 39.21(1.57) 38.94(1.43) 38.72(1.71) 39.58(1.89) 39.09(1.53) 40.03(1.57)	1.932(1) 1.933(1) 1.936(1) 1.942(1) 1.942(1) 1.952(1) 1.952(1)	1.951(2) 1.950(2) 1.948(2) 1.948(2) 1.938(2) 1.934(2) 1.931(2) 1.932(2)	4.70 5.15 5.80 5.65 5.81 5.81	3.43 4.13 4.57 4.64 3.49 3.11
H–Cr 1st run 25 200 400 450	8.1080(12) 8.1180(11) 8.1316(10) 8.1354(11)	$\begin{array}{c} 0.26338(10)\\ 0.26331(12)\\ 0.26343(11)\\ 0.26343(11)\end{array}$	$\begin{array}{c} 0.114 \\ 0.120 \\ 0.107 \\ 0.111 \end{array}$	0.0049(4) 0.0065(4) 0.0088(5) 0.0095(5)	$\begin{array}{c} 0.0048(2)\\ 0.0063(3)\\ 0.0082(3)\\ 0.0089(3)\end{array}$	$\begin{array}{c} 0.0046(4) \\ 0.0068(4) \\ 0.0098(4) \\ 0.0109(5) \end{array}$	14.07(59) 14.09(69) 14.12(60) 13.97(62)	12.27(10) 12.35(12) 12.45(10) 12.40(11)	40.40(1.18) 40.53(1.38) 40.68(1.21) 40.34(1.25)	1.925(1) 1.927(1) 1.929(1) 1.931(1)	1.943(1) 1.945(2) 1.950(2) 1.951(2)	2.98 2.73 4.05 4.32	3.41 4.42 3.34 3.34

Table 1 ((Contd.)												
T (°C)	<i>a</i> (Å)	п	$\chi^{\rm a}$	$U_{eq}(O)\;({\rm \AA}^2)$	$U_{eq}(M)$ (Å ²) 1	$J_{eq}(T) (\dot{A}^2) e$; ⁻ (M)	e ⁻ (T)	e ⁻ tot	MO (Å)	TO (Å)	$R_{all}{}^b$ %	$wR_2^{b} \%$
400 25	8.1315(14) 8.1085(8)	0.26348(11) 0.26327(10)	$\begin{array}{c} 0.107\\ 0.123\end{array}$	0.0085(5) 0.0056(4)	0.0079(3) (0.0045(2) ().0099(4) 1 1.0050(4) 1	(4.01(61) 3.86(56)	12.42(10) 12.18(10)	40.44(1.22) 39.90(1.12)	1.929(1) 1.925(1)	1.950(2) 1.942(1)	3.85 2.99	3.39 3.36
2nd run 500 450	8.1400(14) 8.1358(13)	0.26346(12) 0.26358(10)	$0.109 \\ 0.099$	0.0105(5) 0.0098(5)	0.0093(3) (0.0088(2) (.0114(5) 1 .0105(4) 1	13.86(67) 13.92(55)	12.30(11) 12.31(9)	40.02(1.34) 40.14(1.11)	1.932(1) 1.930(1)	1.952(2) 1.953(1)	3.94 3.93	3.78 2.84
3rd run 550 600 650 700 800 900 1000	8.1445(6) 8.1482(8) 8.1525(6) 8.1559(7) 8.1569(7) 8.1655(5) 8.1726(11) 8.1791(9)	0.26355(10) 0.26342(11) 0.26318(12) 0.26318(12) 0.26242(12) 0.26197(11) 0.26197(11) 0.26171(13)	$\begin{array}{c} 0.101\\ 0.111\\ 0.129\\ 0.187\\ 0.204\\ 0.222\\ 0.242\end{array}$	0.0094(5) 0.0123(5) 0.0130(5) 0.0149(5) 0.0166(7) 0.0187(6) 0.0205(7)	0.0101(3) 0.0107(3) 0.0116(3) 0.0127(3) 0.0127(3) 0.0154(3) 0.0154(3) 0.0168(4) 0.0168(4)	0.0124(4) 1 0.0134(5) 1 0.0142(5) 1 0.0138(5) 1 0.0138(5) 1 0.0162(5) 1 0.0170(5) 1 0.0187(6) 1	4.42(59) (3.95(62) (3.99(67) 4.11(67) 4.09(74) 4.02(63) 4.03(76)	12.69(10) 12.36(10) 12.40(11) 12.43(11) 12.65(12) 12.57(10) 12.62(12)	$\begin{array}{c} 41.52(1.19)\\ 40.26(1.24)\\ 40.39(1.34)\\ 40.64(1.35)\\ 40.83(1.49)\\ 40.60(1.26)\\ 40.67(1.53)\end{array}$	1.932(1) 1.934(1) 1.937(1) 1.943(1) 1.947(1) 1.950(1) 1.954(1)	1.954(1) 1.953(2) 1.951(2) 1.941(2) 1.941(2) 1.939(2) 1.937(2)	3.52 4.22 4.57 5.19 5.19 6.36	2.94 3.11 3.37 3.37 3.27 3.27 3.27 3.27
${}^{a}x$ invers ${}^{b}R_{all}$ and	ion degree 1 wR2 agreeme	int factors acco	rding to S	HELXL97 pro	gram package								
T (°C) T	structural rehr a (Å)	lement results n u	n sample I x^{a}	U_{eq} (O) (Å ²)	ted standard (U_{eq} (M) (Å ²	leviation $m b_{i}$) U_{eq} (T) (Å	rackets). Nu x^2 $e^{-}(M)$	e ⁻ (T)	equivalent ret e ⁻ tot	MO (Å)	or all runs TO (Å)	R_{all}^{b} (%)	wR ₂ ^b (%)
L-Fe ³⁺ 1st run 25	8.0989(10)	0.26326(7)	0.141	0.0047(2)	0.0043(2)	0.0048(2)	13.62(5) 12.67(6)	39.90(11)	1.923(1)	1.926(1)	1.94	2.51
200 400	8.1239(6) 8.1239(6)	0.26329(8) 0.26333(8)	0.139 0.136	0.0062(3) 0.0088(3)	0.0059(2) 0.0081(2)	0.0070(3) 0.0098(3)	13.63(5)	(1) 12.67(7) (12.76(7))	39.93(12) 40.06(12)	1.923(1) 1.929(1)	1.926(1) 1.946(1)	2.30	3.13 2.73
450 25 25	8.1281(7) 8.1243(6) 8.0987(6)	$\begin{array}{c} 0.26323(9) \\ 0.26336(7) \\ 0.26325(8) \end{array}$	$0.142 \\ 0.132 \\ 0.138$	$\begin{array}{c} 0.0099(3) \\ 0.0096(3) \\ 0.0049(2) \end{array}$	$\begin{array}{c} 0.0090(2) \\ 0.0081(2) \\ 0.0045(2) \end{array}$	$\begin{array}{c} 0.0103(3)\\ 0.0094(3)\\ 0.0044(3)\end{array}$	$13.53(5) \\13.42(5) \\13.55(5) \\13.5$	$\begin{array}{cccc} 12.57 (8) \\ 12.42(7) \\ 12.48(7) \\ 12.48(7) \end{array}$	$\begin{array}{ccc} 39.64(13) \\ 39.27(11) \\ 39.58(13) \end{array}$	1.930(1) 1.939(1) 1.923(1)	1.946(1) 1.947(1) 1.939(1)	2.42 2.38 2.30	3.212 2.65 3.05
2nd run 500 450	8.1318(5) 8.1280(6)	0.26331(8) 0.26337(8)	$0.138 \\ 0.132$	$\begin{array}{c} 0.0104(3) \\ 0.0100(3) \end{array}$	0.0096(2) 0.0090(2)	0.0110(3) 0.0106(3)	13.51(5 13.57(5) 12.58(7)) 12.63(7)	39.61(12) 39.78(12)	1.931(1) 1.929(1)	1.948(1) 1.948(1)	2.13 2.17	2.70 3.02
3rd run 550 600 650 700 800 900 1000	8.1359(6) 8.1403(7) 8.1440(7) 8.1449(7) 8.1547(8) 8.1547(8) 8.1627(9) 8.1701(10)	0.26327(10) 0.26292(8) 0.26250(9) 0.26190(10) 0.26159(10) 0.26133(9) 0.26133(9)	$\begin{array}{c} 0.146\\ 0.172\\ 0.210\\ 0.258\\ 0.278\\ 0.297\\ 0.309\end{array}$	0.0112(4) 0.0134(3) 0.0153(4) 0.0167(4) 0.0186(4) 0.0199(4) 0.0219(5)	0.0103(3) 0.0112(2) 0.0118(2) 0.0127(3) 0.0143(2) 0.0143(2) 0.0160(3)	0.0118(4) 0.0121(3) 0.0128(4) 0.0128(4) 0.0139(4) 0.0148(4) 0.0163(4) 0.0163(4)	13.62(7 13.38(5 13.38(5 13.27(5 13.21(6 13.32(5 13.33(6)) 12.80(9) 12.80(9) 12.67(6) 12.78(7) 12.78(7) 12.88(7) 13.01(7) 13.04(9)	40.05(17) 39.43(12) 39.32(13) 39.34(14) 39.34(14) 39.64(13) 39.64(13)	1.932(1) 1.936(1) 1.940(1) 1.945(1) 1.945(1) 1.949(1) 1.953(1)	1.948(1) 1.945(1) 1.940(1) 1.932(1) 1.929(1) 1.927(1) 1.927(2)	3.10 2.86 3.28 3.46 3.46 4.26	3.05 2.53 3.26 3.26 3.63

 a x inversion degree (Al + Fe $^3^+)$ b $R_{\rm all}$ and wR2 agreement factors according to SHELXL97 program package

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Table 3 Chemical compositionby electron microprobe. Average of 20 to 33 analyses. (estimated standard deviation *in brackets*)

Sample	L–Cr	M–Cr	H–Cr	L-Fe ³⁺	
MgO	28.04 (28)	27.85 (37)	27.71 (36)	27.93 (21)	
Al ₂ O ₃	68.14 (56)	65.89 (64)	64.59 (95)	68.06 (42)	
Cr_2O_3	3.51 (14)	6.49 (23)	8.50 (75)	- ``	
MnO	0.03 (3)	0.09 (4)	0.09 (3)	0.07 (5)	
ZnO	0.10(7)	0.41(7)	0.37(7)	0.20(8)	
FeO	-	- ()	-	3.93 (12)	
Σ	99.82	100.73	101.26	100.19	
Cations on ba	sis of four oxygens				
Mg	1.003 (5)	0.999 (9)	0.995 (9)	0.997 (6)	
Al	1.928 (6)	1.868 (13)	1.834 (19)	1.920 (6)	
Cr ³⁺	0.066 (3)	0.123 (4)	0.162 (14)	_	
Mn	0.001(1)	0.002(1)	0.002(1)	0.001(1)	
Zn	0.002(1)	0.007(1)	0.007(1)	0.003(1)	
Fe ³⁺	_	_	_	0.079(3)	
Σ	3.000	2.999	3.000	3.000	

indicator for monitoring the structural state evolution of the spinel as a function of temperature.

In order to identify the temperature at which the cation exchange begins (transition zone), the three crystals with different Cr^{3+} contents and the L-Fe³⁺ sample were subjected to several back-and-forth runs at lower temperatures (from room temperature up to 500 °C). During these heating runs, the oxygen coordinate remained constant and thermal expansion was completely reversible, because no changes in the degree of order were evident.

Figure 2 shows variations in oxygen positional parameter u as a function of temperature for the three Cr^{3+} samples. The u value clearly remains constant for the three samples up to 600 °C, independently of Cr^{3+} contents. From this temperature onwards, the disordering pathways show two different slopes, between



Fig. 1 Cell edge vs. temperature. In each sample, two regression lines before and after transition zone are calculated with slightly different α_1 and α_2 thermal expansion coefficients

600 and 700 °C and 700 and 1000 °C, in all samples. From 600 to 650 °C, the samples slowly begin to disorder, M–Cr and H–Cr slightly more than L–Cr. Instead, from 650 to 700 °C, *u* suddenly drop in all samples, and at 700°C sample L–Cr reverses the above trend, reaching the lowest *u* value. Parameter *u* then decreases continuously reaching, along different pathways, quite different values at 1000 °C: 0.2612(1), 0.2615(1) and 0.2617(1), with a degree of inversion *x* of 0.29, 0.27 and 0.24 in L–Cr, M–Cr and H–Cr respectively. The degree of disorder reached at the highest temperature is inversely correlated with Cr³⁺ contents, which therefore limit Mg–Al exchange, due to the preference of Cr³⁺ for the octahedral site.

On the basis of different kinds of non-aluminium trivalent cations (Cr^{3+} vs Fe^{3+}) but with similar contents ($Cr^{3+} = 0.07$; $Fe^{3+} = 0.08$ atoms per four oxygens), it was possible to compare L–Cr with L–Fe³⁺. Both samples are highly ordered in terms of Mg–Al and, in addition, L–Fe³⁺ shows Fe³⁺ cation mainly ordered in M site (Lucchesi and Della Giusta 1997).

Figure 3 shows variations in oxygen positional parameter *u*, as a function of temperature. It is evident how the transition zone is different between the two samples: in L–Fe³⁺ the exchange reaction starts between 550 and 600 °C, and in L–Cr³⁺ at 650 °C. From these temperatures, both the *u* values show abrupt changes up to 700 °C, whereas from 800 to 1000 °C they show gentle changes, following the same disordering pathways and finally reaching the same values [0.2612(1)]. The same trend is shown by inversion parameter *x*, with the greatest differences between the two samples in the range 550–700 °C (e.g. at 650 °C x = 0.210 in L–Fe³⁺ and 0.129 in L–Cr) and only slight between 800 and 1000 °C (e.g. at 1000 °C x = 0.309 in L–Fe³⁺ and 0.293 in L–Cr). This difference in *x* between 550 and 700 °C is due to the different electron density in M and T sites.

Figure 4 shows qualitatively how the distribution of electrons between M and T sites ($(2eM-eT)/e_{Tot}$) changes in L–Fe³⁺ from 550 to 700 °C and remains constant from 700 to 1000 °C; instead, in L–Cr there are no sig-

Table 4 Best-fit values of thermal expansion coefficient in linear regression $a = a_0$ $(1 + \alpha \Delta T) \alpha_1$: lower temperature range; α_2 : highest temperatures

Sample	T range (°C)	a_0 (Å)	$\alpha_1 \times 10^{-6} \ (^{\circ}C^{-1})$	$\alpha_2 \times 10^{-6} (^{\circ}\text{C}^{-1})$	\mathbb{R}^2
L–Cr	25-650	8.0906 (7)	9.0 (2)		0.995
	700-1000	8.0880 (2)		9.9 (3)	0.998
M–Cr	25-600	8.1011 (5)	8.6 (2)		0.997
	650-1000	8.0977 (3)		9.57 (5)	1.000
H–Cr	25-600	8.1074 (8)	8.5 (2)		0.994
	650-1000	8.1050 (3)		9.6 (4)	0.994
L–Fe ³⁺	25-550	8.0980 (5)	8.7 (2)		0.997
	600-1000	8.0970 (1)		9.2 (2)	0.999

nificant changes at any temperature. Actually, according to detectable or undetectable differences in electrons in the two samples, Fe^{3+} cations migrate from M to T site; alternatively, the exchange mainly involves Mg and Al. Table 5 shows how, in L–Fe³⁺ in the T site, Mg starts exchanging with Fe³⁺ between 500 and 550 °C, while Al remains constant up to 550 °C; between 550 and 600 °C Mg exchanges only with Al, this exchange also remaining dominant from 600 up to 700 °C with respect to that of Fe³⁺–Mg. From 700 to 1000 °C, the only detectable exchange is Mg–Al. Table 1 shows that in L–Cr, Al remains constant up to 650 °C in the T site and then significantly increases from 700 to 1000 °C.

Figure 5 shows Fe^{3+} mainly ordered in the M site (80%) from room temperature up to 500 °C, but from 550 to 700 °C Fe^{3+} it disorders in the T site up to ~ 50%, and then, from 700 to 1000 °C, it reorders slightly in the M site. The same behaviour is confirmed by recent studies performed on synthetic spinel–magnesioferrite series by Mössbauer spectroscopy (Andreozzi et al. 2001), showing Fe^{3+} partitioning as a function of magnesioferrite components and temperature. Unfortunately, these synthetic samples have a final

equilibrium temperature estimated to be around 800 °C, and therefore Fe^{3+} behaviour is comparable with that found in L–Fe³⁺ only at the highest temperatures (800–1000 °C), with a slight preference of the Fe³⁺ cation for the M site.

The behaviour of Fe^{3+} , in spite of its low amount in the sample, supports the idea that at the lowest temperature, our data are consistent with a two-stage kinetic process in which relatively rapid exchange of Fe^{3+} with Mg between tetrahedral and octahedral sites is followed by slower exchange of Mg with Al, as observed by Harrison et al. (1999). For this purpose, in situ heating of the disordering kinetics at different isotherms can highlight such a two-stage process, and study is in progress on this natural sample in the range 500–700 °C. It will be interesting to compare results with those of the equivalent Cr sample.

Summary



In the three Cr samples, oxygen positional parameter u and inversion parameter x values as a function of tem-

Fig. 2 Oxygen positional parameter u as a function of temperature in samples L–Cr, M–Cr and H–Cr. *Inset:* Inversion parameter x vs. temperature



Fig. 3 Oxygen positional parameter *u* as a function of temperature in samples $L-Fe^{3+}$ and L-Cr. *Inset*: Inversion parameter *x* vs. temperature



Fig. 4 Distribution of electrons between M and T sites obtained from structural refinements in samples $L-Fe^{3+}$ and L-Cr, as a function of temperature, for forward runs only

perature are related to Cr content, which therefore affects Mg–Al exchange. In sample L–Cr, u reaches the lowest and x the highest values, opposite to what occurs in sample H–Cr.

In sample L–Fe³⁺, u reaches the same value as in sample L–Cr at the highest temperatures, showing the same disordering pathways from 800 to 1000 °C, and therefore similar values of x, whereas those between 550 and 700 °C are significantly different. In this region, inversion parameter x (Al³⁺ + Fe³⁺) in L–Fe³⁺ is due to rapid Fe³⁺–Mg exchange between M and T sites, and



Fig. 5 Ferric iron distribution (%) as a function of temperature in sample L-Fe³⁺, obtained by cation partitioning (Table 5). *Solid and open squares* Relative proportions of Fe³⁺ in T and M sites respectively, for forward runs only

at 800–1000 °C it is due to prevalent Mg–Al exchange, suggesting a two-stage kinetic process. Therefore Cr and Fe³⁺, when substituting for Al in equal amounts, affect high-temperature behaviour in the spinel in different ways. Whereas Cr contents limit Mg–Al exchange, not starting before 700 °C, Fe³⁺ triggers trivalent–divalent cation exchange starting from 550 °C.

The model adopted here to determine cation distributions from bond lengths determined in situ seems to be reliable also for complex compositions.

Table 5 A1 Ee^{3+} and Mg dis-	
tribution between T and M sites	Τ (
as a function of temperature in	
sample L-Fe ²⁺	

<i>T</i> (°C)	T site			M site		
	Al	Fe ³⁺	Mg	Al	Fe ³⁺	Mg
L-Fe ³⁺						
1st run						
25	0.126	0.015	0.847	1788	0.067	0.143
200	0.126	0.013	0.850	1.791	0.068	0.138
400	0.120	0.016	0.852	1.795	0.066	0.137
450	0.131	0.011	0.849	1.786	0.069	0.143
400	0.120	0.012	0.860	1.799	0.067	0.133
25	0.131	0.007	0.852	1.786	0.073	0.139
2nd run						
500	0.123	0,015	0,853	1,795	0,065	0,138
450	0.118	0,014	0,857	1,798	0,067	0,133
3rd run						
550	0.123	0.023	0.844	1.794	0.057	0.147
600	0.148	0.024	0.819	1.769	0.056	0.174
650	0.180	0.030	0.781	1.735	0.050	0.214
700	0.221	0.037	0.733	1.693	0.043	0.263
800	0.242	0.036	0.714	1.672	0.044	0.282
900	0.263	0.034	0.695	1.649	0.048	0.300
1000	0.277	0.032	0.683	1.634	0.050	0.314

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