

**RADIOCARBON DATING AND PALYNOLOGICAL  
STUDY OF SOIL-PYROCLASTIC DEPOSITS AT THE  
FOOT OF KARYMSKII AND MALYI SEMYACHIK  
VOLCANOES**

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This paper presents radiocarbon dates obtained at the Institute of Volcanology and the Geological Institute, USSR Academy of Sciences, from 77 samples of coal, wood, and buried soil collected during a tephrochronological study of soil-pyroclastic deposits at the foot of Karymskii and Malyi Semyachik. The radiocarbon dates are tabulated with references to major events in the eruptive history of the volcanoes and are also indicated on the geologic columns which show the location of the samples used for dating. The accuracy of the dates is discussed. A spore and pollen diagram is presented for the soil-pyroclastic deposits at the foot of Malyi Semyachik. It is considered to be a reference for the East-Kamchatka volcanic zone. Spore and pollen assemblages were identified, dated and tied to the time intervals of the Blytt-Sernander climatic classification.

This paper is a current report on radiometric dating in a series of publications initiated by the Institute of Volcanology in 1984 [6]. The results of radiocarbon age determination for the eruptive products of Karymskii and Malyi Semyachik Volcanoes, as well as tephrochronological data, were published in [2], [3], [8], and [9], where they were used to determine the age and reconstruct the eruptive histories of the volcanoes. Since then, a large number of new radiocarbon dates have been obtained and we deem it reasonable to summarize the data collected during this stage of the work.

Karymskii and Malyi Semyachik Volcanoes are situated in the East-Kamchatka volcanic zone at a distance of 20 km from each other. The Karymskii volcanic structure is a normal cone located in a young Holocene caldera. Malyi Semyachik is a volcanic ridge consisting of three stratovolcanic cones connected at the base: Paleozoic, Mesozoic and Cenozoic Semyachik cones.

The geologic materials used for radiocarbon dating were pyroclastic-soil deposits at the foot of the volcanoes which record a sequence of explosive events. The pyroclastic-soil deposits consist of tephra beds with sandy loams and buried soils in-between. Dating was undertaken to determine the time of the volcanic events recorded by the tephra. We also determined the age of the pyroclastic flows which had been related to the formation of the Karymskii caldera and to the eruptive activity of Cenozoic Semyachik.

The source materials for carbon-14 dating were buried wood, coals, and soils. Samples were collected in different years by O.A. Braitseva, I.A. Egorova, S.N. Litasova, and L.D. Sulerzhitskii. A sample of carbon extracted from organic matter buried in volcanic ash (Sample GIN-1066 in Table I) was made available to us by G.N. Kovalev. Radiocarbon dating was performed at the Geological Institute, USSR Academy of Sciences (samples labeled GIN), and at the Institute of Volcanology, Far East Division, USSR Academy of Sciences (samples labeled IV AN).

Most of the samples were buried soils. Their ages were determined using successive alkali extracts from one and the same sample, which contained organic remains of different ages and gave an idea of the beginning and end of soil formation. This technique, however, could only be employed with carbon-rich samples. Besides, use was made of wood buried in soil or in tephra above it. Also, coal samples were collected from the material deposited by pyroclastic flows.

Figures 1 and 2 demonstrate the geologic columns of soil-pyroclastic deposits from which samples were collected for dating, and Figure 3 shows their location. On the columns of Figure 2 the radiocarbon dates are tied to particular soil horizons. The materials used for dating (coal, wood, or buried soil) are marked by symbols. The scale plotted on the left indicates the depth of sampling. The boxed figures are the maximum and minimum ages obtained from the successive alkali extracts from one sample (all ages between these limits are presented in Tables I and II). The constant of 5568 years was used and no corrections were applied. To make correlation easier, the radiocarbon ages are plotted in Figure 1 on the composite sections of the soil-pyroclastic deposits at the foot of Karymskii (I) and Malyi Semyachik (II).

Before considering the results of radiocarbon dating it is worth dwelling on the technique of their interpretation. The normal sequence of dates obtained from buried soils, as they gradually become older downward, and the absence of significant outliers and reversals indicate that these dates are close to the actual radiocarbon age. On this basis, soils from soil-pyroclastic deposits can be recognized as a material suitable for radiocarbon dating, because they were safely conserved under the pyroclastics and some of their horizons had a very short life. Another evidence supporting the supposition that the dating was accurate enough is a satisfactory agreement between the dates for the samples collected from the same soil at several sites located along the strike, sometimes at large intervals between them. This is clearly demonstrated by the dates for the soil horizon between a transit ash bed (3) and a cinder layer at the foot of Karymskii and for soils 4, 6, and 7 at the foot of Malyi Semyachik.

At the same time, one can see that for a sequence of humus horizons in a thick soil the age from the youngest extract for the lowermost horizon is frequently younger than the age from the oldest extract for the uppermost horizon. Examples are the lower horizons of soils 3 and 4 in the Malyi Semyachik column and the soils above and below unit PT<sub>6</sub> and the soils under pyroclastic unit K in the Karymskii column. Obviously, this situation may be the result of the downward migration of young humus, which might be quite significant where humus horizons are separated by a small number of thin ash layers and constitute one fairly thick unit (soils 3 and 4 on Malyi Semyachik and soils in the upper and lower part of the Karymskii column). At the same time, one can see that the dates obtained

**Table I** Radiocarbon Dates for Soil-Pyroclastic Cover at Karymskii Base

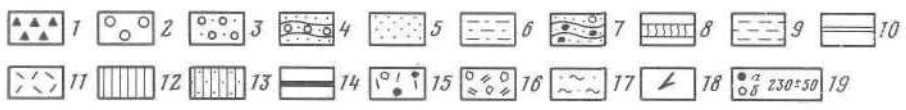
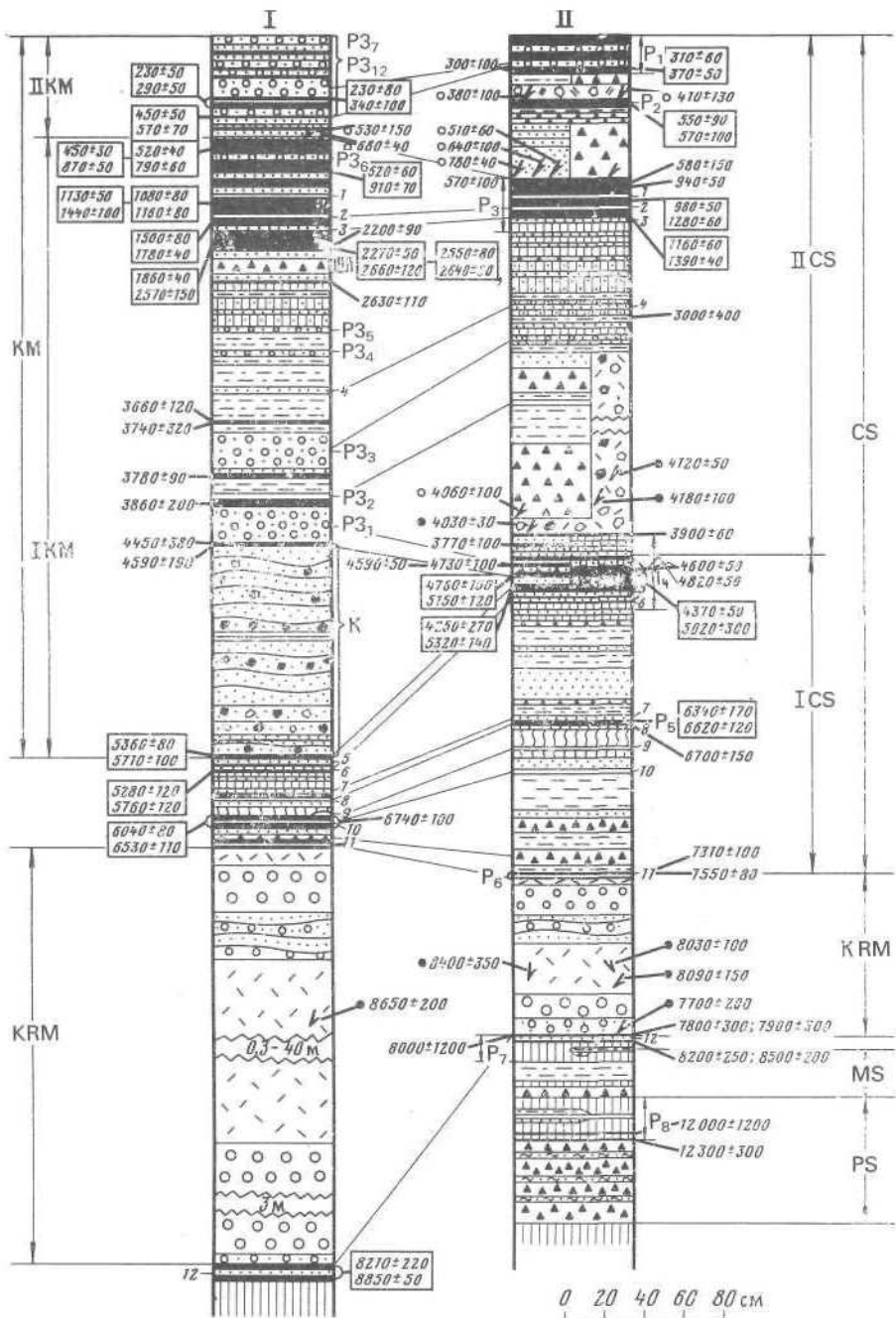
<i>Stratigraphic unit and material sampled for dating</i>	<i>Dated event</i>	<i>Sampl- ing depth</i>	<i>Radiocarbon age, years ago</i>	<i>Sample number</i>	<i>Sampling site number</i>		
Humus horizon under pumiceous tephra PT <sub>11</sub> Humus horizon under pumiceous tephra PT <sub>8-9</sub>	Short breaks between eruptions during II KM active period	0.08	Recent	IV AN-197	336		
		0.14	230±80 <sub>I</sub> 290±50 <sub>II</sub> 340±100 <sub>III</sub>	IV AN-199a IV AN-199b IV AN-199c	336		
Humus horizon under pumiceous tephra PT <sub>7</sub>		0.04	230±50 <sub>II</sub> 290±50 <sub>III</sub>	IV AN-209a IV AN-209b	357		
		0.60	450±50 <sub>I</sub>	IV AN-181a	283		
			570±70 <sub>II+III</sub>	IV AN-181b			
Wood from soil between pumiceous tephra PT <sub>6</sub> and gray-black volcanic sand	Onset of Karymskii activity	0.80	530±50	IV AN-66	283		
		0.90	680±40	IV AN-206	349		
Soil between pumiceous tephra PT <sub>6</sub> and gray- black volcanic sand	Period of subdued activity between IKM and IIKM active periods	0.90	450±30 <sub>II</sub> 770±90 <sub>I</sub>	GIN-1850 GIN-1849	283, sampled in 1972		
			870±50 <sub>III</sub>	IV AN-35		283, sampled in 1978	
		0.86	520±40 <sub>III</sub> 630±40 <sub>II</sub> 790±60 <sub>I</sub>	IV AN-185 IV AN-186 IV AN-184			
Soil under pumiceous tephra PT <sub>6</sub>		1.00	620±60 <sub>II</sub> 680±100 <sub>III</sub> 910±70 <sub>I</sub>	IV AII-188a IV AII-188b IV AII-188c	283		
			0.64	1500±120		IV AII-208	361
				1.24		1080±80 <sub>II</sub> 1160±80 <sub>III</sub>	
Soil between transit ash beds 1 and 2		1.24	1130±50 <sub>III</sub> 1210±40 <sub>I</sub> 1440±100 <sub>II</sub>	IV AH-190 IV AH-189a IV AH-189b	283, sampled in 1978		
			0.26	1200±30		GIH-1167	28
			Soil between transit ash beds 2 and 3			1.34	1500±80 <sub>I</sub> 1760±40 <sub>II</sub>

Table I contd

			1780±40 <sub>III</sub>	IV AH-191c	
Soil between transit ash bed 3 and cinder layer CL		1.40	1860±40 <sub>II</sub>	IV AH-193a	283, sampled in 1978
			2160±60 <sub>I</sub>	IV AH-191b	
			2570±150 <sub>III</sub>	IV AH-191c	
		1.40	2090±60 <sub>I</sub>	GIH-1847	283, sampled in 1975
			2290±110 <sub>III</sub>	IV AH-34	
			2990±60 <sub>II</sub>	GIH-1848	
	1.46	2270±50 <sub>III</sub>	IV AH-5a	283	
		2660±120 <sub>I+II</sub>	IV AH-5b		
		1.44	2200±90 <sub>I</sub>	GIH-1853	278
		1.30	2550±80 <sub>II</sub>	IV AH-177a	285-286
			2640±80 <sub>III</sub>	IV AH-177b	
Humus horizon below cinder layer CL	Short repose period before violent eruption	1.52	2630±110	IV AH-203	343
Soil above pumiceous tephra PT <sub>3</sub>	Repose period after violent eruption which ejected PT <sub>3</sub>	3.00	3660±120	IV AH-476	285
		1.06	3740±320	IV AH-194	336
Soil between PT <sub>3</sub> and PT <sub>2</sub>	Repose period between eruptions which ejected PT <sub>3</sub> and PT <sub>2</sub>	0.70	3780±90	IV AH-480	25
Soil between PT <sub>1</sub> and PT <sub>2</sub>	Repose period between eruptions which ejected PT <sub>1</sub> and PT <sub>2</sub>	1.28	3860±200	GI H-1852	278

Soil below PT <sub>1</sub>		Repose period	1.42	4450±380	IV AH-195	336
		before eruption which ejected PT <sub>1</sub>	1.68	4590±190	IV AH-283	342
Buried soil between Karymskii Volcano (KM) and caldera (KRM) deposits	Humus horizon below pyroclas- tics ejected at Karymskii birth	End of repose	6.90	5360±80 <sub>I</sub> 5710±100 <sub>III</sub>	IV AH-178a IV AH-178b	286
			6.96	5280±120 <sub>I</sub> 5760±120 <sub>II</sub>	IV AH-179a IV AH-179b	286
		Repose period between caldera formation and Karymskii birth	7.06	6040±80 <sub>I</sub> 6280±100 <sub>II</sub> 6530±110 <sub>III</sub>	IV AH-180a IV AH-180b IV AH-180c	286
	Humus horizon below transit tephra 6	3.70	6740±100	GIH-1709	278	
Coals from pyroclastic flow produced during Karymskii caldera formation	Caldera formation	9.00	8650±200	IV AH-519	319	
Soil buried under caldera-related deposits (KRM)		Period preceding onset of caldera- forming eruptions	25, 36	8210±220 <sub>I</sub> 8280±180 <sub>III</sub> 8850±50 <sub>II</sub>	IV AH-1a IV AH-1b IV AH-1c	319

Note: I, II, III in Column 5 are successive alkali extracts from the same soil sample. For location and explanation of symbols see Figure 1.



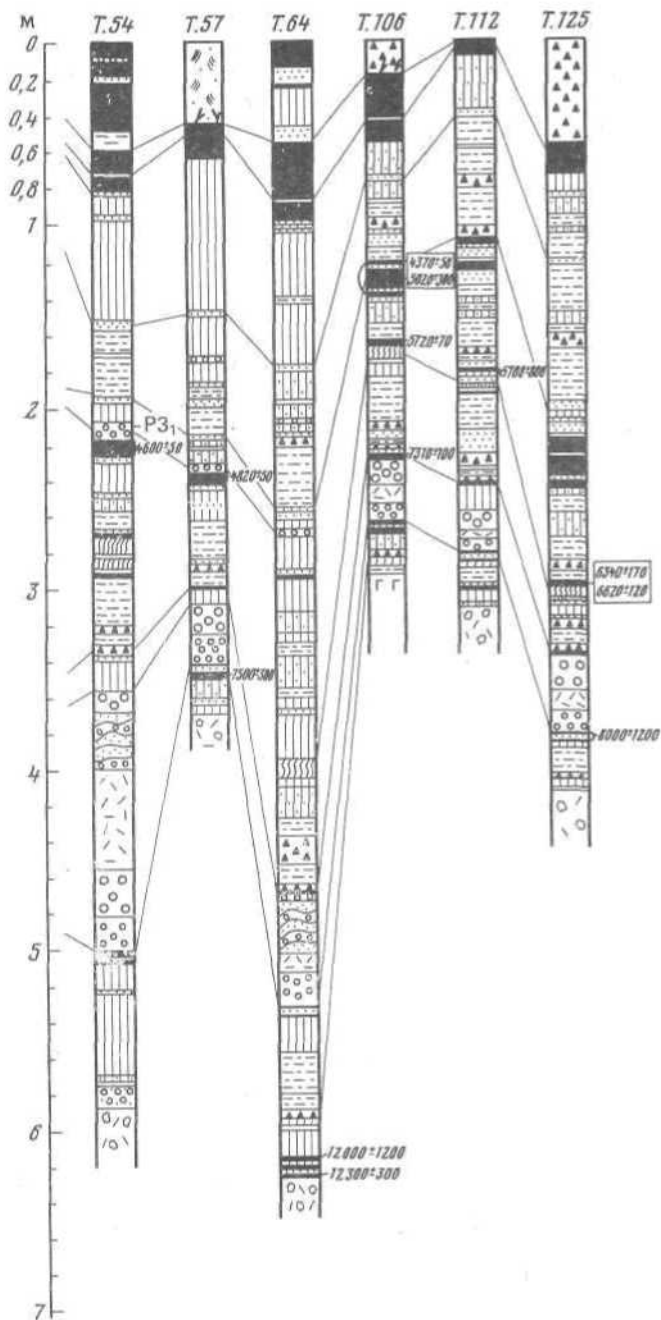
from the oldest extracts for the same humus horizons tend to be definitely older for each successively lower horizon, i.e. the ages of the soil horizons determined from the oldest extracts become progressively older downward.

In some instances, the dates obtained for one and the same soil horizons vary essentially from place to place. For example, two dates of about 5700 years ago and one date of 6700 years ago were obtained earlier [8], [9] for soil 5 at Malyi Semyachik. The age of 5700 years was preferred because it was given by two datings. However, the new data (Sample IV AN-281, Table II) obtained for the same soil after a particularly careful resampling procedure yielded an age close to 6500 years ago. The age of 5700 years seems to have been underestimated because soil 5 was contaminated with young humus. Earlier [9], a date of  $21600 \pm 500$  years ago was reported for the soil at the base of the Malyi Semyachik pyroclastic cover at Site 125, but no geological interpretation was offered. Later this soil was resampled and dated approximately 8000 years ago. This age agrees with the dates obtained for the same soil from the other localities.

Consistent dates were obtained from different wood samples, where the soil from which they were recovered was thin enough (e.g. soil 2 at Malyi Semyachik). The wood samples collected from thick soils (e.g. soil 3 at Malyi Semyachik and the soil including pumiceous tephra unit  $PT_6$  at Karymskil) yielded ages with a spread of 100-300 years. Yet, they fall within the age range obtained for the upper horizon of the soils under consideration (above transit ash unit 1 and pumiceous tephra unit  $PT_6$ ).

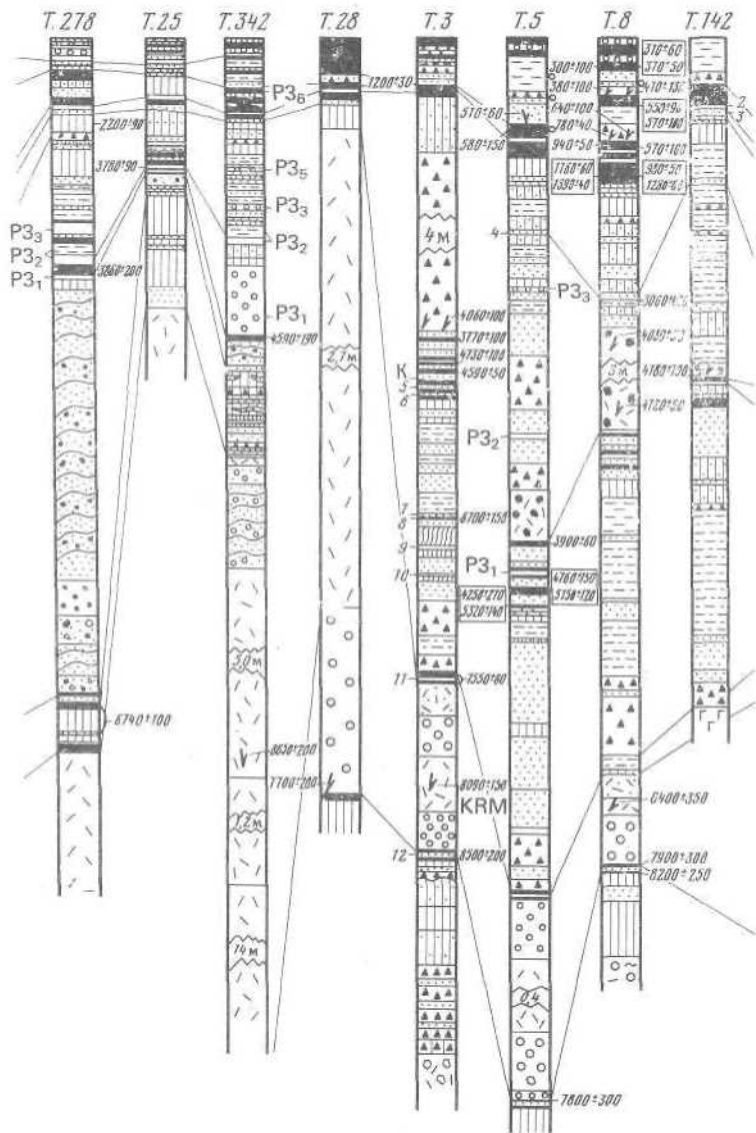
A normal sequence of dates along the column and a good agreement of dates for the same horizons sampled at different sites indicate that the dating was accurate enough and the ages obtained do not differ much from the true radiocarbon ages thanks to the "dead"  $CO_2$  effect of active volcanoes. However, where coals were sampled from pyroclastic flows, dating yielded ages older than those of the underlying soil horizons. Examples are the coals from the pyroclastic flows of the Karymskii caldera dated 8400 and 8600 years with the underlying soil dated 7800-7900 years and the coals from a pyroclastic flow at Cenozoic Semyachik dated 4120 and 4180 years with the underlying soil of 3900 years. Obviously, the overestimated ages resulted from the effect of various amounts of juvenile volcanic carbon entrapped from the carbon dioxide of the gas-rich material of the pyroclastic flows. The age of 7700 years obtained for the coals sampled from the cold-

**Figure 1** Composite sections of soil-pyroclastic deposits at the foot of Karymskii (I) and Malyi Semyachik (II). 1 — bombs, lapilli, volcanic gravel; 2 — bombs lapilli, pumice gravel; 3 — pumiceous lapilli and coarse volcanic gravel; 4 — stratified pumiceous tephra, lapilli, coarse ash; 5 — coarse ash; 6 — stratified coarse ash; 7 — stratified coarse ash with bombs, lapilli, and resurgent material (pyroclastics ejected by early Karymskii eruptions); 8 — "ochreous" bed of oxidized gray and yellow coarse ash; 9 — thin-bedded fine and coarse ash; 10 — fine ash; 11 — pyroclastic flow deposits and fine ash produced by Karymskii eruption responsible for caldera formation; 12 — sandy loam; 13 — sandy loam with coarse ash; 14 — buried soil; 15 — deposits of Cenozoic Semyachik pyroclastic flow; 16 — deposits of explosive origin; 17 — redeposited coarse ash; 18 — wood; 19 — radiocarbon dates: (a) from coal samples, (b) from wood, and remaining — from buried soil. Code: KRM — deposits produced by Karymskii eruptions responsible for caldera formation; KM — deposits ejected by Karymskii eruptions; PS — Paleozoic Semyachik deposits; MS — Mesozoic Semyachik deposits; CS — Cenozoic Semyachik deposits; I KM, II KM and I CS, II CS — deposits of major events in Karymskii and Cenozoic Semyachik activity, respectively. Marker beds:  $PT_1$  to  $PT_{12}$  — Karymskii Pumiceous tephra units; K — early Karymskii pyroclastics; CL — Karymskii cinder layer, 1 to 12 — transit ash beds; Si to S $\alpha$  — buried soils in Malyi Semyachik deposits.



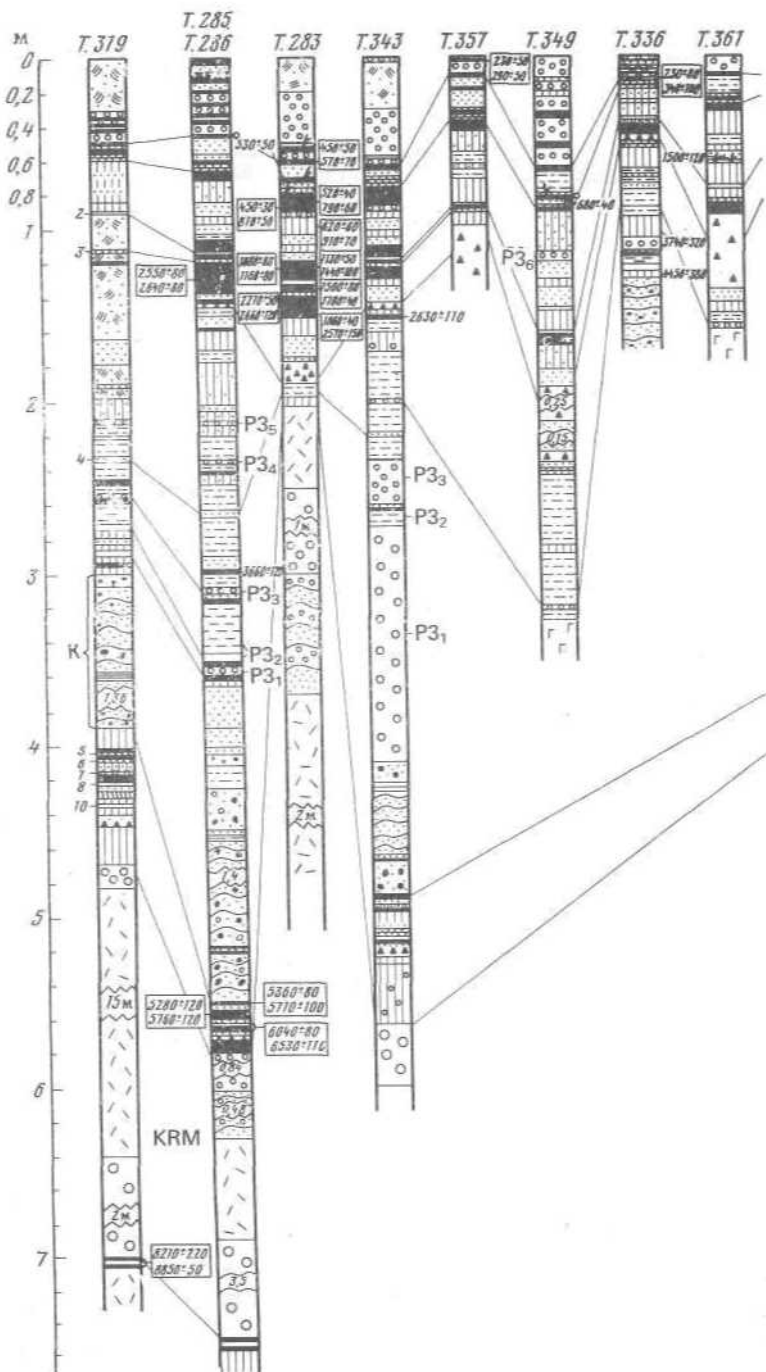
fallen lapilli at the base of the caldera pyroclastic unit is close to the carbon-14 age for the underlying soil horizon. Here, wood was probably coalified when the soil was covered by a thick pyroclastic flow, though heating was not strong enough for





the adsorption of volcanic carbon. It is quite likely that the age of  $1500 \pm 200$  years (see IV AN-202 in Table I) for the coals in the upper soil containing transit ash unit 1 of about 1000 years was overestimated due to the entrapment of "dead" carbon dioxide by the wood as it was coalified.

The tephrostratigraphic correlation of the geologic sections served as a valuable means of checking the accuracy of dating by correlating the ages for the same horizons in many separated areas at the foot of Karymskil and Malyl Semyachik Volcanoes. The soil horizons which were used for dating were identified in the field by tracing ash beds along a profile between the volcanoes. An important



**Figure 2** Soil-pyroclastic columns from the base of Karymskil and Malyi Semyachik. For legend and code see Figure 1.T.54 — sampling site.

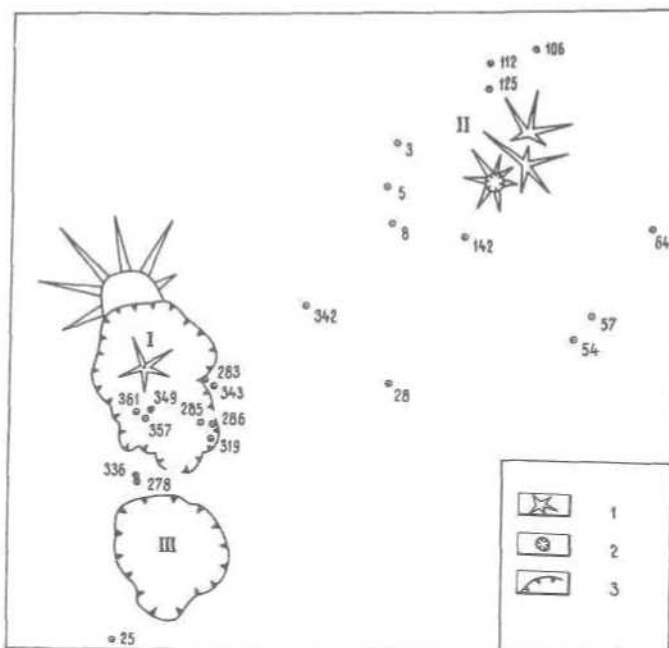


Figure 3 Location of soil-pyroclastic sections sampled for dating. 1 — stratovolcano; 2 — caldera; 3 — caldera wall; I — Karymskii; II — Malyi Semyachik; III — Karymskii caldera lake

feature on which the stratigraphic correlation was based was a sequence of ash beds which could be discriminated by color, thickness, grain size, and chemical and mineral composition. Figure 1 shows correlation of major markers which are present in the tephra of both columns. The markers are transit ash beds and some of the pumiceous tephra horizons at Karymskii and cinders at Malyi Semyachik. The correlation of the dates obtained for the soil horizons mutually correspondent in stratigraphic position (i.e. for the same horizons in two separated areas) showed their fairly good agreement. Examples are the top of soil 3 at Malyi Semyachik and the soil including the  $PT_6$  pumiceous tephra unit at Karymskii, the soil under these horizons, the soil between transit ash beds 1 and 2, the soil under pumiceous tephra unit  $PT_b$ , the soil under pyroclastic unit K, and the soil horizons under the pyroclastic material produced by the explosion responsible for the formation of the Karymskii caldera.

In some cases, the same soil horizons at Karymskii and Malyi Semyachik yielded different ages, e.g. the soil between transit ash beds 2 and 3. The date obtained from the Karymskii soil seems to be closer to the true age because dates of the order of 1400—1500 years were obtained for the soil under transit ash bed 2 (ash from Opala Volcano) elsewhere in the East-Kamchatka volcanic zone [5], [7], [10].

The results of the carbon-14 dating suggest the following ages for some of the transit tephra beds: about 1000 years for bed 1, approximately 1400—1500 years for bed 2 (ash of Opala Volcano), about 1800 years for bed 3 (Ksudach ash),

**Table II** Radiocarbon Dates for Soil-Pyroclastic Cover at Malyĭ Semyachik Base

<i>Stratigraphic unit and material sampled for dating</i>	<i>Dated event</i>	<i>Sampling depth</i>	<i>Radiocarbon date, years ago</i>	<i>Sample number</i>	<i>Sampling site number</i>
Soil 1, lower humus horizon	Onset of relatively subdued activity of Cenozoic Semyachik	0.14	300±100	GIN-1160	8, sampled in 1973
		0.14	310±60 <sub>I</sub> 370±50 <sub>II+III</sub>	IV AH-268a IV AH-268b	8, sampled in 1978
Wood from black volcanic sand above soil 2	Strong eruption of Cenozoic Semyachik	0.32	380±100	GIN-1039	8
		0.34	410±130	GIN-1052	8
Soil 2	Repose period between two strong eruptions	0.36	550±90 <sub>I</sub> 570±100 <sub>III</sub> IV AH-269a IV AH-269b	IV AH-269a IV AH-269b	8
Wood from soil 3 and overlying volcanic sand	Strong eruption of Cenozoic Semyachik	0.50	510±60	IV AH-275	5
		0.60	640±100	GIN-1041	8
Soil 3, upper humus horizon above transit ash 1		0.62	780±40	IV AH-270a	8
		0.62	940±50	IV AH-270b	8
		0.52	580±150	GIN-1368	13
		0.60	570±100	IV AH-270	8
Soil 3, middle humus horizon between transit ash beds 1 and 2	Repose period in Cenozoic Semyachik activity	0.65	1260±70 <sub>I</sub> 1250±60 <sub>III</sub>	IV AH-272a IV AH-272b	8
		0.67	980±50 <sub>II</sub>	IV AH-273a	8
			1060±60 <sub>III</sub>	IV AH-273b	
			1280±60 <sub>I</sub>	IV AH-273c	
		0.70	1160±60 <sub>III</sub> 1300±80 <sub>II</sub> 1390±40 <sub>I</sub>	IV AH-274a IV AH-274b IV AH-274c	8
Plant remains in black volcanic sand	Sequence of	1.46	3000±400	GIN-1066	8

	minor eruptions during relatively subdued Cenozoic Semyachik activity				
Coals from pyroclastic flow	Strong eruptions of Cenozoic	1.7	4030±30	GIN-1040	8
		5.0	4120±50	GIN-1053	9
	Semyachik which produced a pyroclastic flow and cinder ejections in early II CS period	1.84	4180±100	GIN-1051	142
Wood from cinder above soil 4		5.6	4060±100	GIN-1042	3
Soil 4, sequence of humus horizons	Repose period and subdued activity of Cenozoic	2.80	3900±60	IV AH-276	5
		5.70	3770±100	GIN-1043	3
		5.82	4730±100	GIN-1044	3
		5.86	4590±50	GIN-1045	3
	Semyachik in late ICS period	3.06	4760±150 <sub>I</sub>	IV AH-279a	5
			4830±120 <sub>III</sub>	IV AH-279b	
			5150±120 <sub>II</sub>	IV AH-279c	
		2.24	4600±50	GIN-1168	54
		2.44	4820±50	GIN-1170	57
		1.30	4370±50 <sub>I</sub>	GIN-1050a	106
			4800±140 <sub>II</sub>	GIN-1050b	
			5020±300 <sub>III</sub>	GIN-1050c	
		3.16	4250±270 <sub>I</sub>	IV AH-281a	5
		5320±140 <sub>II+III</sub>	IV AH-281b		
Soil 5	Short repose period in the middle of ICS period	1.70	5720±80	GIN-1054	106
		1.88	5700±800	GIN-1374	112
		3.04	6340±170 <sub>I</sub>	IV AH-281a	125
			6420±140 <sub>II</sub>	IV AH-281b	

Table II contd

		6.68	6620±120 <sub>I<sup>II</sup></sub> 6700±150	IV AH-281c GIN-1046	3
Soil 6	End of repose period prior to birth of Cenozoic Semyachik	2.32	7310±100	GIN-1055	106
		7.58	7550±80	GIN-1047	3
Coals from lower KRM lapillic horizon	Eruption responsible for formation of Karymskiĭ caldera	28.0	7700±200	GIN-844	28
Coals from KRM pyroclastic flows			8030±100	GIN-1171	146
		8.10	8090±150	GIN-1162a	3
Soil 7, humus horizon above transit ash 12 Soil 7, humus horizon below transit ash 12	Repose period preceding forma- tion of Karymskiĭ caldera after Mesozoic Semyachik ceased its activity	6.26	7800±300	GIN-1049	13
		7.56	7900±300	GIN-1163	8
		7.60	8200±250	GIN-1164	8
		8.54	8500±200	GIN-1048	3
		3.88	8000±1200	GIN-1371	125
Soil 8, basal humus horizons	Period of subdued activity before Paleozoic Semyachik ceased to erupt	6.20	12000±1200	GIN-1376	64
		6.30	12300±300	GIN-1375	64

Note. I, II, III are successive alkali extracts from the same sample. For location and explanation of symbols see Figure 1.

about 3000 years for bed 4 (Avacha ash), 5500-5600 years for bed 5 (Avacha ash) and bed 6, about 7400 years for bed 11, and approximately 8000 years for bed 12. Bed 10 is the ash produced by Khangar Volcano and has an age older than 6700 years in the Malyt Semyachik column. On Krashennikov Volcano this ash was dated approximately 6900 years [10]. The presence of these markers in the sections of the neighboring areas enables us to use them as time-stratigraphic units for the dating and correlation of different geologic formations and relief forms.

In this paper we report the carbon-14 dates obtained for the sections at the foot of Karymskiy and Malyt Semyachik Volcanoes and for the major eruptive events and repose periods in the activity of these volcanoes dated accordingly on the basis of these ages. As mentioned above, the application of carbon-14 ages for more detailed reconstructions of volcanic histories, such as the determination of the ages of volcanoes and the duration of their active eruptive and repose periods, the dating of most violent eruptions, etc., was discussed earlier in [2], [3], [8], and [9]. To reconstruct the volcanic histories and compare the carbon-14 ages with paleogeographic, paleoclimatic and other data available on the Holocene period, they were calibrated in a conventional manner [1], [12].

Along with the tephrochronological work, we undertook a palynological analysis [4], [9]. We studied pollen and spores in every layer of the soil-pyroclastic cover and found that pollen and spores were present in almost all horizons with the exception of thick interbeds of volcanic sand and lapilli which had been deposited as the result of strong and short-term ejections during volcanic eruptions. The spore and pollen spectra of each horizon are synchronous with the time of its deposition and reflect both the local and regional features of vegetation. Variations in the spore and pollen spectra are associated primarily with climatic changes.

Having analyzed the palynological data, we identified five phases in the evolution of the Holocene vegetation in eastern Kamchatka and, using a great number of radiocarbon dates, determined their age ranges and tied them to the periods of the Blytt-Sernander climatic classification\* (Figure 4). These phases are: phase I — the growth of grass and bush communities with the local distribution of brushwoods 10300—9200 years ago; phase II — the growth of creeping alder with minor birch forests 9200-6000 years ago; phase III — the Holocene peak of birch forest development with the general predominance of bush communities of creeping alder and Japanese stone pine 6000-4500 years ago; phase IV — the dominant development of bush communities of creeping alder and Japanese stone pine 4500-2500 years ago; and phase V — the growth of birch forests with the general predominance of bush communities of creeping alder and Japanese stone pine 2500 years ago to the present time.

The identified phases of the vegetation evolution are characterized by certain spore and pollen assemblages and served as the basis for differentiating the eastern Kamchatka Holocene deposits into five stratigraphic units each of which was tied to the absolute time scale. The characteristic features of the spore and pollen assemblages are: an insignificant content of bush pollen and a predominance of the *Lycopodiaceae* spores in the cold periods of the late glacial stage and the

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\* We referred to the Holocene subunits of the Blytt-Sernander classification as modified by Khotinskii for North Eurasia [11].

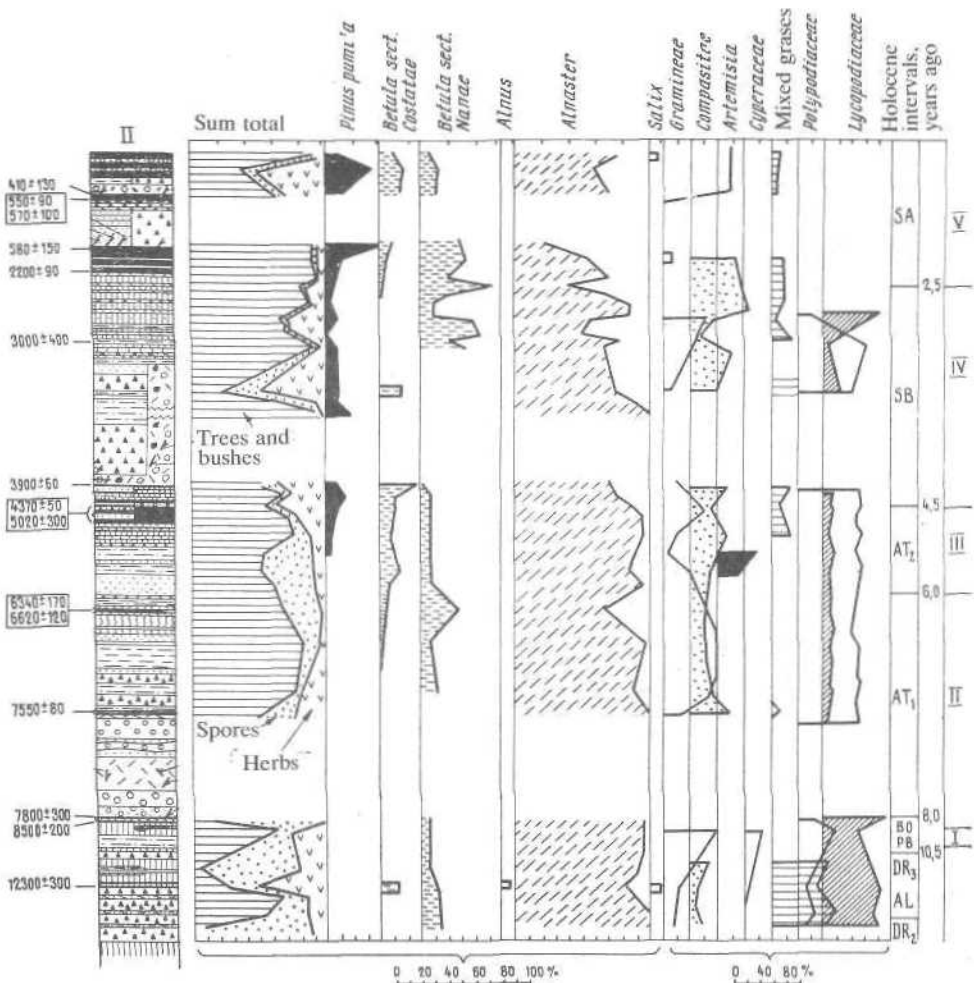


Figure 4 Composite spore and pollen diagram for soil-pyroclastic deposits at Malyl Semyachik base. SA - Subatlantic; SB - Subboreal; AT - Atlantic; BO - Boreal; PB - Preboreal; AL - Allerod; DR — Dryas. I to V — phases of vegetation evolution

early Holocene; an absolute domination of creeping alder pollen (*Alnaster*) during the Boreal interval (BO) and the first half of the Atlantic (AT<sub>1</sub>); the appearance of birch tree pollen (*Betula sect. Costatae*) and Japanese stone pine pollen (*Pinus pumila*) in the second half of the Atlantic (AT<sub>2</sub>); a decrease and disappearance of birch tree pollen during the Subboreal climatic deterioration (SB); and a new increase in the content of birch tree pollen in the spectra of the Subatlantic interval (SA).

The most prominent stratigraphic unit corresponds to the Holocene climatic optimum which occurred 4500-6000 years ago and was characterized by an abundance of birch forests.



The lower Holocene boundary was determined by the beginning of bush growth after a climatic deterioration at the end of the postglacial interval (Dryas-3). It was placed at 10.3—10.5 thousand years, which agrees with the universally adopted time. The stratigraphic and pollen data indicated a climatic amelioration of late-glacial time corresponding to the Allerod (12000 years ago).

The dated spore and pollen diagram (see Figure 4) for the foot of Malyy Semyachik Volcano is a reference for the extensive region of the East-Kamchatka volcanic zone. The pollen assemblages identified in this research and tied to the time scale can be used as a basis to correlate geologic sections in the region and solve an inverse problem — date the soil-pyroclastic covers where carbon-14 dating is precluded by the absence of organic matter [3]. Also, the palynological data will be instrumental in reconstructing the paleogeographic environments that existed during the birth of the volcanoes.

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