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THE PRE-VENDIAN (640–610 MA) GRANITE MAGMATISM IN THE CENTRAL TAIMYR FOLD BELT: THE FINAL STAGE OF THE NEOPROTEROZOIC EVOLUTION OF THE SIBERIAN PALEOCONTINENT ACTIVE MARGIN

A. B. Kuzmichev¹, M. K. Danukalova¹, V. F. Proskurnin²,
A. A. Bagaeva², N. I. Beresyuk², P. A. Gromov²

¹ Geological Institute of RAS, Moscow, Russia

² A.P. Karpinsky Russian Geological Research Institute (VSEGEI), Saint Petersburg, Russia

Abstract: Eastern part of the Central Taimyr belt is composed of Precambrian rocks penetrated by granites of the Snezhnaya complex (845–825 million years) and later overlain by mid-Neoproterozoic sin- postorogenic sedimentary deposits of the Stanovaya-Kolosova Group. Two competing concepts on the Precambrian history of the belt are discussed. The first suggests that by the middle of the Neoproterozoic amalgamation of various terrains formed the Central Taimyr microcontinent, which afterwards collided with Siberia in Vendian. 2) According to the second point of view, which is shared by the authors of this article, the belt was part of the Siberian craton from at least the Mesoproterozoic, and there is no suture that would separate it from the South Taimyr belt. To our surprise, during the field work in the South-Eastern part of the Central Taimyr belt near the proposed “Vendian sutura”, assumed by the first concept, we found a granite pluton (Pregradnaya massif) intruding clastic rocks of Stanovaya-Kolosova Group. Such setting is quite uncommon for the belt and contradicted to publications, describing the mentioned clastic rocks to overlay the granites and contain their debris. Dating of the pluton confirmed the field observations – its SRIMP zircon age has proved to be 609±2 Ma, an unusually young for this region. The pluton is located in a wide deformation zone separating the Precambrian rocks (to the northwest) and the Paleozoic deposits (to the southeast). Two minor bodies of similar porphyritic granite were found in the same zone further to the southwest, and it seemed logical to assume that a chain of Vendian granites marks boundary deformation zone. However, their dating (843±6 и 840±5 Ma) showed that they belong to Snezhnaya complex. In this paper, we discuss two Neoproterozoic magmatic ‘flare-ups’ in the Central Taimyr Belt, which are dated at 845–825 and 640–610 Ma. Both ‘flare-ups’ are evidenced by K-rich peraluminous granite batholiths intruded the upper crust. It is most probable that each flare-up was related to a collision event completing an independent cycle in the evolution of the active margin of the Siberian paleocontinent.

Key words: Precambrian; Taimyr; Neoproterozoic; collisional granite; geochronology; Siberia Palecontinent

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ПРЕДВЕНДСКИЙ (640–610 МЛН ЛЕТ) ЭТАП ГРАНИТНОГО МАГМАТИЗМА В ЦЕНТРАЛЬНО-ТАЙМЫРСКОМ СКЛАДЧАТОМ ПОЯСЕ: ЗАВЕРШАЮЩАЯ СТАДИЯ ЭВОЛЮЦИИ АКТИВНОЙ ОКРАИНЫ НЕОПРОТЕРОЗОЙСКОГО СИБИРСКОГО ПАЛЕОКОНТИНЕНТА

А. Б. Кузьмичев¹, М. К. Данукалова¹, В. Ф. Проскурнин²,
А. А. Багаева², Н. И. Березюк², П. А. Громов²

¹ Геологический институт РАН, Москва, Россия

² Всероссийский научно-исследовательский геологический институт им. А.П. Карпинского, Санкт-Петербург, Россия

Аннотация: При полевом изучении докембрийских пород восточной части Центрально-Таймырского пояса обнаружен гранитный плутон, прорывающий синорогенные обломочные породы (становско-колосовская серия, предположительно 780–750 млн лет). Такие взаимоотношения нетипичны для пояса: считается, что наоборот, указанные орогенные накопления содержат обломки подобных гранитов. Датирование массива подтвердило наблюдавшиеся взаимоотношения: возраст гранитов оказался равным 610 млн лет. Столь молодые граниты ранее здесь не отмечались. Массив приурочен к широкой зоне деформаций, которая отделяет докембрийские породы (к северо-западу) от выходов деформированного палеозойского чехла (с юго-востока). На простирании в этой зоне были найдены еще два гранитных массива, также сложенные калиевыми порфиридовидными гранитами. Их датирование (840 млн лет), однако, показало, что они относятся к распространённому в этой части Таймыра более древнему (снежинскому) комплексу. В статье обсуждаются две вспышки гранитного магматизма, проявившегося в Центрально-Таймырском поясе в интервалах 845–825 и 640–610 млн лет, выраженные во внедрении калиевых слабопералюминиевых гранитных батолитов в верхние горизонты коры. Каждая вспышка завершает самостоятельный этап эволюции активной окраины Сибирского палеоконтинента в неопротерозое и предположительно связана с коллизионным событием. Породы второго этапа эволюции активной окраины преимущественно выходят на северо-восточной окраине Таймыра, где широко распространены надсубдукционные базальты, андезиты и риолиты позднего неопротерозоя (750–610 млн лет). Вулканиды несогласно залегают на породах мезопротерозоя, нижнего неопротерозоя, а также в синорогенных толщах середины неопротерозоя и несогласно перекрыты пестроцветными конгломератами терминального неопротерозоя.

Ключевые слова: докембрий; Таймыр; неопротерозой; коллизионные граниты; геохронология; Сибирский палеоконтинент

1. INTRODUCTION

The Central Taimyr fold belt is composed mainly of the irregularly metamorphosed and deformed Precambrian rocks. At the north, it is bounded by the Main Taimyr and Diabase thrust faults (Fig. 1) [Bezzubtsev et al., 1983]. They were interpreted as a suture between the Central Taimyr Belt and the Kara microcontinent collided in the late Paleozoic [e.g., Zonenshain et al., 1990; Vernikovskiy et al., 1996]. The southern boundary is the Pyasina-Faddey fault separating the Central- and South-Taimyr belts (see Fig. 1). It is widely accepted that the South Taimyr Belt is composed of deformed rocks of the Siberian platform. In some publications, the Pyasina-Faddey thrust is also interpreted as a suture of Cretaceous [Zonenshain et al., 1990; Uflyand et al., 1991] or Vendian [Vernikovskiy, Vernikovskaya, 2001, and their other publications] age. According to the ci-

ted geologists, the Central Taimyr Belt is composed of set of the island-arc, continental and oceanic terranes amalgamated to a solid agglomerate in the Neoproterozoic and thus formed a microcontinent. The time of the accretion-collision was determined by the age of granites and volcanic rocks at 869–823 Ma [Vernikovskiy, Vernikovskaya, 2001; Proskurnin et al., 2014]. In the mid-Neoproterozoic orogeny was completed by accumulation of clastic rocks (containing pebbles of above granites) overlaid with carbonate rocks and shales [Bezzubtsev et al., 1986; Makariev, 2013]. According to [Vernikovskiy, 1996] and his later publications, the Central Taimyr 'accretionary belt' collided with the Siberian paleocontinent in the Vendian or prior to the Vendian, and the Pyasina-Faddey thrust is a suture.

In 2016, we participated in the VSEGEI field survey of the eastern part of the Central Taimyr Belt (see the box in Fig. 1). The studied area includes high-grade

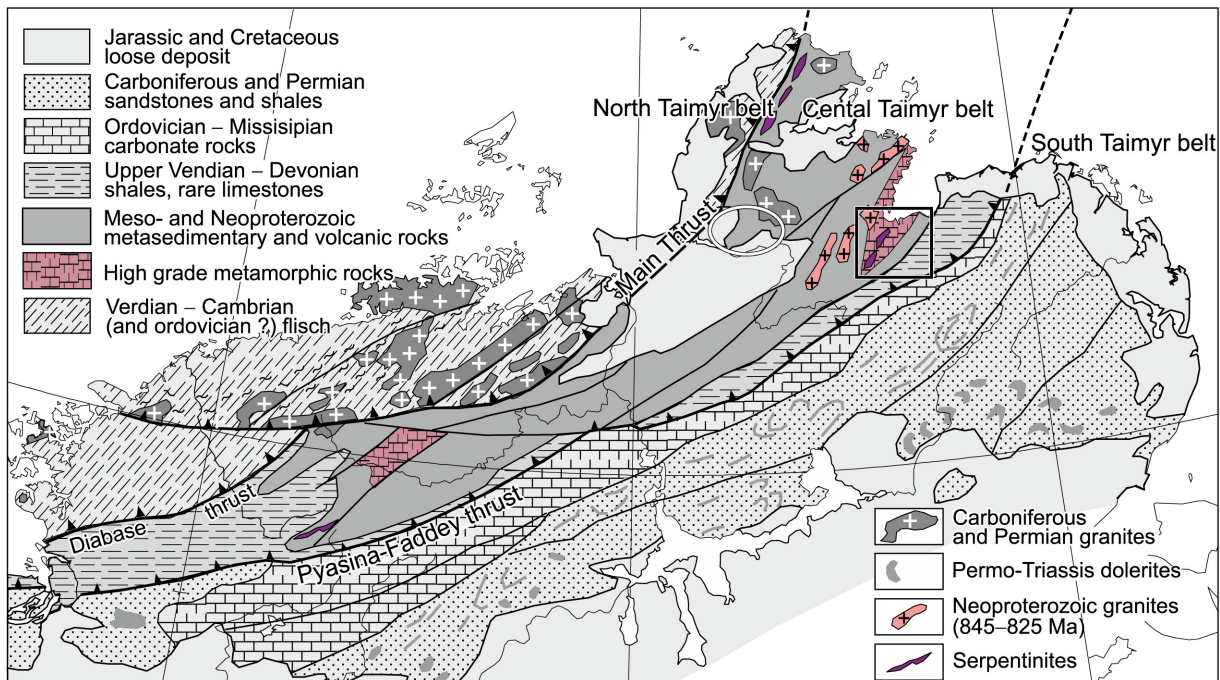


Fig. 1. Position of the Central Taimyr belt in the structural zoning of the Taimyr folded region (after [Bezzubtsev et al., 1983]). Box indicates the study area discussed in this paper; white oval – location of the Nansen (Vil'kitsky) granite suite (see Discussion).

Рис. 1. Положение Центрально-Таймырского пояса в структурной зональности Таймырской складчатой области (по [Bezzubtsev et al., 1983]). Квадратиком показано положение участка, обсуждаемого в данной статье, белым овалом показана область распространения гранитов нансеновского комплекса (см. раздел «Обсуждение»).

metamorphic rocks interpreted in different ways. They were attributed to “Faddey uplift” of the Archean or Paleoproterozoic basement of the Siberian craton [Pogrebitsky; 1971; Zabiya et al., 1986, Bezzubtsev et al., 1986], or the Faddey cratonic terrane being a part of the accretionary belt [Vernikovskiy, Vernikovskaya, 2001; Proskurnin et al., 2014]. V.V. Bezzubtsev did not exclude that the protolith of Faddey metamorphic rocks was Riphean metasedimentary rocks that are widespread in the Central Taimyr belt [Bezzubtsev et al., 1986].

Among other interesting field observations, we found a granitic pluton intruding the above-mentioned synorogenic mid-Neoproterozoic clastic rocks. The discovery could have important implications for understanding the tectonics of Taimyr, as it was consistent with the concept of the Vendian collision of the Central Taimyr belt and the Siberian paleocontinent.

2. OVERVIEW OF THE GEOLOGY OF THE STUDY AREA

Our field observations had shown that in the eastern part of the Central Taimyr Belt no terrain can be delimited, which could be interpreted as an exotic “cratonic terrane” or an ancient basement uplift. Three zones of metamorphism can be roughly outlined (Fig. 2): (i)

the zone of migmatized rocks (anatexis zone) that may correspond to high amphibolite facies and, probably, granulite facies (index minerals are lacking); (ii) the intermediate zone of amphibolite and epidote-amphibolite metamorphism; and (iii) the outer zone of greenschist and epidote-amphibolite metamorphism. Rocks that have experienced mylonitization and retrograde metamorphism are abundant in this area.

In all the three zones metamorphic rocks are represented by metasedimentary crystalline schists and gneisses interlayered with quartzites and marbles and containing amphibolite bodies. On the sites of relatively weak metamorphism, metasedimentary strata were mapped as parts of the October and Zhdanov formations [Bezzubtsev et al., 1983; Zabiya et al., 1986; Markariev, 2013]. The age range of detrital zircons from quartzites sampled from the anatexis zone is similar to that of the October and Zhdanov Formations, and thus supports the above-mentioned assumption of V.V. Bezzubtsev ([Kuzmichev, Danukalova, 2018], and unpublished data of the authors). The detrital zircons are mostly dated to the Paleoproterozoic. The youngest detrital zircons mark the lower age limit for sedimentation at ~1600 Ma ([Proskurnin et al., 2009], and unpublished data of the authors). The upper age limit (1300–1360 Ma) is determined by the age of the intruded metadolerite sills of the North Byrranga complex

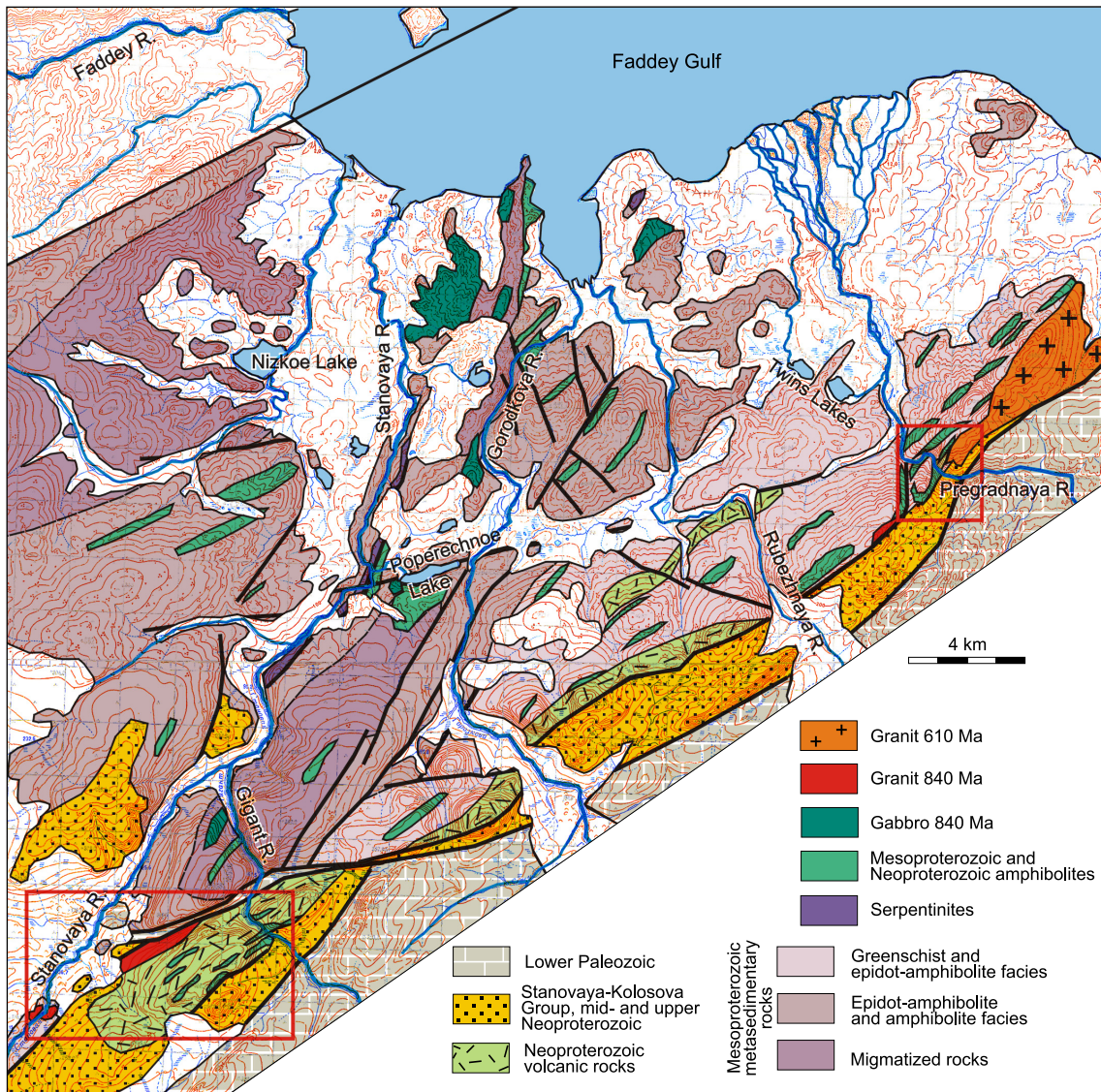


Fig. 2. Schematic geological map of the SE part of the study area. The map is based on the field observations by first two authors, the Planet satellite images (URL: <https://www.planet.com>) and geological maps compiled by [Stepanov et al., 1965; Makariev, 2013]. Rectangular boxes are the areas wherein the discussed granites are located (see more detailed maps in Figs. 3 and 5).

Рис. 2. Схема геологического строения юго-восточной части изученной площади (составлена первыми двумя авторами по полевым наблюдениям с использованием космических снимков Planet (URL: <https://www.planet.com>) и предшествующих геологических карт [Stepanov et al., 1965; Makariev, 2013]). Прямоугольниками выделены участки, включающие обсуждаемые гранитные тела и иллюстрированные более детальными схемами.

([Proskurnin et al., 2014; Priyatkina et al., 2017], and unpublished data of the authors). Similar ages were determined for detrital zircons from the October Fm. sampled further to the west (upper reaches of the Leningradskaya River) outside the Faddey metamorphic complex [Proskurnin et al., 2009] and directly westward of our study area [Priyatkina et al., 2017]. Convincing arguments for the Anabar origin of the Paleoproterozoic detrital zircons are given in the last cited paper. Our data support the conclusion of N.S. Priyatkina and her colleagues that during Meso- and Neopro-

terozoic, the southern part of the Central Taimyr belt belonged to the Siberian paleocontinent and was not separated from it by an ocean.

The early Neoproterozoic volcanic (mostly felsic) rocks occupy the next structural level. Their stratigraphic contact with the Mesoproterozoic rocks has never been observed. Volcanic rocks are partially comagmatic to the Snezhnaya granites (see below).

The Neoproterozoic granites (970–800 Ma [Kuzmichev, Danukalova, 2018]) are represented by different facies reflecting different depth levels. Migmatites and

larger granitoid veins located in the anatexis zone. In the transition zone, there are leucogranitic dikes, generally represented by fine-crystalline rocks (not shown in Fig. 2). The zone of low-grade metamorphism contains similar dikes, small granitic bodies and relatively large medium- to coarse-crystalline granitic plutons (see Fig. 1), typically with K-feldspar phenocrysts. Granites of the latter type are shown in the maps of the Central Taimyr belt as the Snezhnaya complex [Proskurnin *et al.*, 2009; Proskurnin *et al.*, 2014]. Available zircon datings listed in these publications (ID TIMS and SHRIMP) are as follows: 846 ± 11 , 833 ± 14 , 824 ± 3.4 , 827 ± 22 , 830 ± 5.3 , 833 ± 5 , and 843 ± 5 Ma. Thus, the age of the Snezhnaya complex is in the range of 845–825 Ma.

All the above-mentioned rocks are unconformably overlain by mid-Neoproterozoic syn- and post-orogenic deposits of the Stanovaya-Kolosova group (more than 1000 m). It comprises conglomerates and sandstones at the bottom, shallow-water carbonate rocks in the middle, and grey-coloured shales in the upper part. The group is fragmented in the study area and its complete succession is not available. The Stanovaya-Kolosova rocks are deformed and metamorphosed to lower greenschist facies (chlorite zone). To the north of the study area, these rocks are overlain by suprasubduction volcanic rocks of the late Neoproterozoic [Makariev, 2013; Markovsky *et al.*, 2000]. All the Precambrian rocks are unconformably overlain by non-metamorphosed, strongly deformed uppermost Vendian – lower Paleozoic shales and limestones of the Siberian platform.

Deciphering of the complex fold-thrust structure of the study area is possible only for the late Vendian – early Paleozoic structural level. Deformations of Stanovaya-Kolosova mid-Neoproterozoic formation are far more complex. In detailed satellite images only fold hinges composed of contrasting carbonate and shale rocks are visible. These rocks are generally form separate lense-shaped tectonic blocks, each representing only a small fragment of the Stanovaya-Kolosova group. Delineation of the Mesoproterozoic – early Neoproterozoic rock structure can hardly be practicable. In the anatexis zone, the rocks have experienced complex deformation in the plastic state. The satellite images show the general direction of foliation, which is displaced by faults of different orientations. It may be suggested that general structural motif is a set of curved reverse faults concordant with the overall structural trend. The deep-seated rocks are exhumed due to displacements along these faults. The youngest normal faults and right-lateral strike-slip faults (sub-latitudinal and ENE strike) are clearly visible in the satellite images and even on topographic maps. Displacements along these faults continued in the Quaternary.

In the Neoproterozoic, the eastern part of the Central Taimyr belt has experienced orogeny twice in Neo-

proterozoic: at the end of the first half and at the end of the second half. The first event was a full-scale orogeny, including granite batholith intrusion and the accumulation of the Stanovaya clastic deposits. The second event is mentioned in [Vernikovskiy, 1996; Vernikovskiy, Vernikovskaya, 2001; Proskurnin *et al.*, 2014] and discussed in more detail below.

In the area shown in Fig. 2, the only relatively large granitic body is located on the right bank of the lower Pregradnaya River. According to our observations, a unique feature of these granite is its intrusion into the conglomerates of the Stanovaya Formation, i.e. these granites are younger than the above-mentioned Snezhnaya granite suites. Such setting is quite uncommon for the belt, and our discovery was unexpected and contradicting to the previous works, describing the conglomerates overlaying granites and containing granite boulders. It seemed reasonable to relate this granites to a wide tectonic zone bordering the Precambrian rocks (see Fig. 2). Further southwestward, in the same zone (at the upper reaches of the Stanovaya River), two smaller bodies of porphyritic granites similar to the above-mentioned were found, and we had no doubts that they belong to the same young suite. These three granite massifs (Pregradnaya, Stepanov and Magistral'ny) are described and discussed below.

3. PREGRADNAYA MASSIF

Massif is located on the right bank of the Pregradnaya river in its lower reaches (see Fig. 2) and extends to the northeast in accordance with general structural trend. Its area (approx. 12 km²) is clearly detectable by aerial gamma-ray imaging. Several small porphyritic granite bodies were also observed on the left bank of the Pregradnaya River. They are exposed along the fault zone that separates the Precambrian metamorphosed rocks and the Stanovaya-Kolosova Group. It is probable that these are fragments of the SW continuation of the Pregradnaya massif.

The part of the massif exposed on the slopes facing the Pregradnaya River, has been studied in more detail (Fig. 3). This area is a part of a wide fault zone separating the Precambrian and Paleozoic rocks. The Precambrian rocks are mylonitized, and almost all lithological contacts are faulted. Tracing of the most recognizable rocks (e.g., dolomites of the Kolosova Fm.) suggests that, area is cut out to number of lense-shaped blocks by curved reverse faults. These faults are relatively young, considering that the Stanovaya-Kolosova rocks are displaced by them. The dip and strike data shown in the map (see Fig. 3) reflect the actual orientation of metasedimentary rock layering that coincides with foliation in this area. Observations of the meta-dolerites show that banding, foliation and rock contact

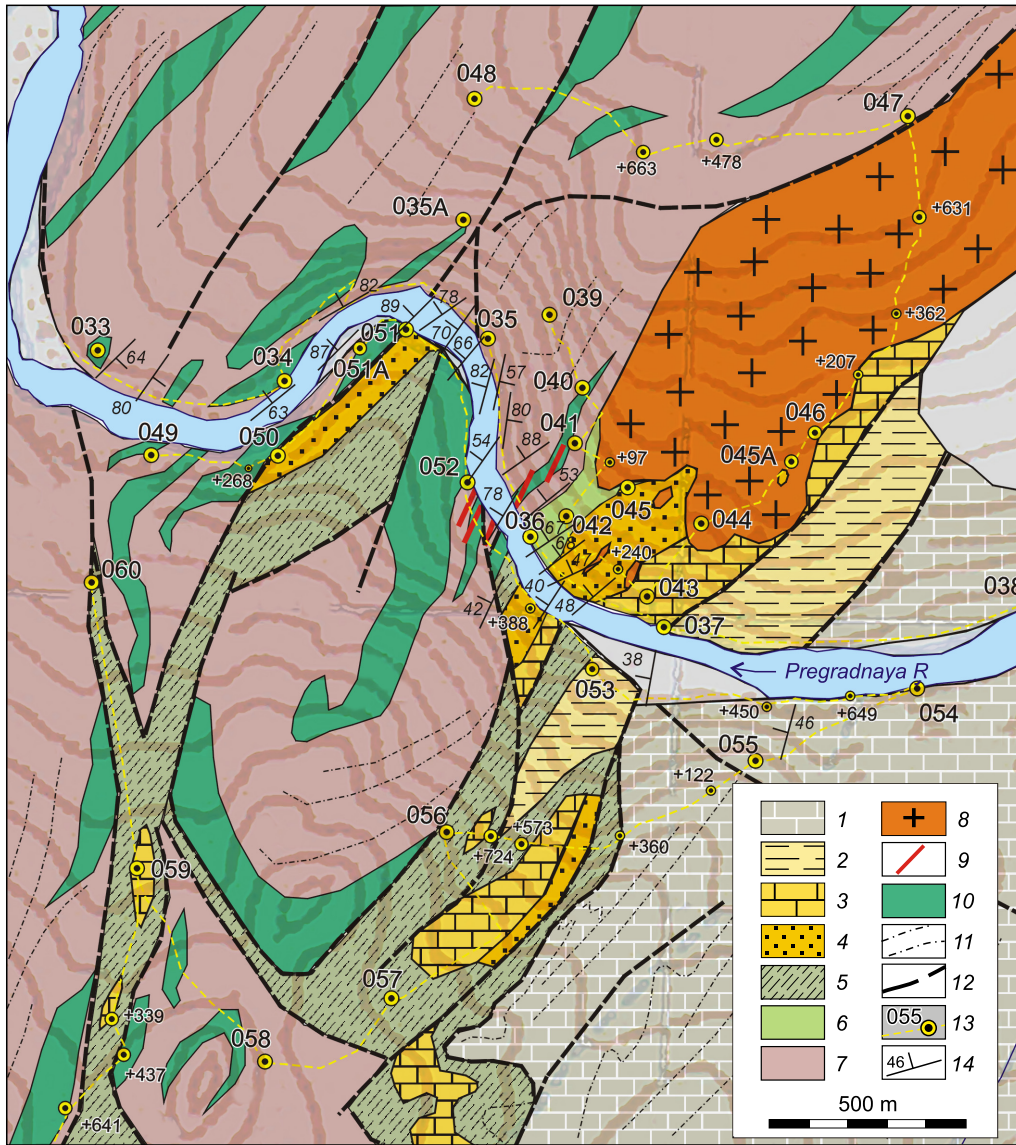


Fig. 3. Geological map of the southwestern termination of the Pregradnaya massif (mapped by first two authors). Contour interval – 20 m.

1 – uppermost Vendian – lower Paleozoic deposits; 2–4 – Stanovaya-Kolosova group (mid-Neoproterozoic): 2 – grey shales, 3 – marbles, 4 – conglomerates, gritstones, marbles; 5 – chlorite and chlorite-sericite green schists, rare gritstones; 6 – felsic metavolcanic rocks (853±11 Ma); 7 – Zhdanova Formation (Mesoproterozoic) – metasedimentary rocks including marbles and quartzites; 8 – Pregradnaya granite (609±2 Ma); 9 – leucogranite dikes; 10 – Mesoproterozoic (1342±14 Ma) amphibolites; 11 – structural lines; 12 – faults; 13 – fixed sites and tracks; 14 – dip and strike of layering, banding, foliation (all the elements have the same orientation).

Рис. 3. Геологическая карта ЮЗ замыкания Преградненского массива (составлена по наблюдениям первых двух авторов). Сечение горизонталей – 20 м.

1 – Верхневендские – нижнепалеозойские отложения; 2–4 – становско-колосовская серия (средний неопротерозой): 2 – серые сланцы, 3 – мраморы, 4 – конгломераты, гравелиты, мраморы; 5 – хлоритовые и хлорит-серицитовые зеленые сланцы, редко – гравелиты; 6 – метавулканыты (853±11 млн лет); 7 – ждановская толща (мезопротерозой) – метасадочные породы, в том числе мраморы и кварциты; 8 – граниты Преградненского массива (610 млн. лет); 9 – дайки лейкогранитов; 10 – мезопротерозойские (1342±14 млн лет) амфиболиты; 11 – структурные линии; 12 – разломы; 13 – фиксированные точки наблюдений и треки; 14 – элементы залегания слоистости, полосчатости, сланцеватости (ориентировка всех перечисленных элементов совпадает).

orientations are coincident wherever such features are observed. Older rocks in Fig. 3 are replaced with younger rocks in SE direction, while most of them are inclined in the opposite direction. Presumably they are predominantly in overturned position. The lower rea-

ches of the Pregradnaya River is the only place in the region, wherein the rock relationships can be observed on the river bank cliffs.

The geological units outlined in the Fig. 3 are as follows: (i) Mesoproterozoic metasedimentary rocks. Pre-

viously, these rocks were mapped as the Zhdanov Fm. due to the presence of carbonate rocks. Carbonate rocks (10–15 % of the visible thickness) represented by grey marble layers (from few centimeters to several meters thick) interbedded with various schists, mica-quartzites and massive quartzites. Widespread are carbonate-silicate rocks. Due to intense deformation, an original lithological succession cannot be restored. (ii) Mesoproterozoic amphibolites of the North Byrranga complex. They form concordant bodies of a few meters up to several dozens of meters thick interpreted as sills. These are massive, slightly schistose, rarely banded black-green rocks composed of hornblende, albite and quartz with varying amounts of Fe-Ti minerals. Indistinct gabbro-dolerite structure is preserved in some sites. (iii) Neoproterozoic metavolcanic rocks with cloudy microcrystalline basis and, vague remnants of feldspar and rare quartz phenocrysts. In places they intercalate with packs of chlorite schists. (iv) Rocks of the Stanovaya-Kolosova Group. These rocks are believed to lie unconformably upon rocks listed above, but the direct contact in the study area is faulted. The lower part of the group comprises conglomerates, gritstones and sandstones interbedded with marbles. Coloured sandstones and gritstones on the left bank of the Pregradnaya River (near point 051) probably compose the lowermost horizons of the Group, which are not preserved on the right bank. These rocks associate with hematitized green schist interpreted as pre-Stanovaya weathering crust. Middle part of the Group consists of marbles. Upper part comprises grey-coloured shales and siltstones with rare sandstones. This shale unit is distinctly mappable several kilometers to the southwest, here only its fragments preserved. (v) All of the above-mentioned rocks are intruded by the Pregradnaya granites (discussed in more detail below). A set of light coloured, almost white granite veins that probably acted as feeder dikes.

Pregradnaya granites crop out at cryoplanation terraces covered with granite blocks (Fig. 4, *a*) and in scarps several meters high. A contour of the southwestern margin of pluton is sinuous. Here granite intrudes all pre-Stanovaya rocks, as well as Stanovaya conglomerates and marbles. Intrusive contacts can hardly be substantiated confidently considering the 'talus' or 'rubble' setting. Nonetheless, all of the above-mentioned host rocks (more or less altered) are observed in xenoliths at the massif's margin. Among them Stanovaya conglomerates are the most detectable. The Pregradnaya massif is composed of medium- to coarse-grained porphyritic granite with yellowish and pink K-feldspar phenocrysts (1–3 cm). Phenocrysts are roughly rounded and less commonly show euhedral habitus (Fig. 4, *b, e, f*). Most granites are more or less cataclased and show some schistosity with preferred phenocrysts orientation. Mafic minerals are completely or almost

completely chloritized. Most intense secondary alterations are noted in the near-contact zone where rocks are pyritized and rusty.

Granites are mainly composed of three minerals: alkali feldspar, plagioclase and quartz. Alkali feldspar phenocrysts vary in number from one site to another and generally amount to 20–30 % of the rock (on some sites, less than 10 % or more than 50 %). The least homogeneous rock composition is observed in the near-contact zone. In thin sections alkali feldspar is represented by perthite. Pelitization is insignificant and patchy, and it generally looks much fresher than plagioclase.

Quartz forms xenomorphic grains (up to 1 cm), which in places are lense-shaped (Fig. 4, *d*). In thin-sections quartz is granulated and shows complex undulose extinction. Cataclasis is manifested by bands of microgranular quartz.

Plagioclase is the main rock component in terms of volume. It is usually poorly distinguishable in field, although its crystals are quite distinct on some sites (Fig. 4, *c*). Plagioclase is sosuritized and looks greenish in hand specimens (Fig. 4, *e, f*). In thin sections it shows blurred polysynthetic twinning with dispersed secondary minerals.

Mafic minerals are represented by biotite and, rarely, amphibole. They are partially or more often completely replaced with chlorite. In thin sections biotite flakes are bent and contain chlorite lamellae, and flecked with pleochroic halos around zircon crystals and possibly some other mineral phase. Some biotite aggregations are heterogeneous in structure and colour, contain numerous inclusions of Fe-Ti minerals, and thus themselves look like a secondary minerals. Probably, they substitute another mafic mineral, and an excessive Fe-Ti forms new mineral phase.

Among accessory minerals (separated from sample 045/1–16) dominate zircon and apatite presented in roughly equal amounts. An unusual feature is the presence of galenite, which also mentioned by [Stepanov *et al.*, 1965]. Therefore, this is not a random mineralization, but a characteristic feature of the Pregradnaya massif.

4. PORPHYRITIC GRANITES IN THE UPPER REACHES OF THE STANOVAYA RIVER

The Neoproterozoic rocks predominate in the upper reaches of the Stanovaya River, and they are separated from the Mesoproterozoic rocks by Magistral'ny fault (Fig. 5). Non-metamorphosed uppermost Vendian and lower Paleozoic rocks are exposed in the fault graben and around it. A significant part of this territory is composed of the Neoproterozoic felsic metavolcanic rocks, according to [Stepanov *et al.*, 1965]. and on some

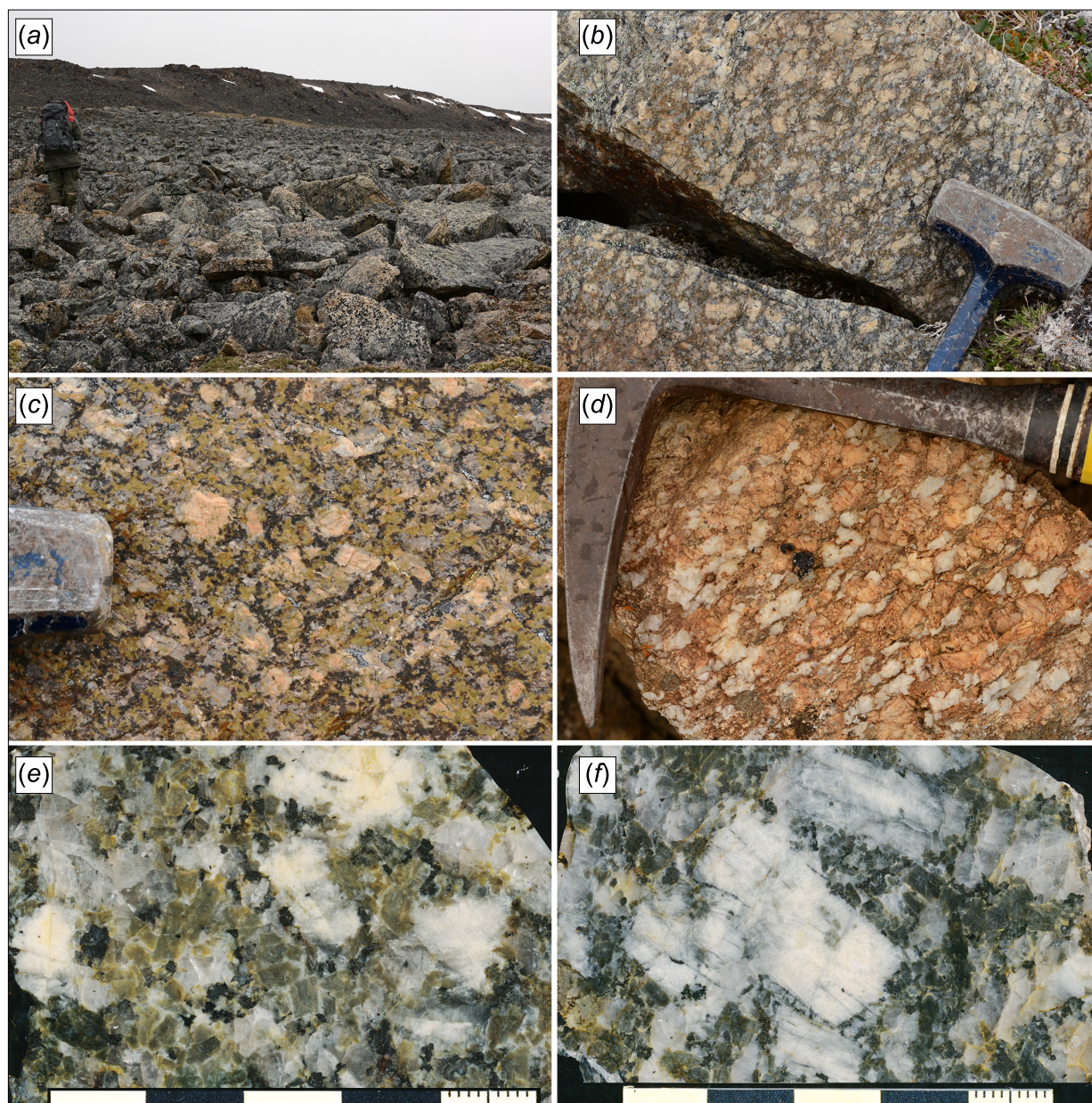


Fig. 4. Photographs of typical granites of the Pregradnaya massif.

(a) – granite talus near site 047 (see Fig. 3). The ridge in the background is composed of amphibolites and the Zhdanova Fm. metasedimentary rocks. (b) – typical appearance of porphyritic granite. Phenocrysts are made of K-feldspar. (c) – naturally cut rock surface showing all the major rock-forming minerals of the Pregradnaya granites: large yellowish-pink mineral – perthitic feldspar; smaller, dirty-green – epidotized plagioclase; grey fuzzy shaped – quartz; black – chloritized biotite. (d) – xenomorphic quartz (white) juts over weathered granite surface. The predominant mineral in this leucocratic variety is yellowish-pink alkali feldspar. (e), (f) – scanned saw-cut granite samples. Below is a ruler in centimeters. (e) – sample 046/3-16. Idiomorphic tabular crystals of the saussuritized plagioclase (greenish) are visible. Light-colored mineral – xenomorphic orthoclase with inclusions of quartz, plagioclase and biotite along the periphery. Black mineral – chloritized biotite. (f) – sample 044/1-16. Similar rock, but orthoclase is more idiomorphic, with quartz inclusions.

Рис. 4. Фотографии типичных гранитов Преградненского массива.

(a) – гранитный курум вблизи точки 047 (см. рис. 3). Хребтик на заднем плане сложен метаосадочными породами ждановской толщи и амфиболитами. (b) – типичный порфировидный облик пород. Вкрапленники сложены калиевым полевым шпатом. (c) – естественный скол глыбы, на котором видны все главные породообразующие минералы гранитов: крупный желторозовый минерал – ортоклаз-пертит, более мелкий грязно-зеленый – эпидотизированный плагиоклаз, серый нечетко оформленный – кварц, черный – хлоритизированный биотит. Слева видна часть кувалды (той же, что и на предыдущей фотографии). (d) – выветрелая поверхность гранита с отпрепарированными ксеноморфными выделениями кварца (белое). Преобладающий минерал в этой лейкократовой разновидности – желтовато-розовый ортоклаз-пертит. (e), (f) – сканированные спилы образцов гранита. Внизу – линейка в сантиметрах. (e) – образец 046/3-16. Видны идиоморфные таблички сосюритизированного плагиоклаза (зеленоватые). Светлое – ксеноморфный ортоклаз с включениями кварца, плагиоклаза и биотита по периферии. Черное – агрегат хлоритизированного биотита. (f) – образец 044/1-16. Сходная порода, но ортоклаз имеет четкие очертания. В нем видны включения кварца.

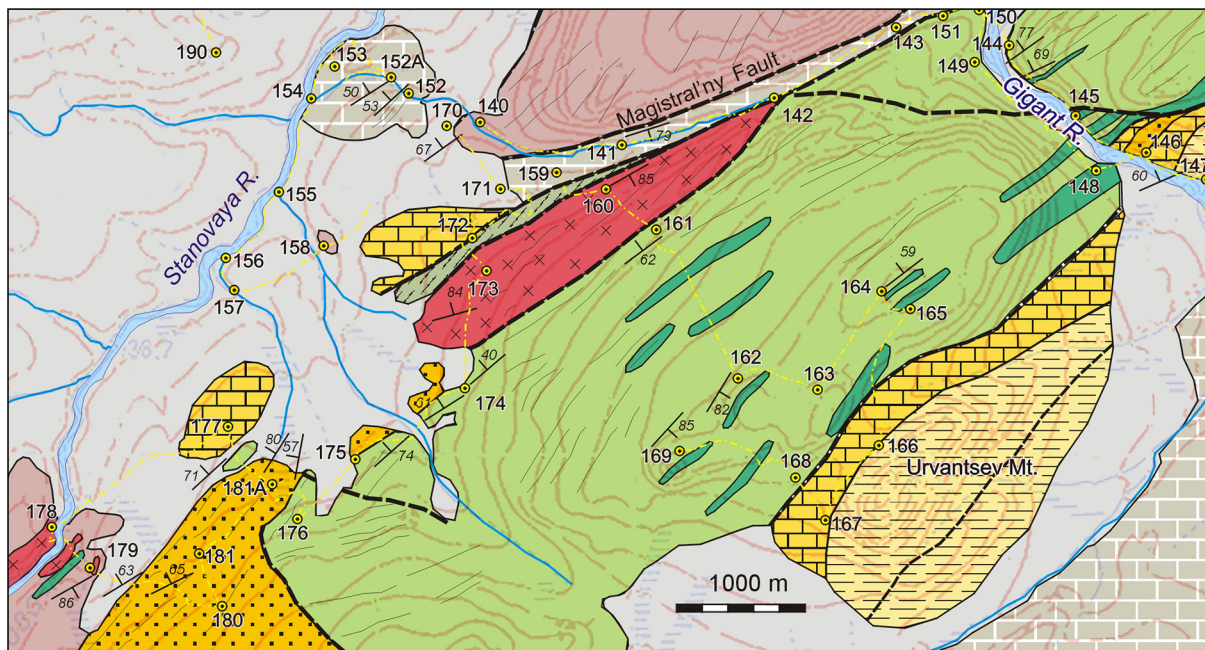


Fig. 5. Geological map illustrating the structural positions of granitic bodies in the upper reaches of the Stanovaya River. Red – granites (~840 Ma); deep green – Neoproterozoic amphibolites. See Fig. 3 for other symbols.

Рис. 5. Геологическая карта, иллюстрирующая структурное положение гранитных тел в верховьях р. Становой. Красным цветом показаны граниты (~840 млн лет), густо-зеленым цветом – неопротерозойские амфиболиты. Остальные условные обозначения см. на рисунке 3.

sites they resemble sheared porphyry. However the bulk of rocks can be rather described as mylonites or blastomylonites. Their volcanic protolith is quite probable: obvious meta-sedimentary rocks are absent in the succession; the chemical composition of the prevailing varieties is homogeneous; the zircon crystals extracted from one of samples show igneous origin and belong to a single age population (840 ± 4 Ma, unpublished data of the authors). There are numerous amphibolite bodies within felsic metavolcanics, which may be interpreted as metabasalts.

In the southeast and southwest, the volcanic rocks are in contact with the mid-Neoproterozoic rocks of the Stanovaya-Kolosova Group. On the slopes facing the Stanovaya River, the lower horizons of the Group were mapped earlier as the Urvantsev Fm. by [Stepanov *et al.*, 1965]. These are green schists with small (1–2 mm) red microcline porphyroblasts and quartz gravel. On some sites, these rocks are replaced by arkose or quartz gritstones. The higher horizons of the Group exposed on the opposite side of the area shown in Fig. 5 comprise conglomerates, limestones, dolomites (including stromatolitic ones) and quartzites. Grey schists composing the upper horizons of the group are observed on the Urvantsev Mountain. Two granitic bodies named as Stepanov (lower left corner in Fig. 5) and Magistral'ny (in the middle of Fig 5) massifs are discussed below.

4.1. STEPANOV MASSIF

This small granite body was selected for detailed study as it was presumable intrude the Urvantsev Fm. [Stepanov *et al.*, 1965], and so its geological position seemed similar to that of Pregradnaya massif. The Stepanov massif is well exposed on the left bank of the Stanovaya River (Fig. 6, *a*). There a deformed porphyritic granite with gneissic structure crops out (Fig. 6, *b*). Some varieties with rounded alkali feldspar resemble an augen gneiss. Feldspar phenocrysts may be a dominating mineral phase (Fig. 6, *c*).

The exposures continue on the right bank of the Stanovaya River (Fig. 6, *d*), where cataclased gneissic granites also crop out. There they stronger altered and associate with amphibolites and green schists. Further on (# 179 in Fig. 5; a scarp visible in Fig. 6, *d*, in the midst of the left half of the picture), there is a rocky exposure of dolomite marble cut with numerous sills and dikes of porphyritic granite (Fig. 6, *e*). Some of them are strongly budinaged (Fig. 6, *f*) due to plasticity of the host carbonate media. Granites are cataclased, carbonatized and pyritized.

4.2. MAGISTRAL'NY MASSIF

The massif crops out in lower scarps of terraced slope (# 160 in Fig. 5 and Fig. 7, *a*). In most exposures

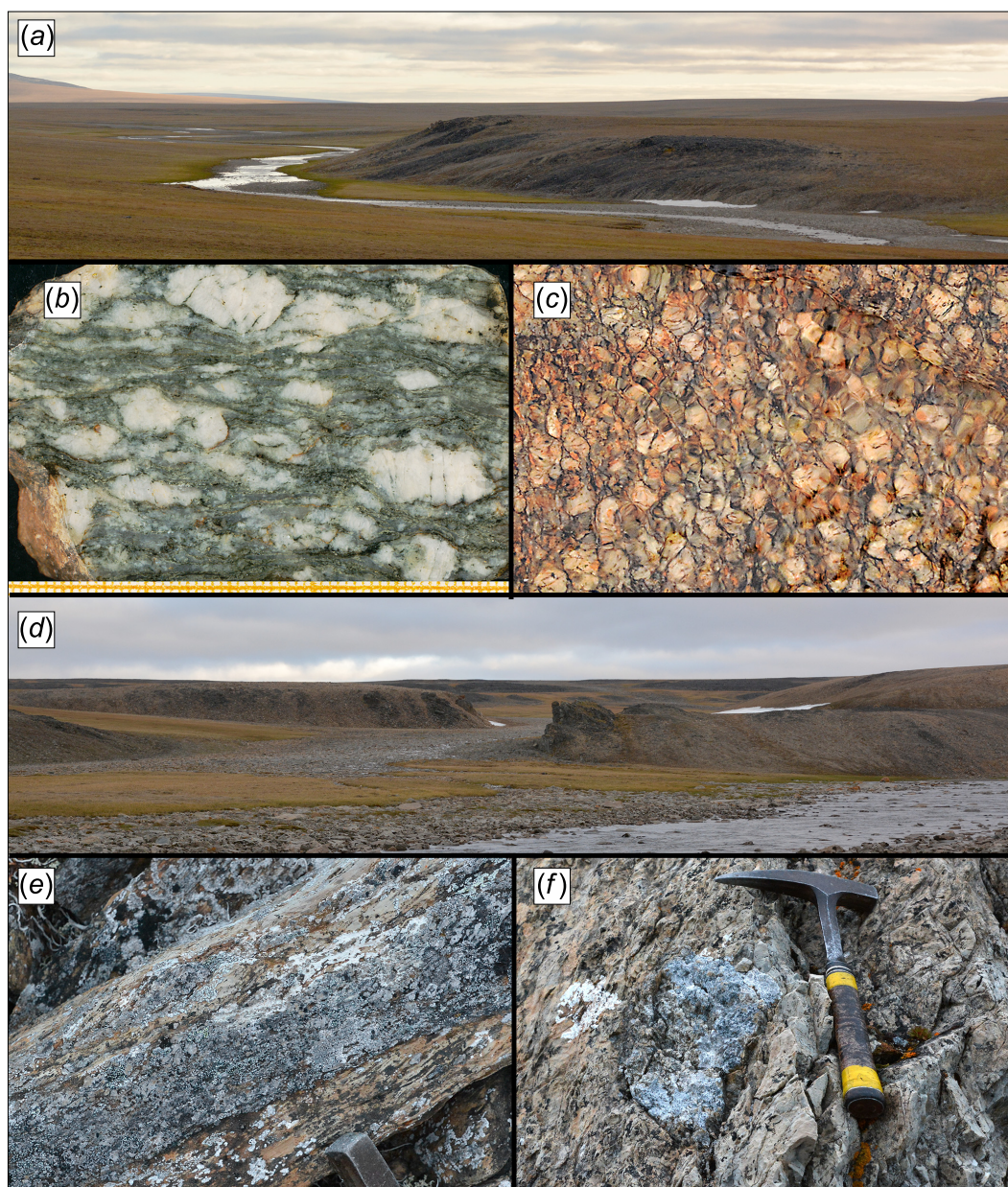


Fig. 6. Stepanov granite massif.

(a) – general view of the gneissic granite outcrop on the left bank of the Stanovaya River (site 178, see Fig. 5). (b) – scanned image of the cut surface of a typical granite sample from the outcrop (sample 178/1). The rock is cataclased and schistose. Only the alkali feldspar shows distinct contours; quartz composes vague veins and lenses, conformable with the general schistosity; plagioclase and biotite are transformed into foliated epidote-chlorite-sericite aggregate. Millimeter paper at the bottom of the photo for scale. (c) – porphyritic granite block below the water surface. It is almost entirely composed of potassium feldspar phenocrysts. (d) – view from site 178 on the right bank of the Stanovaya River. Ridge in the foreground (right half of the picture) composed of altered gneissic granites, amphibolites and schists. The next ridge in the left half is made of marble penetrated by granite veins (site 179). Green schists and gritstones of the Urvantsev Fm. crop out in background (see Fig. 5). (e) – porphyritic granite vein (bluish-grey) in marble of the Zhdanov formation. Dated sample 179/2-16 was taken from a similar vein. (f) – granite boudin in massive dolomite marble.

Рис. 6. Фотографии гранитов Степановского массива.

(a) – общий вид обнажения гнейсогранитов на левом берегу р. Становой (т.н. 178, см. рис. 5). (b) – скан поверхности распила типичного гранита этого обнажения (обр. 178/1). Порода катаклазирована и рассланцована. Оформленные контуры сохранились только у кристаллов щелочного полевого шпата, кварц слагает жилки и линзочки, согласные с общим рассланцеванием, плагиоклаз и биотит превращены в рассланцованный эпидот-хлорит-серичитовый агрегат. Миллиметровка в нижней части фото для масштаба. (c) – глыба порфировидного гранита, почти полностью состоящая из фенокристов КПШ. (d) – вид с точки 178 на обнажения правого берега р. Становой. Скалки на ближнем плане справа сложены сильно измененными гранитогнейсами, амфиболитами и сланцами. Скальные выходы в левой половине фотографии – обнажения мраморов (т.н. 179) с жилами гнейсогранитов. На дальнем плане – гряды зеленых сланцев и гравелитов «урванцевской свиты» (см. рис. 5). (e) – жила порфировидных гранитов (серое) среди рассланцованных мраморов ждановской толщи. Из подобной жилы отобран продатированный образец 179/2-16. (f) – гранитная будина в массивных доломитовых мраморах.

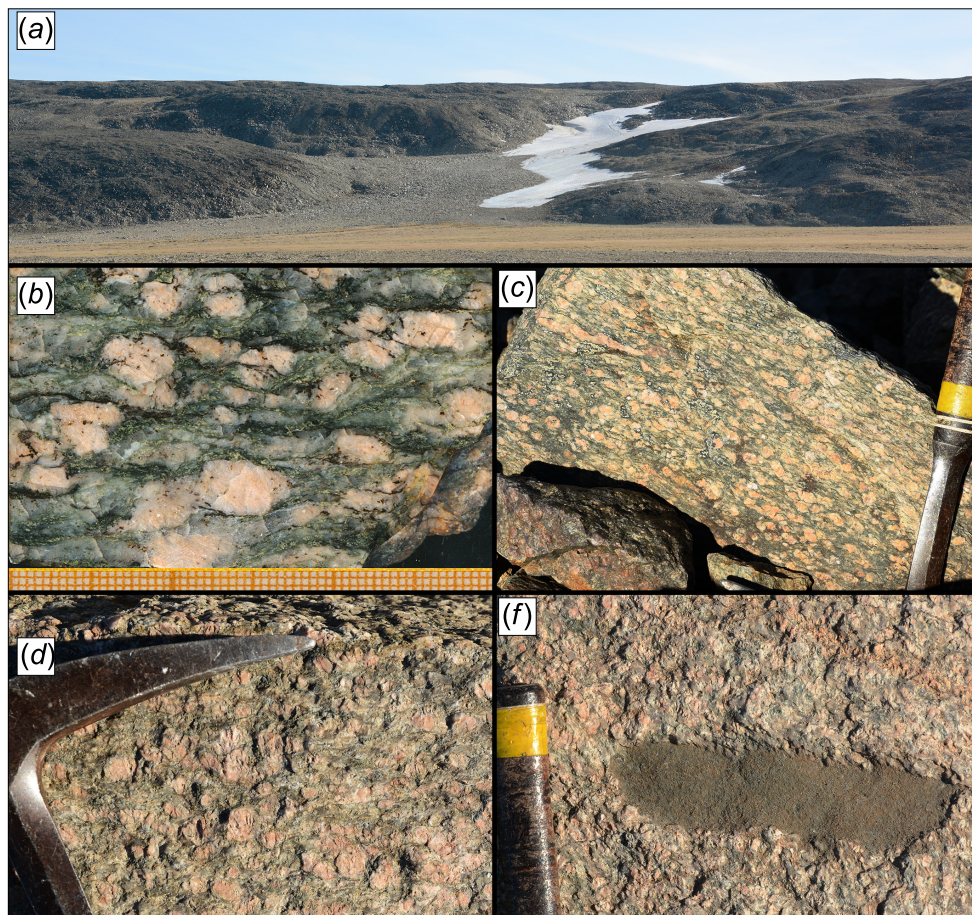


Fig. 7. The Magistral'ny granite.

(a) – scarps and cryoplanation terraces composed of gneissic granites. At the foot of the slope on the left is site 160 (see Fig. 5). (b) – texture of gneissic granite from the lower ledge (sample 160/2, see text). Millimeter paper is at the bottom. (c) – gneissic granite in the lower scarp. The matrix was mylonitized, and the original granite texture is completely lost. (d) – augen gneissic granite. The rock is less foliated and contains plagioclase remnants. (e) – xenolith in granite. The original granitic texture is partly preserved.

Рис. 7. Граниты Магистрального массива.

(a) – вид на уступы (нагорные террасы), сложенные разгнейсованными гранитами. У подножия склона слева – точка 160 (см. рис. 5). (b) – структура гнейсогранитов нижнего уступа на распиле (обр. 160/2) (описана в тексте). Внизу – миллиметровка для масштаба. (c) – сильно расланцованный гнейсогранит первого уступа с милонитизированным матриксом, полностью утратившим гранитную структуру. (d) – очковый гнейсогранит в первом уступе. Порода менее расланцована, и в ней угадываются реликты кристаллов плагиоклаза. (e) – ксенолит в граните. Видна сравнительно слабо измененная гранитная структура субстрата.

granites look like metasomatic rocks with K-feldspar porphyroblasts. Such metasomatism is actually observed at the NW margin of the massif at the foot of slope, where sericite-chlorite green schists contain scattered large (1–3 cm) feldspar porphyroblasts.

The granitic body proper is composed of augen gneissic granite. The rock contains rounded and shapeless, rarely idiomorphic crystals (to 1.5 cm) of pink perthitic feldspar in the schistose matrix. The latter composed of quartz lenses with inclusions of albite, and scabby zoisite-sericite aggregate. This aggregate replaces plagioclase, as evidenced by blurred polysynthetic twinning in some samples. Dark bands and lenses are composed of chlorite and biotite.

In the vicinity of the 160 point (see Fig. 5, Fig 7, a) structure of rocks is clearly visible in the fresh-cut blocks and walls. In the lower part of the slope, the rocks often demonstrate cataclasis and dynamic recrystallization of quartz and K-feldspar to form elongate and lenticular segregations (Fig. 7, b, c) immersed into sericite and chlorite-epidote-sericite aggregate. Higher on the slope occur less sheared rocks with rounded phenocrysts (Fig. 7, d). In the same part of the massif even less altered varieties can be found, which preserve remnants of granite structure, and contours of tabular crystals of plagioclase are clearly visible. Presence of xenoliths in this part of the massif (Fig. 7, e) give unequivocal evidence that this is an igneous rock crystallized from magma. We as-

sume that the augen gneisses of the lower scarp also crystallized as granites, but were sheared afterwards.

5. CHEMICAL COMPOSITION

All the analyzed rocks (Table 1) are the most common potassium granites with concentrations (dry composition, wt. %) of: silica 71–75, alumina 13–16, alkalis 7.2–8.7, and $K_2O/Na_2O=1.0–1.7$ (in one case – 0.4). All granites are weakly peraluminous. The alumina saturation index is 1.05–1.20, except one analysis (0.99). In the geochemical classification scheme of granitic rocks [Frost et al., 2001; Frost B.R., Frost C.D., 2008], all the analyzed rocks are magnesian and can be further classified into calc-alkalic and alkali-calcic, with the exception of one analysis that is calcic.

The gross chemical compositions vary only slightly, and normalized concentrations of minor elements look homogeneous in the diagrams. The pattern of rare earth elements (REE) contents is the same for all samples and is generally typical for the most common varieties of granites: weakly fractionated heavy REE (around 10 chondritic values), stronger fractionated light REE (to 100 chondritic values), and insignificant negative europium anomaly (Fig. 8). Comparison with the Snezhnaya granites (shaded area in Fig. 8) shows that in the latter is somewhat higher REE content. The multi-element diagram shows the Nb and Ti negative anomalies inherited from the source and typical of crustal magmas. The Sr negative anomaly is generally attributed to plagioclase fractionation. Obviously, magma sources for all the analyzed granites were crustal rocks and, possibly, middle-crust rocks in the garnet stability zone, as evidenced by the relatively low concentrations of heavy REEs. According to the above data, the geochemical characteristics of the Pregradnaya, Stepanov and Magistral'ny massifs are quite similar and do not differ from those of the Snezhnaya complex described in [Proskurnin et al., 2014, and references therein]. However, the zircon dating shows that the Pregradnaya granite and the granites from the upper reaches of the Stanovaya River belong to different complexes.

6. U-Pb ISOTOPE ANALYSIS OF ZIRCONS

Zircons were extracted from the granite samples by first two authors in the laboratory of GIN RAS. Isotopic analyses of U, Th and Pb was carried out on a SHRIMP-II mass spectrometer at VSEGEI, using a procedure similar to that described in [Williams, 1998]. Ion beam diameter was set to $\sim 25 \mu\text{m}$; each spot scanned four times; ion countings was taken successively through range of masses in each scan. Temora-2 standard was used for calibration of U/Pb isotopic ratios; standard

91500 to determine U and Pb concentrations; data were processed using SQUID and ISOPLOT software [Ludwig, 2001, 2003].

6.1. SAMPLE 045/1-16, PREGRADNAYA MASSIF

Zircon crystals are dark brown-coloured, euhedral without signs of dissolution or overgrowth; they are heavily fractured and contain numerous inclusions (Fig. 9). Large crystals are short-prismatic, among small ones, occur also long-prismatic to needle-shaped crystals. They have high U concentrations (to 1200 ppm) and look very dark in cathodoluminescent (CL) images. The magmatic oscillatory zoning is clearly visible in backscattered electron (BSE) images and even in optical images (see Fig. 9). Cores are usually recrystallized, often to a complete loss of the original structure. We analyzed 11 spots on 11 crystals. Zircons are characterized by a 'magmatic' Th/U ratio (Table 2). Almost all analyses contain substantial (some very high) proportion of common Pb (Table 2). This is partly due to specific properties of the Pb-rich granite magma, as evidenced by the presence of galenite among heavy minerals. High common Pb concentrations in some analyses makes it impossible to adequately calculate a correction for common Pb by the standard model. In the conventional isotopic diagram ($^{207}\text{Pb}/^{235}\text{U}$ vers. $^{206}\text{Pb}/^{238}\text{U}$), six analyses became reversely discordant if apply a correction for common Pb. The age of the concordant cluster calculated for the remaining five analyses is 607 ± 2 Ma (2s), and MSWD (of concordance) = 0.46. The Tera-Wasserburg isotope diagram operating with total Pb (including common Pb) is more useful for such data. It turned out that all 11 analyses can be approximated by a straight line with very good statistical parameters (Fig. 10). The lower interception with the concordia at a point of 609 ± 2 (2 sigma) million years indicates the time of zircon crystallization.

6.2. SAMPLE 160/2-16, MAGISTRAL'NY MASSIF

Zircons of this sample are quite different from Pregradnaya ones; they are cleaner, without cracks, the prism is not a square in section but closer to an octagon. The CL images show normal oscillatory zoning (Fig. 9). In many crystals, the inner part is more uniform and relatively light (i.e. lower U and REE concentrations), while the outer part is darker, with fine contrasting oscillatory zoning. Some crystals retained more or less explicit cores. The analyzed spots show moderate concentration of uranium (up to 500 ppm) and a 'magmatic' Th/U ratio (Table 2).

Ten spots on nine crystals were set mostly on the external oscillatory zone, which usually occupies most of a crystal. The central part was also analyzed for one crystal and yielded a similar age. All the 10 analyses

Table 1. Chemical compositions of granites (major oxides in wt. %; trace elements in ppm)

Таблица 1. Химический состав гранитов. Породообразующие окислы в мас. %, малые элементы в ppm

Components	Pregradnaya massif				Magisral'ny massif		Stepanov massif
	044/2-16	045/1-16	046/3-16	046/4-16	160/2-16	160/3-16	178/1-16
SiO ₂	71.4	70.8	69.1	72.5	70	71.2	72.6
Al ₂ O ₃	14.1	13.5	14.8	12.8	15.5	14.9	14.7
TiO ₂	0.28	0.26	0.29	0.24	0.3	0.26	0.21
Fe ₂ O _{3t}	2.71	3.79	3.04	2.55	2.79	2.4	2.15
MnO	0.046	0.062	0.064	0.047	0.06	0.052	0.039
MgO	0.74	0.83	0.82	0.7	0.75	0.66	0.63
CaO	1.15	1.64	1.49	2.04	1.34	1.1	0.76
Na ₂ O	2.83	3.1	3.4	3.23	4.18	4.29	5.55
K ₂ O	4.84	4.62	5.15	3.79	4.27	4.21	2.07
P ₂ O ₅	0.08	0.076	0.07	0.064	0.061	0.06	0.056
LOI	1.74	1.24	1.85	2.23	0.85	0.93	1.22
SUM	99.9	99.9	100	100	100	100	100
ASI ^a	1.20	1.05	1.08	0.99	1.12	1.11	1.16
Alkali/Lime ^b	calc-alk	calc-alk	alk-calc	calcic	alk-calc	alk-calc	calc-alk
V	17.6	23.9	16.4	14.1	9.35	8.16	13.7
Cr	13.4	9.96	9.95	11.8	16	12.6	11.1
Rb	162	156	151	116	103	110	69.6
Sr	142	382	164	175	159	149	75.3
Y	13.8	19.7	18.3	13.8	19.9	27.8	16.9
Zr	149	196	127	134	184	170	116
Nb	8.57	10.1	8.87	8.33	9.57	11.1	11.2
Mo	2.19	9.14	3.27	1.98	1.93	1.5	1.46
Cs	2.6	1.81	2.09	1.59	1.09	1.41	0.65
Ba	898	1090	835	685	692	599	548
La	52.3	57.1	47.6	31.5	39.6	39.5	19.6
Ce	91.4	97.8	77.6	53.3	98.2	76.6	42.7
Pr	9.33	9.99	7.62	5.32	9.54	9.53	4.75
Nd	30.4	31.8	28	19	37.8	38.2	19
Sm	4.34	4.8	5.09	3.28	6.93	6.68	3.75
Eu	0.79	0.76	0.62	0.62	1.22	1.42	0.65
Gd	3.67	4.18	3.71	2.53	5.16	6.1	3.33
Tb	0.5	0.59	0.54	0.42	0.73	0.92	0.54
Dy	2.32	3.32	2.75	2.14	3.74	4.79	2.96
Ho	0.44	0.66	0.58	0.4	0.69	0.9	0.57
Er	1.28	1.84	1.51	1.18	2.05	2.56	1.7
Tm	0.21	0.31	0.26	0.21	0.32	0.4	0.27
Yb	1.38	2.03	1.61	1.38	1.94	2.56	1.82
Lu	0.21	0.29	0.28	0.24	0.32	0.4	0.29
Hf	4.63	5.35	3.74	4.09	5.46	4.78	3.66
Ta	0.86	0.73	0.97	1.14	0.66	0.97	1.17
Th	25	29.1	18.2	18.2	10.6	7.42	15.1
U	5.11	4.27	2.69	1.83	1.41	1.42	1.65
Pb	51.6	59.1	47.1	32.7	15.5	14.8	8.35
Li	18.1	15.5	21.6	23.5	11.2	10.1	16.5

Note. Analyzed at VSEGEI: major oxides by XRF, trace elements by ICP MS, Li and Pb by ICP-AES. ^a – Alumina Saturation Index = molar Al₂O₃/((CaO-3.33P₂O₅)+Na₂O+K₂O). ^b – after [Frost et al., 2001; Frost B.R., Frost C.D., 2008].

Примечание. Анализы выполнены в ВСЕГЕИ: основные окислы по XRF, микроэлементы по ICP MS, Li и Pb по ICP-AES. ^a – индекс насыщения оксида алюминия = молярный Al₂O₃/((CaO-3.33P₂O₅)+Na₂O+K₂O). ^b – согласно [Frost et al., 2001; Frost B.R., Frost C.D., 2008].

form a concordant cluster with age of 843±6 Ma (Fig. 10).

6.3. SAMPLE 179/2-16, STEPANOV MASSIF

Zircons from this sample are very similar in every respect to those from the Magisral'ny massif (see Fig. 9),

some crystals exhibit indistinct cores. In the CL images the central part usually is lighter with rough zoning, while the outer part is darker, with finer oscillatory zoning. In some zircons, there is one more relatively bright zone between the above two. The U concentrations and Th/U ratios are approximately similar to those of sample 160/2-16 (Table 2). Central and marginal

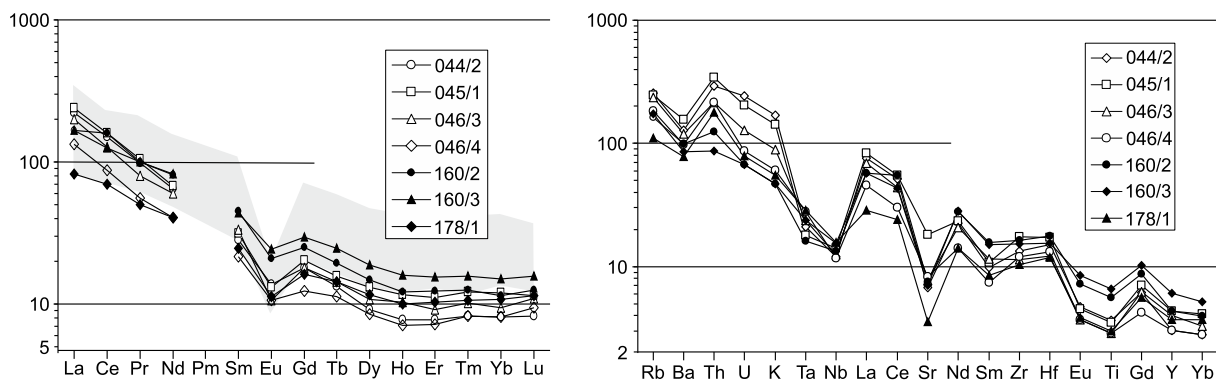


Fig. 8. Trace element diagrams: chondrite-normalized REE concentrations on the left; ‘incompatible’ elements normalized to primitive mantle on the right. Compositions of chondrite and primitive mantle after [Sun, McDonough, 1989]. Shaded contour on the REE graph – Snezhnaya granitic complex at the upper reaches of the Leningradskaya River [Proskurnin et al., 2014].

Рис. 8. Графики распределения малых элементов. Слева – концентрации РЗЭ, нормированные по хондриту, справа – мультиэлементный график концентраций «несовместимых элементов» нормированных по примитивной мантии. Состав хондрита и примитивной мантии по [Sun, McDonough, 1989]. Затемненный контур на графике РЗЭ - гранитоиды снежинского комплекса верховьев р. Ленинградской по [Proskurnin et al., 2014].

parts of the crystals were tested, and all the three zones were analyzed for one crystal. In total, nine analyses were performed for five crystals, and all the CL zones have returned the same age. Eight analyses form a clus-

ter with age of 840 ± 5 Ma (Fig. 10), which coincides with the age of Magistralny massif within the errors. One analysis on the rim gave ~ 810 Ma probably due to Pb loss during late Neoproterozoic thermal impact.

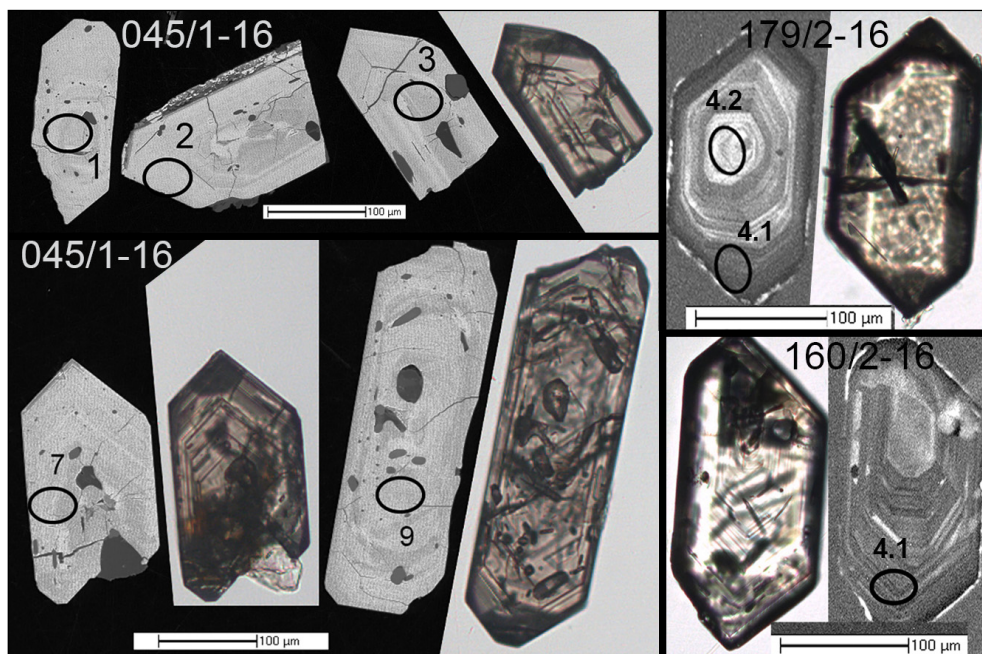


Fig. 9. Examples of the analyzed zircon crystals imaged in transmitted light, backscattering electrons (sample 045/1-16) and cathodoluminescence (samples 179/2-16 and 160/2-16). Left – Pregradnaya granite (045/1-16; upper right – Stepanov granite (179/2-16); lower right – Magistral’ny granite (160/2-16). The spot numbers are labeled.

Рис. 9. Изображения проанализированных кристаллов циркона, выполненные в проходящем свете, в режиме обратно-рассеянных электронов (для образца 045/1-16) и катодолуминесценции. Левая панель – Преградненский массив (045/1-16), правая верхняя панель – Степановский массив (179/2-16), правая нижняя панель – Магистраль-ный массив (160/2-16). Подписаны номера проанализированных точек, соответствующие таблице 2.

Table 2. Zircon U-Pb isotopic data for discussing granites (SHRIMP II, VSEGEI)
Таблица 2. Результаты изотопного анализа цирконов (ВСЕГЕИ, SHRIMP II)

Spot #	% ^{206Pb_c}	ppm U	ppm Th	ppm ^{232Th/238U}	(1) ^{206Pb/238U} Age	(2) ^{206Pb/238U} Age	(3) ^{238U/206Pb}	±%	(1) ^{207Pb/235U}	±%	(1) ^{206Pb*/238U}	±%	err corr		
Sample 045/1-16 Pogradnaya massif															
1.1	16.78	901	587	91.2	577±3	602±3	8.494	0.29	0.19710	1.20	0.4278	15.7	0.093683	0.61	0.04
2.1	4.14	1256	549	111.0	605±2	606±2	9.722	0.26	0.09391	0.65	0.8131	2.5	0.098362	0.31	0.13
3.1	18.39	548	369	56.4	588±5	601±4	8.343	0.39	0.21020	0.68	0.6715	13.7	0.095545	0.94	0.07
4.1	10.64	259	194	24.4	593±5	603±4	9.115	0.54	0.14700	0.98	0.6845	11.8	0.096363	0.92	0.08
5.1	0.44	1287	465	110.0	609±2	609±2	10.057	0.26	0.06370	0.60	0.8297	0.9	0.099058	0.27	0.30
6.1	1.68	449	447	38.9	607±3	609±3	9.919	0.39	0.07390	2.00	0.7845	4.5	0.098706	0.47	0.11
7.1	3.02	1059	527	92.8	606±2	608±2	9.798	0.28	0.08490	1.30	0.7904	3.6	0.098563	0.36	0.10
8.1	2.30	471	221	41.1	603±3	610±3	9.845	0.44	0.07901	0.89	0.6922	4.3	0.097983	0.51	0.12
9.1	0.12	714	366	60.9	608±2	610±2	10.067	0.37	0.06117	0.82	0.7864	1.6	0.098887	0.38	0.24
10.1	9.55	512	490	47.5	581±4	601±4	9.261	0.48	0.13790	2.90	0.4791	17.2	0.094325	0.73	0.04
11.1	6.58	841	497	75.7	593±9	602±2	9.541	0.34	0.11371	0.64	0.6764	22.9	0.096422	1.52	0.07
Sample 160/2 Magistral'ny massif															
1.1	0.02	348	133	41.1	831±6	797±24	7.265	0.78	0.06569	1.2	1.247	1.4	0.13765	0.78	0.56
2.1	0.16	440	353	52.1	832±6	869±19	7.258	0.77	0.06802	0.9	1.292	1.2	0.13779	0.77	0.63
3.1	-	518	300	62.6	848±9	839±17	7.109	1.11	0.06704	0.8	1.300	1.4	0.14067	1.11	0.80
4.1	0.03	505	379	60.2	838±14	819±20	7.201	1.80	0.06638	0.9	1.271	2.0	0.13888	1.80	0.89
5.1	-	246	177	29.7	847±7	894±29	7.124	0.86	0.06886	1.4	1.333	1.7	0.14038	0.86	0.52
5.2	-	297	158	36.0	851±7	776±26	7.085	0.82	0.06506	1.2	1.266	1.5	0.14114	0.82	0.55
6.1	-	257	138	30.9	846±8	817±25	7.134	0.97	0.06634	1.2	1.282	1.5	0.14018	0.97	0.63
7.1	-	558	330	67.8	852±6	825±17	7.075	0.73	0.06661	0.8	1.298	1.1	0.14135	0.73	0.68
8.1	-	378	219	46.1	856±12	835±21	7.046	1.46	0.06691	1.0	1.309	1.8	0.14192	1.46	0.82
9.1	0.04	420	274	50.7	847±6	833±21	7.122	0.77	0.06684	1.0	1.294	1.3	0.14041	0.77	0.60
Sample 179/2 Stepanov massif															
3.1	0.25	681	266	78.0	807±6	819±18	7.495	0.77	0.06639	0.9	1.221	1.2	0.13342	0.77	0.66
1.1	0.06	164	87	19.7	843±7	857±29	7.154	0.94	0.06763	1.4	1.303	1.7	0.13978	0.94	0.56
1.2	0.06	489	266	59.0	848±6	846±19	7.115	0.75	0.06727	0.9	1.304	1.2	0.14055	0.75	0.63
2.1	-	300	152	36.0	845±6	748±29	7.144	0.81	0.06419	1.4	1.239	1.6	0.13998	0.81	0.51
2.2	0.04	147	62	17.4	834±8	798±40	7.241	1.00	0.06575	1.9	1.252	2.2	0.13811	1.00	0.46
2.3	0.16	937	388	110.0	829±9	840±15	7.286	1.10	0.06708	0.7	1.269	1.3	0.13724	1.10	0.83
4.1	0.03	405	189	48.0	833±7	802±23	7.254	0.85	0.06588	1.1	1.252	1.4	0.13786	0.85	0.61
4.2	0.07	172	104	21.1	859±13	833±35	7.017	1.67	0.06685	1.7	1.314	2.4	0.14250	1.67	0.70
5.1	0.01	438	358	52.5	842±6	833±21	7.165	0.75	0.06683	1.0	1.286	1.2	0.13957	0.75	0.60

Note. Errors are 1-sigma. Pb_c and Pb* indicate the common and radiogenic portions, respectively. (1) Common Pb corrected using measured ^{204Pb}. (2) Common Pb corrected by ^{207Pb}. (3) Uncorrected for common Pb. Error in Temora-2 standard calibration was 0.27%, 0.27% and 0.32% for samples 160/2-16, 179/2-16 and 045/1-16 respectively.

Примечание. Ошибки приведены в интервале 1 сигма. Pb_c и Pb* обозначают обыкновенный и радиогенный свинец. (1) – введена коррекция на обыкновенный свинец на основании измеренного количества ^{204Pb}. (2) – поправка на обыкновенный свинец введена на основании ^{207Pb}. (3) – без введения поправки на обыкновенный свинец. Ошибки в определении возраста стандарта (Темора-2) составили 0.27%, 0.27% и 0.32% для образцов 160/2-16, 179/2-16 и 045/1-16 соответственно.

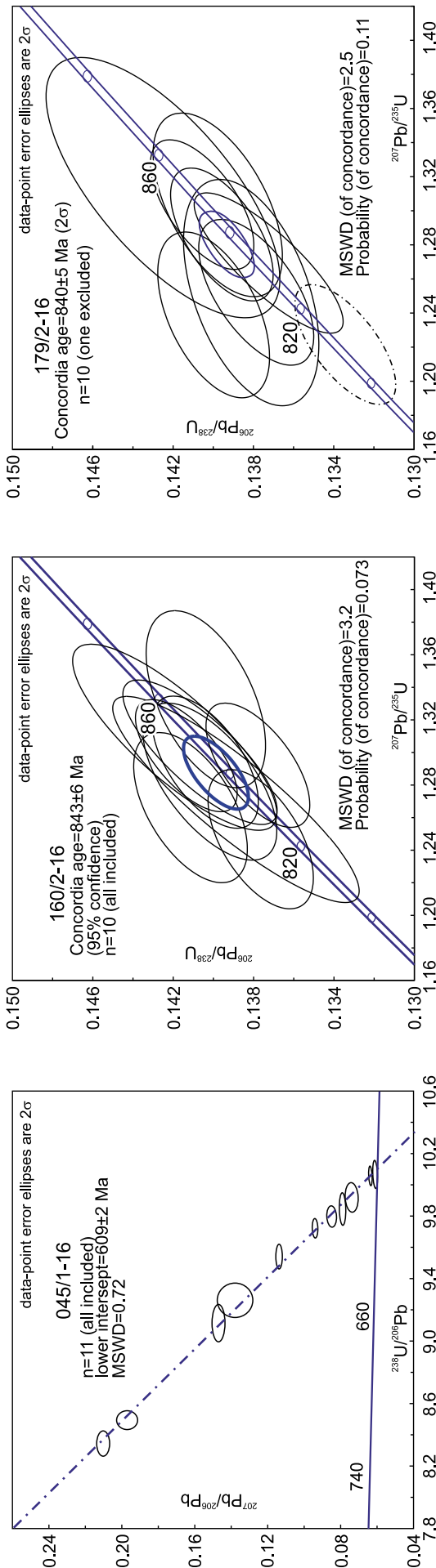


Fig. 10. U-Pb isotope diagrams for zircons of the Pregradnaya, Stepanov and Magistral'ny massifs.

Рис. 10. Изотопные диаграммы, иллюстрирующие результаты изотопного анализа цирконов из трех описанных в статье гранитных массивов.

7. DISCUSSION

7.1. CLASSIFICATION OF THE GRANITIC ROCKS AND MAGMA SOURCES

It seems reasonable in this section to indicate the type of granite our massifs belong to. Among the proposed classifications of granitoids (see the review in [Barbarin, 1990; Frost et al., 2001]), the most popular (at least in Russia) is the 'alphabetic' one, in which granitoids are tried to be related to one of the three most common types (I, S or A) according to recommendations in [Whalen et al., 1987; Chappel, White, 2001]. Granites can be readily discriminated in case of clearly manifested characteristic features, but the most common granites, including three massifs discussing in the paper, do not demonstrate such features. In our study, the following characteristics of the analyzed granites are symptomatic: (i) homogeneous composition with a narrow range of SiO₂ variations (70–75 wt. %); (ii) weakly-peraluminous composition (ASI=1.0–1.2) and potassium specialization of most varieties; (iii) absence of differentiates of dioritic and gabbroic compositions, at least, in obvious relation to granites. These characteristics are valid for described massifs and for other rather large plutons of the Snezhnaya complex, which we observed later in the Peka and Faddey river valleys (to the north of area shown in Fig. 2) as well. Such granites are close to the S-type (e.g., [Pearce et al., 1984; Chappell, White, 2001; Frost et al., 2001; Frost B.R., Frost C.D., 2008], and references therein), however, they differ from high-alumina muscovite granites typical for the S-type.

Anatexis of Mesoproterozoic meta-sedimentary rocks we observed at many points (see Fig. 2), and apparently these rocks were the main source of magma of the studied intrusions. Dating of leucosome of migmatites and related granitic veins (unpublished data of the authors) shows that among them there are rocks of the same ages as both discussed granite complexes.

7.2. GEODYNAMIC SETTING OF GRANITES

Plutons similar to the studied granite massifs can be emplaced into the upper crust in different geodynamic settings such as an active continental margin, a continent–continent or continent–island arc collision, post-collisional extension, and even within-plate ([Moyen et al., 2017], and references therein). It is impossible to distinguish those settings solely on the basis of chemical analyses of rocks, without additional geological information.

The following data should be taken into account. Dating of zircons from migmatite leucosome and leucogranite veins yields a range of ages from 980 to 800 Ma, with the most intense peak at 850–840 Ma ([Kuzmichev, Danukalova, 2018], and unpublished data of the authors). The protracted anatexis we associate with the

setting of an active continental margin, which is partially confirmed by the dating of volcanic rocks. The age range of the Snezhnaya granites is much narrower, 845–825 Ma. Obviously, melted materials from the depths did not always reach the upper crust, and extensive granite emplacement at the end of the first half of the Neoproterozoic was related to a discrete event. In our assumption, it was a collision event, which entail orogeny and subsequent deposition of the Stanovaya clastic rocks. In the course of the orogeny the Snezhnaya granites were exhumed and eroded, which is registered in the ages of detrital zircons [Priyatkina *et al.*, 2017]. Since the main characteristic features of Pregradnaya complex are similar to the Snezhnaya ones, the same conclusions are valid for both in terms of the geodynamic setting, although the scale of the last collision event was probably lower. The final stage of the Precambrian granite magmatism was also accompanied by orogeny. The synorogenic clastic rocks (Chekin Fm.) are known on the Chelyuskin Peninsula (see below).

7.3. OCCURENCE OF LATE NEOPROTEROZOIC GRANITIC ROCKS IN TAIMYR

The Pregradnaya massif is not the only evidence of the pre-Vendian crustal rocks fusion and granite emplacement. According to the authors' preliminary results, migmatites of close ages are found in the cape facing the north, located between the mouths of the Stanovaya and Gorodkova rivers (Fig. 2). The ages of almost half (48 spots) of zircon grains analyzed by LA ICPMS technique are in the range of 600–660 Ma with a peak value of 630 million years. Zircons with multi-stage overgrowth rims and old cores are challenging objects for laser ablation since material from several zones of different ages gets into the analysis. The data should be confirmed by the SIMS method with separate dating of the each zone. Perhaps here we met a root zone of the late Neoproterozoic granites.

Another granite pluton, located to the north of the Fig. 2 frame and studied later, may also be of the same age as the Pregradnaya massif. Among heavy minerals separated along with zircon (not dated yet) there was galenite, which is a typical feature of the Pregradnaya granite.

Further to the northeast, in the Zimovochnaya Bay area there is one more granite massif with the zircon SIMS age of ~630 Ma [Pease, Vernikovsky, 2000]. It is embedded into metamorphic rocks comparable to those in the study area and composed of deformed two-mica granite. The zircons contain early Neoproterozoic cores and show complex zoning.

Several late Neoproterozoic granite massifs (Nansen complex) are mapped approx. 30 km north of the Leningradskaya river mouth (see oval in Fig. 1). Three zircon samples from different massifs were analyzed using ID

TIMS technique (Paderin *et al.*, unpublished data). The results are not very reliable due to large errors that are not typical of the method. In total, eight analyses were done, and seven of eight were discordant due to Pb loss. They show different $^{207}\text{Pb}/^{206}\text{Pb}$ values and apparently do not belong to the same-age suite. One sub-concordant analysis yielded a 'concordant' age of 636.6 ± 1.1 Ma. According to [Makariev, 2013], these massifs belong to the Chukcha complex, and the SHRIMP dating (10 spots per sample) of two granite samples from the bank of Gafner Fjord (bottom part of the oval in Fig. 1) yields 640 ± 3 Ma and 635 ± 8 Ma. In general, this means that the pre-Vendian granites are more abundant in the north-eastern part of the Central Taimyr belt rather than in the territory covered by our study.

7.4. TAIMYR TECTONICS IN THE NEOPROTEROZOIC

Our study confirms two orogenic events in the Neoproterozoic evolution of the Central Taimyr belt (e.g., [Vernikovsky, 1996; Vernikovsky, Vernikovskaya, 2001]) and we assume that both events were related to collisions. The first one was manifested by emplacement of the Snezhnaya granites with the age of 845–825 Ma. The collision could have started earlier than 845 Ma and finished earlier than 825 Ma as the most of Snezhnaya plutons are discordant, i.e. late- or post-collisional ones. On the other hand, orogeny could lasted more long. The youngest zircons from migmatites and granite veins in studied area have shown the age of 780–770 Ma [Kuzmichev Danukalova, 2018]. This is the age of the youngest metamorphic rims, which usually do not constitute a distinct cluster in isotope diagram, but distributed along concordia line. The youngest ages correspond to final stages of a long period of repeated dissolution and overgrowth of zircons and these final stages may include Pb loss due to late Neoproterozoic thermal effect, when regime of active margin in the northern Central Taimyr belt was reestablished. Under such uncertainty we tentatively assign termination of early Neoproterozoic orogeny to 800 Ma, that is to the date more confidently proved by isotopic analyses.

V.A. Vernikovsky relates mid-Neoproterozoic orogeny to 'accretion – collision' processes that formed the Central Taimyr microcontinent [Vernikovsky, 1996; Vernikovsky, Vernikovskaya, 2001; Proskurnin *et al.*, 2014]. We, however believe that a substantial part of the Central Taimyr belt area belonged to the Siberian craton at least from Mesoproterozoic, and the 'accretion – collision' processes could have developed only along the northern margin of the belt [Priyatkina *et al.*, 2017; Kuzmichev, Danukalova, 2018].

In the second half of the Neoproterozoic, magmatism was predominantly active in the northern and northeastern parts of the modern Central Taimyr belt, wherein suprasubduction volcanic rocks of the late Ne-

oproterozoic age are widespread. Their isotopic ages that are summarized in [Makariev, 2013] are in the range of 760–610 Ma. According to published data [Markovsky et al., 2000; Makariev, 2013] these supra-subduction volcanic rocks unconformably lie on the Zhdanova and Kolosova formations. As shown above, the first formation is Mesoproterozoic in age, the second is mid-Neoproterozoic, and both were formed within bounds of the Siberian craton. This means that if the indicated correlations with the Zhdanova and Kolosova Formations are correct, then in the late Neoproterozoic, the NE part of the Central Taimyr belt was also a part of the Siberian paleocontinent. The presence of differentiated (basalt-andesite-rhyolite) suprasubduction volcanics indicates the active margin regime in the second half of the Neoproterozoic that was concluded by collision event in Vendian.

The synorogenic clastic deposits of this stage are preserved only in the NE part of the Central Taimyr belt. These are variegated clastic deposits of about 900 m-thick composed mostly of fragmented volcanic rocks mentioned in the previous paragraph [Makariev, 2013]. The Vendian orogen was completely peneplained by the beginning of Paleozoic, when Siberian Platform covered the entire Central Taimyr belt [Sobolevskaya, Kaban'kov, 2014].

The question of possible involvement of the Central Taimyr belt into Vendian collision event was initiated by V.A. Vernikovskiy ([Vernikovskiy, 1996] and his subsequent publications) and discussed in [Priyatkina et al., 2017, p. 1650] The issue is important for understanding the tectonic evolution of the Siberian craton and its probable collision with the Baltica. The currently available information is still insufficient and it is unknown what continent or microcontinent the Siberian craton was came upon.

8. CONCLUSIONS

1. The granite pluton (Pregradnaya massif) that intruded the mid-Neoproterozoic orogenic deposits of Stanovaya-Kolosova Group was discovered in the SE part of the Central Taimyr belt. This finding was unex-

pected as these orogenic deposits are known to contain pebbles and boulders of similar porphyritic granites attributed to the Snezhnaya complex (845–825 Ma). As it turned out the U-Pb zircon age (SHRIMP) of the Pregradnaya granites amounted to 609 ± 2 Ma, which is quite uncommon for the belt.

2. The Pregradnaya massif and several other granitic bodies with similar textural and geochemical features, are located within deformation zone in the SE part of the Central Taimyr belt. Their position seemed to be consistent with the idea by V.A. Vernikovskiy that the 'accretionary belt' collided with the Siberian paleocontinent in Vendian or pre-Vendian times. However, other granite massifs in this deformation zone turned out to be older, belonging to the Snezhnaya (845–825 Ma) complex. According to published data, the late Neoproterozoic granites (640–610 Ma) are more abundant in the northern part of the belt, and are not confined to its southeastern margin.

3. The petrological and geochemical similarities between the Snezhnaya and Pregradnaya granites (845–825 Ma and 640–610 Ma, respectively) indicate that they were produced from similar melt sources in presumably similar conditions (P, T, H₂O). We suggest that the two granite magmatism 'flare-ups' separated by an interval of 200 Ma were related to two collision events that took place along the Taimyrian margin of the Siberian paleocontinent. The first event resulted in the massive intrusions of granite batholiths and the formation of a full-scale long-term orogeny. Its erosion products are widespread in the Central Taimyr belt. The second event was a more modest one. It is not known what terranes collided with Siberia during these events.

9. ACKNOWLEDGEMENTS

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INFORMATION ABOUT AUTHORS | СВЕДЕНИЯ ОБ АВТОРАХ

Aleksander B. Kuzmichev

Doctor of Geology and Mineralogy

Geological Institute of RAS
7 Pyzhevsky lane, Moscow 119017, Russia

✉ e-mail: nsi.kuzmich@yandex.ru

 <https://orcid.org/0000-0002-1595-1744>

Александр Борисович Кузьмичев

докт. геол.-мин. наук


Геологический институт РАН
119017, Москва, Пыжевский пер., 7, Россия

Maria K. Danukalova

Candidate of Geology and Mineralogy

Geological Institute of RAS
7 Pyzhevsky lane, Moscow 119017, Russia

e-mail: danukalovamk@yandex.ru

 <https://orcid.org/0000-0002-2974-3065>

Мария Константиновна Данукалова

канд. геол.-мин. наук

Геологический институт РАН
119017, Москва, Пыжевский пер., 7, Россия

Vasily F. Proskurnin

Doctor of Geology and Mineralogy

A.P. Karpinsky Russian Geological Research Institute (VSEGEI)
74 Sredny ave. V.O., Saint Petersburg 199106, Russia

e-mail: vasily_proskurnin@vsegei.ru

 <https://orcid.org/0000-0002-6424-0068>

Василий Федорович Проскурнин

докт. геол.-мин. наук

Всероссийский научно-исследовательский геологический
институт им. А.П. Карпинского
199106, Санкт-Петербург, Средний проспект В.О., 74, Россия

Aleksandra A. Bagaeva

Candidate of Geology and Mineralogy

A.P. Karpinsky Russian Geological Research Institute (VSEGEI)
74 Sredny ave. V.O., Saint Petersburg 199106, Russia

 <https://orcid.org/0000-0001-7675-2020>

Александра Александровна Багаева

канд. геол.-мин. наук

Всероссийский научно-исследовательский геологический
институт им. А.П. Карпинского
199106, Санкт-Петербург, Средний проспект В.О., 74, Россия

Nikolai I. Berezyuk

A.P. Karpinsky Russian Geological Research Institute (VSEGEI)
74 Sredny ave. V.O., Saint Petersburg 199106, Russia

 <https://orcid.org/0000-0002-3109-2943>


Николай Игоревич Березюк

Всероссийский научно-исследовательский геологический
институт им. А.П. Карпинского
199106, Санкт-Петербург, Средний проспект В.О., 74, Россия

Petr A. Gromov

A.P. Karpinsky Russian Geological Research Institute (VSEGEI)
74 Sredny ave. V.O., Saint Petersburg 199106, Russia

e-mail: petr_gromov@vsegei.ru

 <https://orcid.org/0000-0001-6957-5716>

Петр Андреевич Громов

Всероссийский научно-исследовательский геологический
институт им. А.П. Карпинского
199106, Санкт-Петербург, Средний проспект В.О., 74, Россия