Potential problems in the interpretation of powder X-ray diffraction patterns from fine-dispersed $2M_1$ and 3T dioctahedral micas

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Abstract: XRD patterns were calculated for $2M_1$ and 3T dioctahedral mica structural models in which *trans*-vacant (tv) and *cis*-vacant (cv) layers were interstratified and arranged, independent of octahedral cation distribution, as in the periodic $2M_1$ and 3T mica polytypes. The XRD patterns of the interstratified $cv/tv2M_1$ and cv/tv3T models were shown to be similar to those of the $tv2M_1$ and tv3T mica structures, respectively. Because both the positions and shapes of *hkl* reflections in the XRD patterns of the $tv2M_1$ and cv/tv3T as well as of the $tv2M_1$ and $cv/tv2M_1$, respectively, are the same qualitative X-ray identification of the $tv2M_1$, $cv/tv2M_1$, tv3T and cv/tv3T mica varieties should be a complex problem, especially when the amount of cv layers is either rather low (10-15 %) or high (30-40 %) with significant tendency to segregation of the layer types.

The diffraction features of $cv/tv2M_1$ and cv/tv3T mica structures make it possible to use the Rietveld technique for the structural refinement. Correct application of this technique to the $cv/tv2M_1$ and cv/tv3T structures should reveal partial occupancy of *trans*-octahedra because of the ability of diffraction to average structural parameters. Pavese *et al.* (2001) found partial occupancies of *trans*-octahedra in 3T and $2M_1$ powder phengite samples. To explain these results the authors postulated the $\pm b/3$ stacking disorder. However, the calculations have shown that such disorder strongly modifies the XRD patterns. Therefore, in the paper an alternative, crystal-chemically more justified hypothesis is suggested according to which the occupancy of *trans*-sites in 3T and $2M_1$ phengites results from coexistence of tv and cv layers in their structures. Diffraction features a dioctahedral mica sample given as a tv3T standard in textbooks (Brindley, 1980; Bailey, 1984) are quite similar to those characteristic for a 3T structure in which tv and cv 2:1 layers are interstratified.

Key-words: *cis*-vacant layers, *trans*-vacant layers, $\pm b/3$ disorder, interstratification, $2M_1$, 3T mica polytypes, simulation of XRD patterns.

1. Introduction

Micas are rock-forming minerals which occur in quite different geological environments and are often used as markers of metamorphic grade. The mica structure consists of 2:1 layers separated by large interlayer cations, such as K or Na. In the general case, the octahedral sheet of a 2:1 layer contains three symmetrically independent sites differing in mutual dispositions of OH groups and oxygen atoms coordinating octahedral cations. They are termed as M1 or *trans*-sites, and M2 and M2', or *cis*-sites.

For a long time it was commonly accepted that in dioctahedral phyllosilicates and in micas, in particular, the *trans*-sites are vacant. This supposition was based on single crystal structural refinements where octahedral cations in 2:1 layers of various polytypes of muscovite, phengite, paragonite and margarite were shown to occupy only *cis*sites (Bailey, 1984). Therefore, identification of fine dispersed mica polytypes and illites, in particular, was usually based on the comparison between experimental positions and intensities of *hkl* reflections and those calculated for *trans*-vacant (*tv*) polytype structural models given in reference books (Brindley, 1980; Bailey, 1984).

The existence of dioctahedral phyllosilicates with different distributions of octahedral cations was first reported by Méring & Oberlin (1967) for a sample of Wyoming montmorillonite, for which one of the two symmetrically independent cis-octahedra of the 2:1 layer was shown to be vacant. Drits et al. (1984) deduced unit cell parameters and atomic coordinates for a one-layer monoclinic *cis*-vacant (cv) illite model (cv1M). These authors calculated powder XRD patterns for periodic cv1M and tv1M models, as well as for models in which cv and tv mica-like layers are interstratified. Based on these data they formulated diffraction criteria for identification of the periodic and interstratified mica-like structures as well as for those containing rotational stacking faults. Zvyagin et al. (1985) were the first to describe a monomineral Al-rich cv1M mica sample. Reynolds & Thompson (1993) and Drits et al. (1993) described other occurrences of cv1M illites. The state of the

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art in structural and crystallochemical study of micas was recently presented by Brigatti & Guggenheim (2000).

Tsipursky & Drits (1984), using oblique-texture electron diffraction (OTED), showed that dioctahedral smectites have a wide range of proportions of cv and tv layers. Reynolds (1993), McCarty & Reynolds (1995) and Drits et al. (1996) demonstrated that the cv structure exists in illite fundamental particles of the mixed-layer illite-smectites (I-S) exhibiting the 1M and 1Md stacking sequences, and that cv 2:1 layers are commonly interstratified with tv layers in these illite stackings. Coexistence of cv and tv layers in I-S formed in volcanic and sedimentary rocks was recently described in many papers (McCarty & Reynolds, 2001; Drits et al., 1998a, 2002; Ylagan et al., 2000; Lindgreen et al., 2000; Cuadros & Altaner, 1998). Drits & McCarty (1996) proposed a method for semiquantitative determination of coexisting tv and cv layers in 1M illite and I-S using experimental d(11l) values.

Drits et al. (1993) noted that tv3T and cv1M mica polymorphs produce similar diffraction effects, that is, in their XRD patterns hkl reflections have similar intensities and close positions. These features explain the reason for confusion in the identification of tv3T illite. Careful analysis of the experimental data published in the literature showed that illite varieties described as 3*T* (Warshaw, 1959; Ey, 1984; Halter, 1988) in fact, correspond to cv1M (Drits et al., 1993). Additional problems with reliable identification of the actual structure of illites and other fine dispersed dioctahedral micas may arise because of the similarity of XRD patterns corresponding to different mica varieties. For example, coexistence of tv and cv layers was found in a nonswelling illite sample from Amethyst Vein system, Creede Mining District, Colorado (Drits & McCarty, 1996) which was initially described as 3T illite (Horton, 1983). The reason is that powder XRD patterns corresponding to tv3T, cv1M and interstratification of tvand cv layers in 1Md illite in which cv layers prevail are very similar.

Recently Pavese et al. (2001) studied 3T and $2M_1$ powder phengite samples using low-temperature neutron diffraction and the Rietveld technique. Phengites have the idealized composition $K(Al_{1.5}Mg_{0.5})(Si_{3.5}Al_{0.5})O_{10}(OH)_2$ and attract special attention because crystallization of a specific polytype is supposed to depend on the thermobaric conditions, which, in their turn, control the cation ordering occurring in tetrahedral and octahedral sheets of the 2:1 layers (Sassi *et al.*, 1994). For the studied 3T and $2M_1$ samples Pavese et al. (2001) found that the occupancies of trans-octahedra of these micas are not zero but roughly equal to 0.18 and 0.12 atoms per octahedron, respectively. The title of their paper contains a question: are these occupancies real or an artefact? To explain the observed results, Pavese et al. (2001) postulated stacking disorder resulting from $\pm b/3$ slips in the octahedral sheet of the 2:1 layer that occur in $2M_1$ and 3T phengites. At the same time, the authors noted that the $\pm b/3$ disorder was observed experimentally only for trioctahedral micas (Noe & Veblen, 1999), as for dioctahedral micas, there are energy hindrances against this type of disorder. One has to note that a partial occupancy of trans-sites in Mg-Fe-bearing muscovites $2M_1$ was determined by Brigatti *et al.* (1998, 2000).

Taking into account that interstratification of tv and cv layers in illite stacking sequences exists, one may assume that the occupancy of *trans*-sites found in 3T and $2M_1$ phengite samples results from coexistence of both layer types in their structure. The main aim of this paper is to consider this alternative hypothesis using simulation of powder XRD patterns from different structural models. Special attention is paid to potential problems related with identification of dispersed dioctahedral mica varieties based on their powder XRD patterns.

2. Simulation of XRD patterns

The presence of stacking faults often prevents crystalchemical studies of dispersed layer minerals by conventional methods of structure analysis, including the Rietveld technique. One of the effective ways to obtain structural and crystal chemical information from such minerals is to simulate diffraction effects from realistic structural models and then to compare the calculated and experimental XRD patterns (Drits & Tchoubar, 1990). This approach was successfully applied to natural and synthetic phyllosilicates (Sakharov et al., 1990; Plançon et al., 1989; Reynolds, 1993; McCarty & Reynolds, 1995, 2001; Manceau et al., 2000; Drits et al., 1993) and phyllomanganates (Drits et al., 1998b; Lanson et al., 2000, 2002; Manceau et al., 1997). The mathematical formalism used to calculate powder diffraction patterns was described by Plançon & Tchoubar (1977), Plançon (1981), Sakharov et al. (1982) and, in greater detail, by Drits & Tchoubar (1990). One of the main advantages of this approach is that it can predict diffraction effects from diverse, theoretically possible and crystal chemically justified layer structural models. Its application makes it possible to analyze diffraction effects of various parameters describing specific features of periodic and defective structural models. For each particular model, the following structural and probability parameters should be defined: the structure and chemical composition of coexisting layers; the azimuthial orientations, translations, reflection operations and modes of alternation of the layers; probability parameters that describe quantitatively the model (e.g. proportion of each layer type, W_i, and conditional probability, P_{ij}, denoting the probability of a layer type j to follow layer type i, when the short-range factor is equal to 1). It is assumed that the interstratification of the layer types follows the Marcovian statistics (Drits & Tchoubar, 1990).

3. Structural models

Cartesian coordinates and parameters for the cv and tvunit cells used in this study are taken from Drits *et al.* (1984) and correspond to their designation of C and T layers. Let us term C^r and C^l the cv layers in which the fractional x, y coordinates of their vacant sites in the projection on the *ab* plane are equal to (0.346, -0.333) and (0.346, 0.333) respectively. In the notation of Zvyagin (2001) Cr and C^{l} layers are denoted as 51 and 15, respectively. These numbers correspond to intersheet displacements S_iS_i characterized by their normal projections on the *ab* plane in the fixed coordinate system. These displacement vectors are relating the upward succeeding sheets TO and OT in a 2:1 layer, respectively. The origins of the tetrahedral sheets are in the centers of the ditrigonal cavities and the origins of the octahedral sheets are in the centers of vacant octahedra. The unit cells of the C^r , C^l and T layers have the same dimensions a = 5.200 Å, b = 9.007 Å, $a\sqrt{3}$, $c^* =$ 9.98 Å and the same cation composition $K_{0.8}(Al_{1.70} Mg_{0.30})$ $(Si_{3,5} Al_{0,5}) O_{10} (OH)_2$. In projection on the *ab* plane, the position of K cation located in the ditrigonal hole of the upper tetrahedral sheet of the 2:1 layer is shifted with respect to the closest K site located in the lower tetrahedral sheet by -0.383a and -0.3085a along the *a* axis for the *tv* and cv unit cells, respectively.

3.1 The $\pm b/3$ disordered model

In accordance with Nespolo (2001) and Nespolo & Ferraris (2001), Pavese *et al.* (2001) assumed that stacking disorder occurring in phengite samples consists in $\pm b/3$ slips in the octahedral sheets of their 2:1 layers. These slips move the upper plane of the apical oxygen atoms with respect to the lower plane in the octahedral sheet. In order to preserve the octahedral coordination for interlayer K cations, the $\pm b/3$ shift in the nth layer requires the same shift for the layer located just above it. As a result, in dioctahedral micas the position of a filled *cis*-site is superposed on to that corresponding to a vacant *trans*-site in the defect-free structure.

If all layers have the same azimuthual orientation, then five types of 2:1 layers are required to describe the $\pm b/3$ disordered structural model. Along with a defect-free T layer, they include two "defective" layers, in which their upper parts including the upper plane of the apical oxygen atoms of the octahedral sheets are shifted with respect to the lower planes by + b/3 (T^u layer) or -b/3 (T^u layer) along the b axis, respectively. Atomic coordinates for T^{u}_{+} and T^{u} layers are obtained by adding and substracting 0.333b from the Y-coordinates of atoms located above the octahedral cation plane in T layer, respectively. For the next two defective layers in which their lower parts are shifted with respect to the upper one by + b/3 and - b/3, (T^l₊ and T^{l} layers), respectively, the 0.333b value is added and subtracted from the Y-coordinates of atoms located below the octahedral cation plane. To preserve the octahedral coordination of interlayer K cations some layer pairs must be excluded from the consideration and only the following pairs may exist:

TT; TT^u₊; TT^u₋; T^I₊T; T^I₊T; T^u₊, T^I₊, T^I

The layer pairs having the same upper indexes but different lower ones (e.g. TT^u_+ and TT^u_- ; $T^u_+T^l_+$ and $T^u_-T^l_-$) can be shown to be equivalent from the diffraction point of view. Therefore, to describe the $\pm b/3$ stacking disorder in tv1M

Table 1. Matrice of conditional probabilities for the $\pm b/3$ disordered $2M_1$ models.

	T ₆₀	T60	T^{u}_{60}	T ^u -60	T_{60}^{1}	T ¹ 60	
T ₆₀	0	Р	0	1-P	0	0	
T60	Р	0	1-P	0	0	0	
T^{u}_{60}	0	0	0	0	0	1	
T ^u -60	0	0	0	0	1	0	
T^1_{60}	0	Р	0	1 - P	0	0	
T ¹ -60	Р	0	1 - P	0	0	0	

structure, three layer types, T, T^u_+ and T^l_+ or T, T^u_- and T^l_- can be used. For simplicity, the lower subindexes in the layer symbols may be omitted.

For a defective $2M_1$ structure, adjacent layers should be rotated with respect to each other by $\pm 120^{\circ}$ or related by glide plane as in the periodic $tv2M_1$ mica. Therefore, six layer types, T_{60} , T_{-60} , T_{60}^{u} , T_{60}^{u} , T_{60}^{l} and T_{-60}^{l} should be used to describe the $\pm b/3$ stacking sequence in defective $2M_1$ micas. Lower indices 60 and -60 correspond to rotations of T, T^u and T^l layers by 60° counter-clockwise and clockwise, respectively, relative to the fixed a axis. One may assume that the occurrence probabilities $W(T_{60}) = W(T_{-60}) = W$ and W(T₆₀) = W(T₋₆₀) = W(T₆₀) = W(T₋₆₀) - W(T₋₆₀) - W and W(T₆₀) = W(T₋₆₀) = W(T₆₀) = W(T₋₆₀) = W_i. Table 1 shows that among the 36 values of P_{ij} only eight should be deter-mined, because the other P_{ij} values are equal to zero or to 1.0. In the first case, P_{ij} = 0 because the corresponding layer pairs are forbidden, whereas in the second case, $P_{ii} = 1$ because the T^{*l*} layers should always follow the T^{*u*} layers for definition. If there are no additional constraints for the layer stacking, then, as can be seen in Table 1, only one parameter, P, should be defined. Its value is equal to the total amount of T layers, that is, to 2W. In this case equations of the type W(T₆₀) = W(T₋₆₀) P (T₋₆₀ T₆₀) + W(T^{*l*}₋₆₀) P(T^{*l*}₋₆₀ T₆₀), P(T₋₆₀ T₆₀) + P(T₋₆₀ T^u₆₀) = 1, and 2(W + W_i) = 1, satisfying the Marcovian statistics are fulfilled.

For a defective 3*T* structural model, each layer should be rotated with respect to the preceding one by 120°, as in the periodic *tv*3*T* mica. In this case, nine layer types, T_0 , T_{120} , T_{240} , T_0^u , T_{120}^u , T_{240}^u , T_0^l , T_{120}^l and T_{240}^l should be used. As in the previous case, one can assume that W(T_0) = W(T_{120}) = W(T_{240}) = W and portions of the other 6 layer types, W_i , are also equal to each other, so that 3W + 6 W_i = 1. Within the constraints described above, only W and W_i values are needed to describe the $\pm b/3$ defective 3*T* model. Indeed, in this case the conditional probability of the T_0 (or T_{120} , or T_{240}) layer to follow the T_{240}^l (or T_0^l , T_{120}^l) is equal to 3W, whereas the conditional probability of the T^u_0 (or T^u_{120} , T^u_{240}) layer to follow the T_{240} (or T_0 , T_{120}) is equal to 6 W_i . The rest of P_{ij} values differing from 0 and 1 can be found using the equations: $P(T_{240}^l, T_0) + P(T_{240}^l, T_0^u) = 1$ or $P(T_0 T_{120}) + P(T_0 T_{120}^u) = 1$, *etc*.

Thus, in order to describe the $\pm b/3$ defective $2M_1$ and 3T models, the occurrence probability for each layer type should be defined.

3.2 Interstratification of cv and tv layers (tv/cv models)

In a defective $tv/cv 2M_1$ mica model, four layer types, T₆₀, T₋₆₀, C^r₆₀, and C^r₋₆₀, or T₆₀, T₋₆₀, C^l₆₀ and C^l₋₆₀ are interstratified in such a way that the layers in any layer pair, independent of octahedral cation distribution, are rotated with respect to each other by $\pm 120^{\circ}$ as in a periodic $tv2M_1$ mica. One may assume, that the occurrence probabilities $W(T_{60}) = W(T_{-60}) = W$ and $W(C^r_{60}) = W(C^r_{-60}) = W_i$. In the case of random interstratification of the layer types having identical azimuthial orientations, the conditional probabilities, P, for the layer T₋₆₀ to follow the layer T₋₆₀ and C^r₋₆₀ as well as for the layer C^r₋₆₀ to follow the layers T₋₆₀ and C^r₋₆₀ as well as for the layer C^r₋₆₀ to follow the layers T₋₆₀ and C^r₋₆₀ as well as for the layer C^r₋₆₀ to follow the layers T₋₆₀ and C^r₋₆₀ as well as for the layer C^r₋₆₀ to follow the layers T₋₆₀ and C^r₋₆₀ are both equal to 2W_i and P + P' = 2(W + W') = 1.

In defective cv/tv 3T mica models, six layer types, T_0 , T_{120} , T_{240} , C_0^r , C_{120}^r and C_{240}^r or T_0 , T_{120} , T_{240} , C_0^l , C_{120}^l and C_{240}^l are interstratified in such a way that each layer is rotated, with respect to the preceding one, by 120°, counter-clockwise for the first six layer types and clockwise for the second layer sextet. For periodic cv3T structures, in terms of Zvyagin's notations, the counter-clockwise and clockwise periodic rotations of C^r layers by 120° are represented as 51 13 35 and 51 35 13, respectively, whereas the corresponding rotations of C^l layers, as 15 31 53 and 15 53 31. Let us assume that $W(T_0) = W(T_{120}) = W(T_{240}) = W$ and $W(C_0^r) = W(C_{120}^r) = W(C_{240}^r) = W_i$. As in the previous case, if the layer types having identical azimuthal orientation are interstratified at random then the conditional probabilities, P, for the layer T_{120} to follow the layers T_0 and C_0^r , and for the layer T_{240} to follow the layers T_{120} and C_{120}^r are equal to 3W. The other nonzero conditional probabilities, P', are equal to $3W_i$, so that $3(W + W_i) = 1$.

In the general case, additional constraints on the layer sequences may be imposed. For example, the layer types may interstratify with some tendency to segregation. In this case, the P and P' values equal to 3W and 3W_i for the random interstratified model may vary within $3W < P \le 1$ and $3W_i > P' \ge 0$, respectively. The case P = 1 and P' = 0 corresponds to a physical mixture of periodic tv3T and cv3T micas. In the case when P > 3W and $P' < 3W_i$ the occurrence probabilities for the layer subsequences $T_0T_{120}T_{240}T_0...$ and $C_0^rC_{120}^rC_{240}C_0^r$ increase in comparison with those for the layer subsequences in the random model. Simultaneously, the occurrence probabilities of layer subsequences of mixed layer type compositions decrease.

More complex cv/tv3T structural models may consist of nine interstratified layer types, T_0 , T_{120} , T_{240} , C_0^r , C_{120}^r , C_{240}^r , C_0^r , C_{120}^r , C_{240}^r , C_0^r , C_{120}^r and C_{240}^l . These layers may be distributed at random or with some tendency to segregation.

3.3 Models with partial occupancy of trans-sites

In these models *trans*-sites in the octahedral sheets of $2M_1$ and 3T mica models are occupied by Al or Mg cations. These models will be denoted as $m2M_1$ and m3T where m is the occupancy factor equal to 0.10; 0.20, *etc.* atoms per *trans*-site whereas the occupancy of *cis*-sites is equal to 0.95; 0.90, *etc.* atoms per *cis*-site. The atomic coordinates remain the same independent of the occupancy of *trans*-and *cis*-sites.

4. Results

Powder XRD patterns have been calculated for the models described above in the interval $19.0^{\circ} - 65.0^{\circ} 2\theta$. Intensities were calculated with a step equal to $0.02^{\circ} 2\theta$. Thicknesses of coherent scattering domains (CSDs) were distributed log-normally and the parameters of this distribution were determined using the mean number of mica layers in CSDs and the regression given by Drits et al. (1997) with mean (N) and maximum (N_{max}) numbers of mica layers in CSDs as variable parameters. For all models, $N_{max} = 50$ and N varies from 15 to 12. CSDs were assumed to have disc-like shape, with the diameter of 500 Å. In order to compare XRD patterns calculated for different models, reflection intensities were normalized to those corresponding to reflections whose intensities do not depend on the distribution of octahedral cations over cisand trans-sites. These intensities correspond to 20l, 13l and 11*l* reflections of the $2M_1$ and 3T structural models, respectively. In order to estimate the relative difference between the intensities of XRD patterns calculated for two compared models, the standard factor R_{wp} was used.

4.1 Diffraction effects from m3T and $m2M_1$ models

Figure 1 compares XRD patterns calculated for periodic tv3T and $tv2M_1$ micas with those calculated for m3T and $m2M_1$ structural models with m = 0.20 atoms per M1-site. Figure 1 shows that this partial occupancy is accompanied by three principal diffraction features. First, the intensities of 110 and 100 reflections corresponding to the $m2M_1$ and m3T models, respectively, decrease. Second, 11*l*, 02*l* and 10*l* reflections of the $m2M_1$ and m3T models have different sensitivity to redistribution of octahedral cations over trans- and cis-sites. For example, the intensities of some 02l, 11l reflections decrease, whereas the intensities of the other reflections are similar for the $m2M_1$ and periodic $tv2M_1$ micas (Fig.1a). In the case of the m3Tmodel, the cation redistribution significantly decreases intensities the first four 10*l* reflections with l < 5 (Fig. 1b). Third, the intensities of *hkl* reflections located at $2\theta > 34^{\circ}$ are almost insensitive to the cation distribution. It is obvious that the difference between the compared XRD patterns depends on the occupancy of the trans-sites. The calculations have shown that for m = 0.20 atoms, $R_{wp} =$ 7.4 % and 7.5 % for the m3T and $m2M_1$, respectively; but $R_{\rm wp}$ values decrease down to 3.7% for both models for m = 0.10 and increase up to 11.2 % and 11.1 % when m = 0.30.

4.2 Diffraction effects from tv/cv3T models

XRD patterns calculated for tv/cv3T models consisting of different amounts of interstratified cv and tv layers and differing from each other in the modes in the layer distri-



Fig. 1. Comparison of XRD patterns calculated for the $tv2M_1$ (solid line) and $m2M_1$ models (a) as well as for the tv3T (solid line) and m3T models (b). m = 20 %. (N = 15).

bution were compared with that calculated for a periodic tv3T mica. It was found that the XRD patterns of the tv3Tand cv/tv3T containing 10 % of cv layers are almost undistinguishable ($R_{wp} = 6.3$ %). Further increase of cv layers is accompanied more significant modification (Fig. 2a). In particular, intensity of 102 and 103 reflections dramatically decreases whereas 106 reflection becomes more strong in comparison with corresponding reflections of the tv3T. Accordingly, the $R_{\rm wp}$ values increase and equal to 9.0 %, 12.5 % and 15.3 % when amount of cv layers are equal to 20 %, 30 % and 40 %. It is remarkable, however, that segregation of the interstratified layer types in tv/cv3Tmodels significantly decreases the difference between XRD patterns calculated for the models containing different amount of cv layers. For example the XRD patterns corresponding to cv/tv3T models consisting of 20 % and 10 % cv layers almost coincide (R = 2.1%) when layers in both models are distributed with P = 0.90. Similarly, XRD patterns calculated for tv/cv3T models containing 20, 30 and 40 % of cv layers in which the layer types are distributed with P = 0.80 are also very similar. Comparison of these patterns shows that for any two patterns the R_{wp} value does not exceed 4.3 %. Therefore, it is not surprising that XRD patterns calculated for a segregated tv/cv3T model containing 40 % of cv layers remain quite similar to that of tv3T mica except for 10*l* reflections with $l \le 4$ (Fig. 2b). For the compared XRD patterns, R_{wp} values vary from 7.5 to 10.8 % depending on the cv layer amount.

Redistribution of 10*l* reflections observed for cv/tv3T models consisting in nine interstratified layer types, T₀, T₁₂₀, T₂₄₀, C'₀, C'₁₂₀, C'₂₄₀, C'₀, C'₁₂₀, and C^r₂₄₀ is quite similar to that observed for the cv/tv3T models containing six layer types. However, increase of cv layers in these nine-component systems is accompanied more significant decreasing



Fig. 4. Comparison of XRD patterns corresponding to the tv3T (a) and $tv2M_1$ (b) (solid lines)and to the $\pm b/3$ disordered 3T and $2M_1$ models containing 20 % of defective T^u and T^l layers (N = 15) (see text).

of 10*l* reflections with l < 5. Therefore, R_{wp} values increase up to 10.0, 14.2 and 17.8 % when total amount of *cv* layers are equal to 20, 30 and 40 %.

4.3 Diffraction effects from $tv/cv2M_1$ model

XRD patterns calculated for $tv/cv2M_1$ models containing different total amounts of cv layers, W(C), were compared with those calculated for periodic $tv2M_1$ mica structures. An increase in cv layers is accompanied mainly by a redistribution of 02*l*, 11*l* reflections intensities. In particular, the intensities of 110, 111, 022, 113, 024 and 115 reflections decrease, whereas the other reflections are not so sensitive to variations in cv layers content. However, as in the case of cv/tv3T models when the amounts of cv layers are relatively small (*e.g.* W(C) = 10 %), the XRD patterns of the $tv/cv2M_1$ and pure $tv2M_1$ are practically undistinguishable ($R_{wp} = 7.3$ %). Figure 3 shows that even when W(C) = 20 % the difference between the compared

XRD patterns remains relatively small ($R_{wp} = 9.7$ %). For $tv/cv2M_1$ models with W(C) > 25 %, the intensity redistribution becomes stronger, and the models may be identified taking account of very low intensities of 111, 022, 112 and 113 reflections.

As mentioned above, the interstratification of cv and tv layers is accompanied by the interstratification of different interlayer translations equal to -0.383a and -0.3085a which are associated with tv and cv layers, respectively. It is well known that the remarkable feature of diffraction effects is their ability to average interstratified structural parameters (Méring, 1949; Drits & McCarty, 1996). Similar effects are observed in the case of $tv/cv2M_1$ mica models. Both layer types have two azimuthal orientations rotated with respect to each other by 120°. Along each of these directions, as well as along the resulting a axis, the two interlayer translations are interstratified in proportion determined by the amount of tv and cv layers. Therefore diffraction averages interstratified -0.383a and -0.3085a

translations. It means that each given $tv/cv2M_1$ structure is characterized by a statistically weighted interlayer displacement equal to $T_{ef} = -0.383 W(T) - 0.3085 W(C)$. Therefore, the positions of hkl reflections in XRD patterns of the $tv/cv2M_1$ models depend on the contents of the interstratified layers. Moreover, for each given $tv/cv2M_1$ structure, the positions of hkl reflections should coincide with those of $tv2M_1$ structure, for which the projection of the c axis on the *ab* plane ($c \cos \beta/a$) is equal to T_{ef} calculated for a given tv/cv ratio. The validity of these considerations is confirmed by full coincidence of hkl reflections' positions in the compared XRD patterns shown in Fig. 3. In particular, the XRD patterns of $tv2M_1$ micas were calculated with c cos β/a equal to -0.3756a, -0.3681a and -0.3607ain order to be compared with the XRD patterns of the $tv/cv2M_1$ models containing 10 %, 20 % and 30 % of cvlayers.

4.4 Diffraction effects from $\pm b/3$ disordered models

XRD patterns calculated for the $\pm b/3$ disordered 3T and $2M_1$ models having 10 % and 20 % of defective T^u and T¹ layers were compared with XRD patterns of defect-free tv3T and $tv2M_1$ micas. An increase in the amount of defective layers is accompanied by increasing background within of 2θ from 19° to 34° and decreasing intensities of *hkl* reflections located in this diagnostic interval of 2θ . Note that *hkl* reflections located at $2\theta > 34^\circ$ have low sensitivity to variation in the "defective" layer content. At low amount of T^u and T¹ layers (e.g. 10 %), modifications of the XRD patterns are relatively small and R_{wp} factor is equal to 12.8 % and 9.0 % when the XRD patterns of the defective 3T and $2M_1$ structural models are compared with those of the defect-free tv3T and $tv2M_1$. However, Fig. 4a and 4b shows that the difference between the compared XRD patterns increases dramatically when the amount of the "defective" layers is equal to 20 % and R_{wp} values are equal to 22.2 % and 16.0 for the 3T and $2M_1$ models.

5. Discussion

The results of the calculations allow us to answer the question of Pavese et al. (2001) concerning the origin of partial occupancy of *trans*-sites in the phengite samples, as well as to reveal certain problems which should be taken into account to determine the actual structure of finedispersed dioctahedral 3T and $2M_1$ mica polytypes using their powder XRD patterns. The presence of the T^u and T^l layers in the $2M_1$ and 3T structures is accompanied by displacements of some octahedral and tetrahedral cations as well as oxygen atoms from their standard positions to those which are unoccupied in the defect-free polytypes. The contribution of these "shifted" atoms disturbs the reflections' intensities. Due to the ability of diffraction to average structural parameters, a XRD pattern from the \pm b/3 disordered mica structure should be treated in terms of the unit cell in which occupancies can appear in positions generated by the structural disorder. Pavese et al. (2001) assumed that the scaling of the structural factors should greatly smooth the effect of the contribution of the shifted atoms to the reflections' intensities, but they did not estimate this effect quantitatively.

Analysis of the XRD patterns calculated from the $\pm b/3$ disordered $2M_1$ and 3T models shows that the main contribution to the modification of the patterns results not from the partial occupancy of trans-sites in the defective T^u and T¹ layers but from the disordered displacements of the tetrahedral sheets and oxygen atoms of the octahedral sheets in these layers by $\pm b/3$. Indeed, as shown in Fig. 1, partial occupancy of trans-sites decreases the intensities of only some of the 11*l*, 02*l* and 10*l* reflections of the $m2M_1$ and m3T models and does not change the background. In contrast, the $\pm b/3$ layer displacements decrease the intensities of all 111, 021 and 101 reflections and dramatically increase the background (Fig. 4). These features are typical for the XRD patterns of the $\pm b/3$ disordered phyllosilicates (Drits & Tchoubar, 1990; Drits et al., 1984). In particular, to get a fractional occupancy of trans-sites of about 20 %, the $\pm b/3$ defective 3T structure should contain similar amounts of T^u and T^l layers. As a consequence, the background increases and 10l reflections' intensities decrease so dramatically that the scaling of the structural factors cannot smooth this effect (Fig. 4).

Even when the amounts of T^l and T^u layers in the $2M_1$ and 3T micas are rather small, and corresponding XRD patterns can be used for the Rietveld refinement the presence of the $\pm b/3$ disorder can be detected using the fact that background and intensities of 11l, 02l and 10l reflections located in the interval $19^\circ\text{-}34^\circ$ 2θ are most sensitive to this disorder. Therefore, R_{wp} factors determined for two different intervals of the XRD pattern (19°- 34°20 and $2\theta \ge 34^{\circ}2\theta$) should be significantly different if the intensity in each interval is scaled using reflections which are not sensitive to the $\pm b/3$ disorder. In this case, $R_{\rm wp}$ factor for the first interval should be much higher than that for the second one. For example, comparison of XRD patterns calculated for 3T model containing 10% of the defective layers and for tv3T model shows $R_{wp} = 14.7$ % for the first and $R_{wp} = 8.55$ % for the second interval.

Diffraction effects from $tv/cv2M_1$ and tv/cv3T models are, in general, quite similar to those observed for $m2M_1$ and m3T models especially when W(C) = m. Indeed, an increase in the *cv*-layer amount, as well as in the occupancy of *trans*-sites in the models is accompanied by decreasing intensity of reflections having the same *hkl* indices. This similarity does not mean that the XRD patterns of the tv/cv3T and m3T completely coincide because average overlapping of atomic positions of the interstratified cv and tv layers is not equivalent to the atomic sites of the layers with a partial occupancy of trans-sites. It means, however, that the main modification of the XRD patterns from the tv/cv3T structures appears as a result of the partial occupancy of *trans*-sites mimicking the presence of cv layers in the actual structure. For example, 20 % of cv layers in the tv/cv3T (Fig. 2b) and 20 % atoms in trans-sites in the m3T (Fig. 1a) result in a similar decrease in the intensities of 10lreflections having the same l values. Similar diffraction features are observed for $tv/cv2M_1$ models. For example, comparison of Fig. 1b and 3 shows that an increase in the

Fig. 5. XRD patterns calculated for the tv3T(a), cv3T(b) and as well as for the tv3T (solid line) and cv/tv3T (c) structural models. In the last one 80 % of T and 20 % of C^r layers are clockwise rotated with respect to preceding one by 120°.

amount of cv layers in the $tv/cv2M_1$ structures as well as partial occupancy of *trans*-sites in the $m2M_1$ ones are accompanied by similar decrease in the intensities of reflections having the same hkl indices. A remarkable diffraction feature of the $tv/cv2M_1$ is that the positions of hkl reflections observed in each XRD pattern can be described in terms of the unique unit cell despite the interstratification of two different interlayer translations. The fact that in this cell, the effective interlayer displacement along the *a* axis, $|c \cos \beta/a|$, decreases with the amount of cv layers, can be used as an additional criterion for the identification of the $tv/cv2M_1$ micas, taking into account that $|c \cos \beta/a| \ge 0.383$ for micas of muscovite-phengite composition (Bailey, 1984).

Qualitative identification of the tv/cv3T and $tv/cv2M_1$ micas should be especially problematic in the following cases: low content of cv layers (< 10 %); a strong tendency of cv and tv layers to be segregated in tv/cv3T; very small sizes of CSDs; presence of rotational stacking faults. In the first case, the XRD patterns from the interstratified and defect-free 3T and $2M_1$ models are quite similar or even undistinguishable. As was mentioned, a strong tendency to segregation of the layer types in the tv/cv3T significantly decreases the difference between the XRD patterns corresponding to the models having quite different amounts of cv layers. The main reason is that XRD patterns corresponding to the periodic $T_0T_{120}T_{240}T_0$... and $C_0^rC_{120}^rC_{240}^rC_0^r$... layer sequences (space group $P3_112$) are rather similar. As can be seen in Fig. 5a and 5b, they have the same positions and shapes of hkl reflections and differ from each other only by lower intensities of 105 and 108 reflections corresponding to the cv3T structure. The significant difference between the intensities of the 106 reflections in the compared XRD patterns has no diagnostic meaning because of their overlapping with 006 reflection. Finally, small sizes of CSDs and stacking faults are accompanied by broadening of *hkl* reflections and decreasing difference between the compared XRD patterns.

However, when the amount of cv layers in $tv/cv2M_1$ and tv/cv3T is rather high (> 20 %) and they are distributed "at random" among tv layers, then the following criteria can be used for their identification: absence or extremely low intensities of 102 and 103 reflections for tv/cv3T (Fig. 2) and of 111 and 022 ones for $tv/cv2M_1$ (Fig. 3). Because of the absence of 102 and 103 reflections one may suspect that a dioctahedral mica sample given as a tv3T standard in the textbooks (Brindley, 1980; Bailey, 1980, 1984) consists of interstratified cv and tv layers.

Because both positions and profiles of hkl reflections of the $tv/cv2M_1$ and tv/cv3T models are the same as those of periodic and defect-free structures, the Rietveld technique may be used to refine mica samples consisting of tv and cvlayers using both X-ray and neutron diffraction. However, in the general case, along with conventional tv mica polytypes, initial models for the Rietveld refinement should include a partial occupancy of *trans*-sites. This requirement may have a practical meaning. It is well known that cation ordering over the tetrahedral and octahedral sites in dioctahedral micas is the subject of many publications (see review in Pavese, 2002). Special interest to this problem for phengites is derived from the assumption that the stability of $2M_1$ and 3T phengites is related to the cation ordering occurring in T and M sites. However, it would be difficult to achieve reliable results in terms of cation partitioning if the actual structure of phengite samples consists of tv and cv layers.

Let us consider some crystal-chemical features of the tv3T, cv3T and tv/cv3T structures which, most probably, make their XRD patterns similar. As was mentioned, in the tv3T, cv3T and cv/tv3T structures, T and/or Cr layers are counter-clockwise rotated with respect to the preceding one by 120° . Therefore, in the normal projection on the *ab* plane one of the two sets of symmetrically independent octahedral cations of the subsequent T and/or Cr layer almost superimpose on the OH groups located in the lower oxygen plane of the octahedral sheet of the preceding layer. The other half of the octahedral cations, as well as vacant octahedral sites of the subsequent layer, almost coincide with the tetrahedral cations of the lower sheet of the preceding layer. Such homogeneous disposition of the octahedral and tetrahedral cations, as well as OH groups in the adjacent layers is more favourable in comparison with a cation distribution when in the projection on the *ab* plane all octahedral cations of the subsequent layer superimpose on the tetrahedral cations of the following layer. Such an unfavourable mutual cation disposition takes place in cv/tv3T mica models in which the subsequent T or Cr layers are clockwise rotated with respect to the preceding one by 120°. In such structures, mutual dispositions of octahedral and tetrahedral cations in the adjacent layer pairs are different. In particular, in the projection on the *ab* plane, all octahedral cations of the following layer in the $C^r - C^r$ and $C^r - T$ layer pairs almost coincide with tetrahedral cations of the lower sheet of the preceding layer. In contrast, for layer pairs T - T and $T - C^r$, only half of the octahedral cations overlap with tetrahedral ones. That is why the XRD patterns corresponding to such structural models (Fig. 5c) are quite similar to those characteristic for disordered micas containing rotational stacking faults.

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References

- Bailey, S.W. (1980): Structures of layer silicates. *in* "Crystal structures of clay minerals and their X-ray identification". Eds: G.W. Brindley & G. Brown, Mineralogical Society, London SW7 5HR, 495p.
- (1984): Crystal chemistry of the true mica: *in* "Micas, Reviews in Mineralogy". S.W. Bailey, ed., Mineralogical Society of America, 13-66.
- Brigatti, M.F. & Guggenheim, S. (2000): Mica crystal chemistry and the influence of intensive variables on atomistic models. *Advanced on Micas (Problems, Methods, Applications in Geodynamics)*. Accademia Nazionale dei Lincei, Roma, Italy, 29-56.

- Brigatti, M.F., Frigieri, P., Poppi, L. (1998): Crystal-chemistry of Mg-, Fe-bearing muscovites 2M₁. Am. Mineral., 83, 775-785.
- Brigatti, M.F., Frigieri, P., Ghezzo, C., Poppi, L. (2000): Crystal chemistry of Al-rich biotites coexisting with muscovites in peraluminous granites. *Am. Mineral.*, 85, 436-448.
- Brindley, G.W. (1980): Order-disorder in clay mineral structures. *In*: Brindley, G.W. and Brown, G. eds., "Crystal structures of clay minerals and their X-ray identification". *Mineral. Society, London*, 125-196.
- Cuadros, J. & Altaner, S.P. (1998): Characterization of mixed-layer illite-smectite from bentonites using microscopic, chemical and X-ray methods: constrains on the smectite-to-illite transformation mechanism. *Am. Mineral.*, 83, 762-774.
- Drits, V.A. & McCarty, D. (1996): A simple technique for a semiquantitative determination of the *trans*-vacant and *cis*-vacant 2:1 layer contents in illites and illite-smectites. *Am. Mineral*, **81**, 852-863.
- Drits, V.A. & Tchoubar, C. (1990): "X-ray Diffraction of Disordered Lamellar Structures. Theory and Application to Microdivided Silicates and Carbons". Springer Verlag, 242 pp.
- Drits, V.A., Plançon, A., Sakharov, B.A., Besson, G., Tsipursky, S.I., Tchoubar, C. (1984): Diffraction effects calculated for structural models of K-saturated montmorillonite containing different types of defects. *Clay Minerals*, 19, 541-562.
- Drits, V.A., Weber, F., Salyn, A., Tsipursky, S. (1993): X-ray identification of 1*M* illite varieties. *Clays and Clay Minerals*, 28, 185-207.
- Drits, V.A., Salyn, A.L., Sucha, V. (1996): Dynamic and mechanism in the structural transformation of illite-smectites from Dolna Ves hydrothermal deposits. *Clays and Clay Minerals*, 44, 181-190.
- Drits, V.A., Środoń, J., Eberl, D.D. (1997): XRD measurement of mean illite crystallite thickness: Reappraisal of the Kübler index and the Scherrer equation. *Clays and Clay Minerals*, 45, 461-475.
- Drits, V.A., Lindgreen, H., Salyn, A.L., Ylagan, R., McCarty, D.K. (1998a): Semiquantitative determination of *trans*-vacant and *cis*-vacant 2:1 layers in illites and illite-smectites by thermal analysis and X-ray diffraction. *Am. Mineral.*, 83, 31-73.
- Drits, V.A., Lanson, B., Gorshkov, A.I., Manceau, A. (1998b): Suband super-structure of four-layer Ca-exchanged birnessite. *Am. Mineral.*, 83, 97-118.
- Drits, V.A., Sakharov, B.A., Dainyak, L.G., Salyn, A.L., Lindgreen, H. (2002): Structural and chemical heterogeneity of illite-smectites from Upper Jurrasic mudstones of East Greenland related to volcanic and weathered parent rocks. *Am. Mineral.*, 87, 1590-1607.
- Ey, F. (1984): Un exemple de gisement d'uranium sous discordance: les minéralisations protérozoiques de Cluff Lake, Saskatchewan, Canada: Thèse doctorale, Université Louis Pasteur, Strasbourg 1.
- Halter, G. (1988): Zonalité des altérations dans l'environement des gisements d'uranium associée à la discordance du Protérozoique moyen (Saskatchewan, Canada): Thèse doctorale, Université Louis Pasteur, Strasbourg 1.
- Horton, D. (1983): Argillitic alteration associated with the amethyst vein system, Creede Mining District, Colorado. Ph.D. dissertation, University of Illinois, Urbana-Champaign.
- Lanson, B., Drits, V.A., Silvester, E., Manceau A. (2000): Structure of H-exchanged hexagonal birnessite and its mechanism of formation from Na-rich monoclinic buserite at low pH. Am. Mineral., 85, 826-838.
- Lanson, B., Drits, V.A., Gaillot, A-C., Silvester, E., Plancon, A., Manceau, A. (2002): Structure of heavy metal sorbed birnes-

site: Part I. Results from X-ray diffraction. Am. Mineral., 87, 1631-1645.

- Lindgreen, H., Drits, V.A., Sakharov, B.A., Salyn, A.L., P.Wrang, Dainyak, L.G. (2000): Illite-smectite structural changes during metamorphism in black Cambrian Alum shales from the Baltic area. *Am. Mineral.*, 85, 1223-1238.
- Manceau, A., Drits, V.A., Silvester, E.J., Bartoli, C., Lanson, B. (1997): Structural mechanism of Co³⁺ oxidation by the phyllomanganate buserite. *Am. Mineral.*, **82**, 1150-1175.
- Manceau, A., Lanson, B., Drits, V.A., Chateigner, D., Gates, W.P., Wu, J., Huo, D., Stucki, J.W. (2000): Oxidation-reduction mechanism of iron in dioctahedral smectites. I. Crystal chemistry of oxidized reference nontronites. *Am. Mineral.*, 85, 133-152.
- McCarty, D.K. & Reynolds, R.C. (1995): Rotationally disordered illite-smectite in Paleozoic K-bentonites. *Clays and Clay Minerals*, 43, 271-284.
- —, (2001): Three-dimensional crystal structures of illite-smectite minerals in Paleozoic K-bentonites from the Appalachian basin. *Clays and Clay Minerals*, **49**, 24-35.
- Mering, J. (1949): L'interférence des rayons X dans les systèmes à stratification désordonnée. *Acta Cryst.*, **2**, 371-377.
- Mering, J. & Oberlin, A. (1967): Electron-optical study of smectites. Clays and Clay Minerals, 15, 3-25.
- Nespolo, M. (2001): Perturbative theory of mica polytypism. Role of the M2 layer in the formation of inhomogeneous polytypes. *Clays and Clay Minerals*, **49**, 1-23.
- Nespolo, M. & Ferraris, G. (2001): Effects of the stacking faults on the calculated electron density of mica polytypes – the Ďurovič effect. *Eur. J. Mineral.*, **13**, 1035-1045.
- Noe, D.C. & Veblen, D.R. (1999): HRTEM analysis of dislocation cores and stacking faults in naturally deformed biotite crystals. *Am. Mineral.*, 84, 1925-1931.
- Pavese, A. (2002): Neutron powder diffraction and Rietveld analysis; application to crystal-chemical studies of minerals at non-ambient conditions. *Eur. J. Mineral.*, 14, 241-249.
- Pavese, A., Ferraris, G., Pishedda, V., Fauth, F. (2001): M1-site occupancy in 3T and $2M_1$ phengites by low temperature neutron powder diffraction: reality or artifact? *Eur. J. Mineral.*, **13**, 1071-1078.
- Plançon, A. (1981): Diffraction by layer structures containing different kinds of layers and stacking faults. J. Appl. Cryst., 14, 300-304.
- Plançon, A. & Tchoubar, C. (1977): Determination of structural defects in phyllosilicates by X-ray powder diffraction. I. Principle of calculation of the diffraction phenomenon. *Clays* and Clay Minerals, 25, 430-435.
- Plançon, A., Giese, R.F., Snyder, R., Drits, V.A., Bookin, A.S. (1989): Stacking faults in the kaolin-group minerals: Defect structures of kaolinite. *Clays and Clay Minerals*, **37**, 203-210.
- Reynolds, R.C., Jr. (1993): Three-dimensional X-ray diffraction from disordered illite: simulation and interpretation of the diffraction patterns. *In*: "Computer Applications to X-ray Diffraction Methods" (R.C. Reynolds and J. Walker, eds.). Clay Minerals Society, Workshop Lectures 5, 44-78.
- Reynolds, R.C. & Thompson, C.H. (1993): Illites from the Postam sandstone of New York, a probable noncentrosymmetric mica structure. *Clays and Clay Minerals*, **41**, 66-72.
- Sakharov, B.A., Naumov, A.S., Drits, V.A. (1982): X-ray diffraction by mixed-layer structures with random distribution of stacking faults. *Doklady Akademii Nauk SSSR*, 265, 339-343.
- Sakharov, B.A., Besson, G., Drits, V.A., Kameneva, M.Y., Salyn A.L., Smoliar, B.B. (1990): X-ray study of the nature of stac-

king faults in the structure of glauconites. *Clay Minerals*, **25**, 419-435.

- Sassi, P.F., Guidotti, C., Rieder, M., De Pieri, R. (1994): On the occurrence of methamorphic $2M_1$ phengites: some thoughts on polytypism and crystallization conditions of 3T phengites. *Eur. J. Mineral.*, **6**, 151-166.
- Tsipursky, S.I. & Drits, V.A. (1984): The distribution of octahedral cations in the 2:1 layers of dioctahedral smectites studied by oblique texture electron diffraction. *Clay Minerals*, **19**, 177-193.
- Warshaw, C.M. (1959): Experimental studies of illites. *Clays and Clay Minerals*, 7, 303-316.
- Ylagan, R.F., Altaner, S.P., Pozzuoli, A. (2000): Reaction mechanisms of smectite illitization associated with hydrothermal alte-

ration from Ponza island, Italy. Clays and Clay Minerals, 48, 610-631.

- Zvyagin, B.B. (2001): Current problems with the nomenclature of phyllosilicates. *Clays and Clay Minerals*, **49**, 492-499.
- Zvyagin, B.B., Rabotnov, V.T., Sidorenko, O.V., Kotelnikov, D.D. (1985): Unique mica consisting of noncentrosymmetric layers. *Izvestiya Akademii Nauk S.S.S.R, Seriya Geologicheskaya*, 35, 121-124 (in Russian).

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