Radiolarians, bivalves and the J/K boundary in the Birafu Formation, southern Kurosegawa Belt, Central Shikoku, SW Japan

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ABSTRACT: The Oxfordian-Berriasian Birafu Formation in the southern Kurosegawa Belt (Permian accretion terrane), Central Shikoku, SW Japan, yields micro- and mega-faunas that contribute to the biostratigraphy across the Jurassic-Cretaceous boundary. The type section consists of turbiditic sandstone and mudstone with an upward-fining tendency in the lower members (A1-A3). The middle members (B1-B2) are coarse sandstone and mudstone with muddy limestone intercalations, whereas the upper member C is characterized by a fine-grained turbiditic succession. Radiolarians define the assemblage zones (AZ): Kilinora spiralis AZ (Oxfordian) in member A2, Loopus primitivus AZ (Tithonian) in the upper part of member A3, and Pseudodictyomitra carpatica AZ (Berriasian - lower Valanginian) in members B2 and C. These give at least an Oxfordian - Berriasian total time-range to the Birafu Formation. The mixed marine and brackish bivalve assemblage of members B1 and B2 (lower part) comprises Grammatodon takiensis Tamura, Pterotrigonia toyamai (Yehara) and Ctenoides tosanus Kimura, the range of which is Late Jurassic to Earliest Cretaceous, and Aguilerella nagatoensis (Ohta), Miltha japonica Tashiro and Isocyprina japonica Tashiro and Kozai, having an Early Cretaceous range. The concurrent range of Jurassic to Cretaceous bivalves and exclusively Cretaceous species is significant for the clarification of bivalve evolution across the Jurassic - Cretaceous boundary. The Berriasian appearance of Cretaceous marine and non-marine bivalves takes place while Late Jurassic marine bivalves still survived. We consider Member B1 as Berriasian with the J/K boundary situated at its base.

Key words: Bivalves, radiolarians, J/K boundary, Kurosegawa, Japan

1. INTRODUCTION

Jurassic to Cretaceous non-marine, shallow marine and slope basin deposits are widely scattered over Japan. In Northeast Japan, some formations straddle the J/K boundary. The Isokusa Formation of the Karakuwa Group yields latest Jurassic and earliest Cretaceous ammonites with abundant bivalves (Sato, 1958; Hayami, 1960; Takahashi, 1973). The Ayukawa Formation of the Ojika Group yields Berriasian ammonites and bivalves. The Koyamada Formation of the Somanakamura Group yields ammonites and bivalves that indicate Tithonian and Berriasian ages. All these Late Jurassic and earliest Cretaceous bivalve faunas are very similar.

Southwest Japan is subdivided structurally into the Inner and Outer zones. In the Inner Zone, the J/K boundary falls within the non-marine deposits (Kobayashi and Suzuki, 1939; Hirano, 1971). The J/K boundary remains often concealed in the marine deposits of the Outer Zone (Tanaka, 1989), except in the Birafu Formation of the Sakashu Belt (a part of the Chichibu Superbelt), in the Monobe area of Central Shikoku (Ishida and Kozai, 2004).

Complementing the initial find of bivalves and ammonites by Yehara (1923), Kobayashi and Fukada (1947), the discovery of radiolarians, ranging from Oxfordian to Berriasian and the further identification of Jurassic marine bivalves and of an association of marine and brackish-water Cretaceous mollusks (Morino *et al.*, 1989; Kozai *et al.*, 2004) initiated the discussion of the remarkable correlation between radiolarians and particularly the earliest Cretaceous bivalves in the Birafu Formation.

The radiolarian biostratigraphy sets accurate age constraints for the completely different bivalve associations in the late Jurassic (Kimmeridgian-Tithonian) and Early Cretaceous (Hauterivian - Barremian). This exceptional situation may shed light on the evolution of the bivalve faunal association across the Jurassic-Cretaceous boundary in Japan. Detailed studies of Late Jurassic (Oxfordian-Tithonian) bivalves were carried out by Tamura (1959a-d, 1960ac), Hayami (1961) and Kimura (1951, 1956), whereas for the Early Cretaceous (Hauterivian-Barremian) bivalves, investigations were done by Hayami (1965a,b), Ohta (1973, 1974) and Tashiro and Kozai (1984, 1986, 1988, 1989, 1991, 1994,). Hayami (1961) identified three bivalve association groups: Hettangian-Aalenian, Bajocian-Bathonian and Callovian-Tithonian. He also pointed out that some of the younger Jurassic elements survived into the earliest Cretaceous. Later, the J/L faunal turnover was placed within the "Neocomian" (Hayami, 1975).

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2. GEOLOGICAL SETTING

The Inner and Outer zones of southwest Japan are separated by the Median Tectonic Line. The Outer Zone is subdivided into the Sanbagawa, Chichibu and Shimanto superbelts. In eastern Shikoku, the Chichibu Superbelt has been subdivided from north to south into five belts: the Kumosoyama Belt, the Masaki, Yoshigahira and Sakashu belts and the Nakagawa Belt (Ishida and Kozai, 2003). The Masaki Belt is composed of Permian accretionary complex (PAC) with Upper and Lower Cretaceous slope basin deposits (SL), e.g. the Monobegawa Group. The Yoshigahira Belt is a tectonic unit that was formed after the sedimentation of the Lower Cretaceous SL; it consists of the Siluro-Devonian Kurosegawa Tectonic Zone sensu stricto, Permian-Triassic and Jurassic accretionary complexes (AC), and Lower Cretaceous SL. The Sakashu Belt is composed of Permian AC with Siluro-Devonian blocks, uppermost Permian, Middle-Upper Triassic, Jurassic and Lower Cretaceous SLs. The Birafu Formation (Morino et al., 1989) is one of the slope-basin deposits within the Sakashu Belt (Kozai et al., 2004). The Birafu Formation extends between the Nankai Group (Yoshigahira Belt) and Monobegawa Group (Masaki Belt) to the North and the Nakagawa Belt to the South (Fig. 1).

3. DISTRIBUTION AND STRATIGRAPHY OF THE BIRAFU FORMATION

The Birafu Formation is exposed in Central Shikoku, east of Tosamyamada (Kochi Prefecture), striking generally ENE-WSW and dipping steeply south. The type section of the Birafu Formation is located at Ogawa along the Nishinokawa River, where a thickness of 725 m was measured (Fig. 2). The formation is subdivided into six members (Fig. 3).

Member A1 is composed of alternating beds of dark gray mudstone and sandstone. In the lower part of the member, 1 cm to 10 cm thick sandstone layers are well graded. Parallel and current ripple laminations are also observed. In the upper part of the member, the sandstone beds become thicker. The total thickness of the unit A1 is approximately 40 m.

Member A2 contains medium-grained massive sandstone, muddy sandstone, sandy mudstone, mudstone and tuff. The muddy sandstone bed of unit A2 contains fragments of bivalves. The Kimmeridgian ammonite *Ataxioceras kurisakense* Kobayashi and Fukada was described from this member by Kobayashi and Fukada (1947). The member is topped by a mudstone (32601R) with tuff intercalation. The mudstone yields Oxfordian radiolarians. This unit is approximately 80 m thick.



Fig. 1. Tectonic subdivision of the Chichibu Superbelt in East Shikoku (modified after Ishida and Kozai, 2003; Kozai and Ishida, 2003)



Fig. 2. Map showing the distribution of the Birafu Formation in Central Shikoku (modified after Kozai et al., 2004).

Member A3 is characterized by sandy mudstone and mudstone. The upper part of the member is again intercalated with tuff and the mudstone yields rare bivalves and Tithonian radiolarians (32602R). This unit is approximately 40 m thick.

The bivalve yielding Member B1 consists of conglomerate, sandstone, mudstone, limestone and alternating beds of sandstone and mudstone. The conglomerate contains chert, shale and granite pebbles. The medium grained poorly sorted sandstone beds contain coaly inclusions and abundant nonmarine and marine mollusks. Limestone beds are ooliticpeloidal and bioclastic grainstones (Morino, 1993). This unit is approximately 170 m thick.

Member B2 is composed of alternating sandstone, mudstone and limestone beds. The well-sorted, fine- to mediumgrained sandstone is arkosic. The mudstone is intercalated with tuff and yields Berriasian radiolarians (32604R, 32605R). The limestone is sandy and laminated. This unit is approximately 115 m thick.

Member C is composed of sandstone, mudstone, tuff and alternating beds of sandstone and mudstone. The sandstone is medium-grained, graded and ranging from 1 cm to 10 cm in thickness. Parallel and current ripple laminations are observed. The mudstone yielded radiolarians (32606R, 32607R, 32608R, 32609R). This unit is approximately 280 m thick.

4. RADIOLARIAN BIOSTRATIGRAPHY AND GEO-LOGICAL AGE OF THE BIRAFU FORMATION

The Late Jurassic–Early Cretaceous radiolarian biostratigraphy in Japan was established by Matsuoka and Yao (1986). In the type section of the Birafu Formation, three assemblage zones were found in ascending order: the *Kilinora spiralis* Assemblage Zone, the *Loopus primitivus* Assemblage Zone and the *Pseudodictyomitra carpatica* Assemblage Zone. The *Kilinora spiralis* Assemblage Zone is found near the top of Member A2 (32601R). The *Loopus primitivus* Assemblage Zone occurs in the uppermost part of Member A3 (32602R). The *Pseudodictyomitra carpatica* Assemblage Zone ranges from the base of Member B2 (32604-5R) and throughout Member C (32606-9R). The presence of Late Jurassic beds in the Nishinokawa section was first established by the occurrence of *Ataxioceras kurisakense*, a Lower Kimmeridgian ammonite (Kobayashi and Fukada, 1947).

4.1. Kilinora spiralis Assemblage Zone

Although the details of the top and base of this zone have not yet been determined, the *Kilinora spiralis* Assemblage Zone is characterized by *Kilinora spiralis* (Matsuoka) and the common occurrence of *Hsuum maxwelli* Pessagno, *Sti*-



Fig. 3. Lithological section of the Birafu Formation with location of Radiolarian assemblage zones.

chocapsa robusta Matsuoka, Tethysetta dhimenaensis (Baumgartner), Tricolocapsa conexa Matsuoka, Tricolocapsa sp. cf. T. parvipora Tan, Archaeodictyomitra amabilis Aita, Eucyrtidiellum nodosum Wakita, Ristola procera (Pessagno), Stylocapsa tecta Matsuoka, and Tricolocapsa plicarum Yao (Fig. 4). Cinguloturris carpatica Dumitrica and Mirifusus guadalupensis Pessagno are rare.

Correlation and Age: The present radiolarian assemblage corresponds to the *Kilinora spiralis* Zone (Matsuoka and Yao, 1986) of Oxfordian age (Late Jurassic). Based on the co-occurrence of *Eucyrtidiellum ptyctum* (Riedel and Sanfilippo) with *Kilinora spiralis* (Matsuoka), the top horizon (32601R) of Member A2 corresponds to Subzone 2γ (Middle Oxfordian) of Hull (1997) at Stanley Mountain, California. Baumgartner et al. (1995) considered *Kilinora spiralis* (Matsuoka) to range from Middle Bathonian to Early Callovian (Unitary Association zones 6 and 7).

4.2. Loopus primitivus Assemblage Zone

Although the details of the top and base of this zone have not yet been determined, the Loopus primitivus Assemblage Zone is characterized by the common occurrence of Loopus primitivus (Matsuoka and Yao) associated with Cinguloturris carpatica Dumitrica, Protunuma japonicus Matsuoka and Yao, Ristola altissima (Rüst), Svinitzium depressum (Baumgartner), Svinitzium mizutanii Dumitrica, Svinitzium pseudopuga Dumitrica, and Xitus gifuensis Mizutani. Archaeodictyomitra apiarium (Rüst), Archaeodictyomitra brouweri (Tan), Archeodictyomitra minoensis (Mizutani), Loopus doliolum Dumitrica, Loopus primitivus Matsuoka and Yao, Loopus yangi Dumitrica, Neorelumbra kiesslingi Dumitrica, Pantanellium lanceola (Parona), Protunuma japonicus Matsuoka and Yao, Solenotryma (?) ichikawai Matsuoka and Yao, Svinitzium mizutanii Dumitrica, Svinitzium pseudopuga Dumitrica, Tethysetta boesii (Parona), Tethysetta pygmaea Dumitrica, and Xitus gifuensis Mizutani appear first in this zone (Fig. 4). Eucyrtidiellum ptyctum (Riedel and Sanfilippo) is rare, whereas Hsuum maxwelli Pessagno and Pseudodictyomitra carpatica (Lozyniak) are absent in the Loopus primitivus Assemblage Zone.

Correlation and Age: The *Loopus primitivus* Assemblage Zone corresponds to the *Loopus primitivus* Zone of Matsuoka (1995b), and the age is regarded Tithonian (Late Jurassic). In Japan and western Pacific region (ODP Site 801), the *Loopus primitivus* Zone is recognized by the abundant occurrence of *Loopus primitivus* (Matsuoka & Yao), above the last appearance of *Hsuum maxwelli* Pessagno and below the first appearance of *Pseudodictyomitra carpatica* (Matsuoka, 1995a).

4.3. Pseudodictyomitra carpatica Assemblage Zone

Although the top and base of this zone have not yet been determined, the assemblage found in the Nishinokawa section near the base of Member B2 can be attributed to the Pseudodictyomitra carpatica Assemblage, characterized by the common occurrence of Pseudodictyomitra carpatica (Lozyniak). The first appearance of *Pseudodictyomitra* carpatica (Lozyniak) accompanied by Archaeodictyomitra mitra Dumitrica, Archaeodictyomitra pseudomulticostata (Tan), Sethocapsa kaminogoensis Aita, Sethocapsa pseudouterculus Aita, and Tethysetta columna (Rüst). Archaeodictyomitra apiarium (Rüst), Archaeodictyomitra mitra Dumitrica, Loopus doliolum Dumitrica, Loopus vangi Dumitrica, Neorelumbra kiesslingi Dumitrica, Pantanellium lanceola (Parona), Svinitzium depressum (Baumgartner), Svinitzium mizutanii Dumitrica, and Svinitzium pseudopuga Dumitrica are relatively common in the zone. Archaeodictvomitra brouweri (Tan), Archeodictvomitra minoensis (Mizutani), Mirifusus mediodilatatus (Rüst),

JURASSIC		CRETACEOUS						PERIOD			
OXFORDIAN	TITHONIAN	BERRIASIAN						AGE			
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A2	A3	B	220050	22606 D	22607 D						
32001K	32002K	32004K	32000K	32000 R	32007R	32008 R	32009 R	Archaeodictuomitra amabilia Aita			
I.	С	С	С		С			Archaeodictyomitra aniarium (Rüst)			
	C C	R						Archaeodictyomitra browneri (Tan)			
	R	R						Archeodictyomitra minoensis (Mizutani)			
		C	C	C	C	C		Archaeodictyomitra mitra Dumitrica			
		R			<u> </u>	•		Archaeodictyomitra pseudomulticostata (Tan)			
					R	C		Archaeodictyomitra tumandae Dumitrica			
R								Archaeodictyomitra sp. aff A pseudoscalaris (Tan)			
					С			<i>Becus rotula</i> Dumitrica			
R	C		C					<i>Cinguloturris carpatica</i> Dumitrica			
	R	R						Cinguloturris sp. aff. S. carpatica Dumitrica			
				C	С			Cinguloturris cylindra Kemkin & Rudenko			
					С			Cinguloturris sp. aff. S. cylindra Kemkin & Rudenko			
C	R							Dictyomitrella (?) sp. aff. D. kamoensis Mizutani & Kido			
C								<i>Eucyrtidiellum nodosum</i> Wakita			
R	R	R						Eucyrtidiellum ptyctum (Riedel & Sanfilippo)			
				R				<i>Eucyrtidiellum pyramis</i> Aita			
C								Hsuum maxwelli Pessagno			
C								Hsuum spp.			
R								<i>Kilinora spiralis</i> (Matsuoka)			
	R	R	R	C	C			<i>Loopus doliolum</i> Dumitrica			
	C	R	R		R			<i>Loopus primitivus</i> (Matsuoka & Yao)			
	R	R	C				C	<i>Loopus yangi</i> Dumitrica			
R		R						Mirifusus cheodes (Renz)			
R								Mirifusus guadalupensis Pessagno			
				R				Mirifusus mediodilatatus (Rüst)			
	R			C C			C	Neorelumbra kiesslingi Dumitrica			
	<u> </u>	<u> </u>		<u> </u>	0			Obesacapsula rotunda (Hinde)			
	ĸ	к		ĸ	ĸ			Pantane/lium lanceo/a (Parona)			
	0				ĸ			Praecaneta cosmocon/ca (Foreman)			
	G				0		0	Protunuma japonicus Matsuoka & Yao			
			D				<u> </u>	Pseudodictyomitra altiturris Dumitrica			
		G	ĸ	U	U	U	U	Pseudodictyomitra carpatica (Lozyniak)			
				ĸ	0			Pseudodictyomitra sp. att. P. thurowi Dumitrica			
	0				U			<i>Pietola altianima (Dilata)</i>			
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	0	R			5	Svinitzium mizutanii Dumitrice					
\vdash	C.					Svinitzium nseudonuga Dumitrica					
	R		R					Svinitzium sp.			
	R	R	R R					Tethysetta boesii (Parona)			
		R	R Jathysetta columna (Diset)			Tethysetta columna (Riist)					
		N N						reurysella corunna (Rust)			

Fig. 4. Occurrence and inferred range of radiolarians in the type section of the Birafu Formation. Ks: Kilinora spiralis Assemblage Zone; Lp: Loopus primitivus assemblage Zone; Pseudodictyomitra carpatica Assemblage Zone; R: rare (<1%); C: common (1-5%); A: abundant (5%<) in total number of counted specimens; gray bar showing the inferred range of species in the section.

Tethysetta boesii (Parona), Tethysetta pygmaea Dumitrica, the zone. The lower part of the zone (Member B2) still and Tethysetta usotanensis (Tumanda) are relatively few in yields Cinguloturris carpatica Dumitrica, Eucyrtidiellum

ptyctum (Riedel and Sanfilippo), Loopus primitivus ÅiMatsuoka and Yao, and Solenotryma (?) ichikawai Matsuoka and Yao (Fig. 4). The younger part of the zone (in Member C) is characterized by the first occurrence of Archaeodictyomitra tumandae Dumitrica, Becus rotula Dumitrica, Cinguloturris cylindra Kemkin and Rudenko, Eucyrtidiellum pyramis Aita, Pseudodictyomitra altiturris Dumitrica, Pseudodictyomitra sp. aff. P. thurowi Dumitrica, Pseudoxitus sp. aff. P. seriola Dumitrica, Ristola sp. aff. R. cretacea (Baumgartner), Sethocapsa (?) subcrassitestata Aita, and Xitus robustus Wu (Fig. 4).

Correlation and Age: In general, the Pseudodictyomitra carpatica Assemblage Zone corresponds well with the Pseudodictyomitra carpatica Zone (Matsuoka, 1995b) and the Ditrabs sansalvadorensis Zone (Aita and Okada, 1986), the ages of which are regarded as latest Tithonian to earliest Cretaceous (Berriasian-Early Valanginian). In the Nishinokawa section, Sethocapsa pseudouterculus (Aita) and S. kaminogoensis (Aita) appear with Pseudodictyomitra carpatica (Lozyniak) nearly at the base of the zone. In terms of UA (unitary associations), Jud (1994) established for Pseudodictyomitra carpatica (Lozyniak) a Tithonian to Barremian range, or more precisely Late Kimmeridgian/ Early Tithonian to Early Barremian as established by Baumgartner et al. (1995). Matsuoka (1995b) estimated that the first appearance of Pseudodictyomitra carpatica (Lozyniak) took place slightly below the J/K boundary. In the Breggia Gorge (Switzerland) Sethocapsa pseudouterculus (Aita) appears according to Aita and Okada (1986), near the base of the Ditrabs sansalvadorensis Zone, the base of which is also slightly below the J/K boundary. S. kaminogoensis appears however above the J/K boundary. In addition, Dumitrica et al. (1997) reported the occurrences of Archaeodictyomitra mitra Dumitrica, Archaeodictyomitra pseudomulticostata (Tan), and Tethysetta columna (Rüst) from Berriasian and younger samples from Mashirah Ophiolite sections in Oman. Therefore, we regard the base of the present Pseudodictyomitra carpatica Assemblage Zone to be Berriasian, and that the J/K boundary is probably below the base of Member B2.

Concerning the younger age of the zone, Dumitrica *et al.* (1997) reported the occurrences of *Archaeodictyomitra tumandae* Dumitrica, *Becus rotula* Dumitrica, *Cinguloturris cylindra* Kemkin and Rudenko, *Pseudodictyomitra altiturris* Dumitrica from Berriasian–Valanginian samples of the Mashirah Ophiolite sections in Oman. According to Matsuoka and Yang (2000), the first appearance of *Cecrops septemporatus* (Parona) is situated near the base of the *Cecrops septemporatus* Zone, and is Valanginian. We have not found *Cecrops septemporatus* (Parona) in our section. As such, we still consider the uppermost part of Member C as part of the *Pseudodictyomitra carpatica* Assemblage Zone, the age of which can be regarded to be older than Late Valanginian.

5. DEPOSITIONAL ENVIRONMENT OF THE BER-RIASIAN MEMBER B1

Member B1 is well exposed along the banks of the Nishinokawa River, where it yields abundant bivalves, gastropods and corals. It consists of alternating beds of limestone and muddy sediments (Fig. 4). Although faults are present in the area, the exposed sequence seems mostly continuous. The upper contact of the limestone beds with the overlying muddy deposits are sharp. This contrasts with the gradual nature of the lower contact of the limestone beds. We have subdivided member B1 into seven cycles (sub-units I-VII) based on lithofacies.

The lower three consist of the regular cyclic recurrence of mudstone at base, followed by fossiliferous muddy sandstone, calcareous sandstone with nerineacean gastropods and ooid-bearing limestone at top. In each of the three lower subunits, the proportion of clastic particles decreases upward. While the base of sub-unit I is still part of Member A3, containing the L. primitivus radiolarian AZ (Tithonian), the base of Member B1 corresponds to the base of the fossiliferous sandy mudstone that contains disarticulated and poorly preserved bivalves. The other layers are fossiliferous calcareous sandstone bed with undeterminable nerineacean gastropods, ooid-peloid-grainstone, calcareous ooid-grainstone, sandstone, ooid-grainstone/packstone and thin limy mudstone containing algal-coated grains. Within sub-units II and III, the stratigraphy becomes unclear due to faulting, but the observed lithologic succession appears similar to that of sub-unit I. The limestone beds of sub-unit III along the river, probably correspond to the limestone that Morino (1993) described in a cliff about 20 m above the southern riverbank.

In the middle part of member B1, sub-units IV and V are thinner than the lower sub-units. Detailed stratigraphy of the limestone is unclear, although similar lithofacies to that in the lower sub-units.

In the upper part of member B1, Sub-unit VI consists of alternating beds of sandy mudstone and thin- and very finegrained sandstone, calcareous sandstone and pisolite. Subunit VII consists of a 50 m thick sequence of muddy and sandy sediments, showing upward coarsening of the sequence. Higher up, the base of Member B2 is exposed, consisting of a basal sandstone, overlain by calcareous fine-grained sandstone and boundstone.

6. BIVALVE FAUNA

Bivalves are rare in members A2, A3 and B2, and only Member B1 yields abundant bivalves. As such the geological age discussion will center on the bivalves from Member B1.

Members A2 and A3 are well dated as Jurassic by radiolarians and Member B1 starts immediately above the horizon yielding Tithonian radiolarians. Member B1 is overlain



Fig. 5. Route map along the type section of the Birafu Formation, Nishinokara River in Kahoku. Also shown are the Micro- and Mega fossil localities.

by member B2 that contains a Berriasian radiolarian fauna. This indicates that the J/K boundary must be within Member B1. Abundant bivalves are mostly found in the sandstone beds that underlie the limestones in most of the cycles of Member B1. Bivalves are rarely found in the mudstone beds. Bivalve distribution occurs in the six sub-units of Member B1 (Fig. 5).

6.1. Sub-unit I

The sandstone yielded abundant but generally poorly preserved forms of *Grammatodon takiensis* Kimura, *Cucullaea* sp. aff. *C. Acuticarinata* Nagao, *Arcomytilus* sp. aff. *A. Laitmairensis* (de Loriol), *Aguilerella nagatoensis* (Ohta), *Ctenoides tosanus* (Kurata and Kimura), *Pterotrigonia toyamai* (Yehara), *Miltha japonica* Tashiro, *Protocardia* sp. cf. *P. tosensis* Kimura and *Isocyprina japonica* Tashiro and Kozai. *Grammatodon takiensis* Kimura and *Ctenoides tosanus* (Kurata and Kimura) occur in the Jurassic to the Lowest Cretaceous formations of Japan, e.g. the Torinosu Group in Shikoku, Nakanosawa and Koyamada formations of the Somanakamura Group and Isokusa Formation of the Karakuwa Group in NE Japan (Kimura, 1951; Tamura, 1960a; Masatani and Tamura, 1959; Hayami *et al.*, 1960). *Pterotrigonia toyamai* (Yehara) was originally described from the Birafu Formation by Yehara (1923). It was also found in the Oxfordian - Kimmeridgian formations of the Torinosu Group in Shikoku, as well as in the Kurosaki and Ebirase formations in Kyushu (Tamura, 1960a, 1984; Tamura and Murakami, 1986). The Cretaceous is represented by the species *Aguilerella nagatoensis* (Ohta), *Miltha japonica* Tashiro and *Isocyprina japonica* Tashiro and Kozai that occur in the non-marine and marine formations (Tashiro and Kozai, 1988, 1989).

6.2. Sub-unit II

The sandstone bed of this sub-unit yielded *Grammatodon* takiensis Kimura, Isocyprina japonica Tashiro, Corbula globosa Tamura, C. imamurae Hase and Caestocorbula morinoi Tashiro and Kozai. Corbula globosa Tamura occurs in the Upper Jurassic Sakamoto Formation in Kyushu (Tamura, 1959a, 1959b) and in the Torinosu Group in Shikoku (Tamura, 1960a) as well as in Upper Jurassic to Berriasian formations, like the Nakanosawa and Koyamada formations in NE Japan. C. imamurae Hase was described from the Lower Cretaceous Yoshimo Formation in Chugoku (Hase, 1960).

6.3. Sub-unit III

The sandstone yielded *Eomiodon nipponicus* Ohta and *Corbula imamurae* Hase that occur in Lower Cretaceous nonmarine formations (Hase, 1960; Ohta, 1973; Kozai *et al.*, 2005).

6.4. Sub-unit IV

The fossiliferous sandstone is overlain by limestone along the river bed. It contains the Aguilerella nagatoensis-Corbula imamurae Association, a highly diversified brackishwater molluscan association. The horizon is identical with the one described by Morino (1993). This sub-unit yields Parallelodon sp. cf. P. koikensis Tamura, Aguilerella nagatoensis (Ohta), Parvamussium habunokawense (Kimura), Ctenoides tosanus (Kurata and Kimura), Miltha japonica Tashiro, Protocardia sp. cf. P. tosensis Kimura, Isocyprina japonica Tashiro and Kozai, Linearia sp. aff. L. nankaiana Tashiro and Kozai, Eomiodon nipponicus Ohta, Eomiodon kumamotoensis Tamura, Crenotrapezium kitakamiense Hayami, Corbula globosa Tamura and Caestocorbula morinoi Tashiro and Kozai. Eomiodon kumamotoensis Tamura was described from the Upper Jurassic Sakamoto Formation in Kyushu by Tamura (1959c). Crenotrapezium kitakamiense was described from the Lower Cretaceous in the Jusanhama Group in NE Japan by Hayami (1960).

6.5. Sub-unit V

The sub-unit yields *Pterotrigonia* sp., *Isocyprina japonica* Tashiro *and Ostrea* sp.

6.6. Sub-unit VI

The sub-unit yields Aguilerella nagatoensis (Ohta), Protocardia sp. cf. P. tosensis Kimura, Eomiodon nipponicus Ohta, Eomiodon kumamotoensis Tamura, Crenotrapezium kitakamiense Hayami and Corbula imamurae Hase. While four species, Aguilerella nagatoensis (Ohta), Eomiodon nipponicus Ohta, Crenotrapezium kitakamiense Hayami and Corbula imamurae Hase are Lower Cretaceous, two species, Protocardia sp. cf. P. tosensis Kimura and Eomiodon kumamotoensis Tamura are still Jurassic.

7. THE J/K BOUNDARY BIVALVE FAUNAL TURN-OVER

The bivalves found in Member B1 are characteristically mixed Jurassic marine and Cretaceous marine and non-marine (Fig. 6). Hayami (1961) summarizing Jurassic marine bivalve faunas in Japan, included *Grammatodon takiensis Kimura*, *Arcomytilus* sp. cf. *A. laitmairensis* (de Loriol), *Ctenoides tosanus* (Kurata and Kimura), *Pterotrigonia toyamai* (Yehara), *Protocardia tosensis* Kimura, *Eomiodon kumamotoensis*

Tamura, Corbula globosa Tamura, Parvamussium habunokawense (Kimura). With exception of Pterotrigonia toyamai (Yehara), all these marine species also occur in the Berriasian formations. Furthermore, the marine *Isocyprina japonica* Tashiro and Kozai and Miltha japonica Tashiro, as well as the brackish Eomiodon nipponicus Ohta, Crenotrapezium kitakamiense Hayami, Aguilerella nagatoensis (Ohta) and Corbula imamurae Hase are exclusively Early Cretaceous in Japan. The presence of both the marine and brackish Cretaceous forms from the very base of member B1 strongly suggests that the member can be regarded as Berriasian. Since Hayami (1961, 1975) pointed out that some Upper Jurassic marine bivalves survived until the Berriasian, the concurrent range of Jurassic and Lower Cretaceous bivalves is not surprising. Although Although Ohta and Moji (1976) assigned on Early Cretaceous age to the Uminoura Formation based on Pulsidis globosa Tamura and Eomiodon kumamotoensis Tamura (Late Jurassic-Berriasian) and Eomiodon matsumotoi Ohta and Pulsidis nagatoensis Ota (Early Cretaceous), regarding this formation as Early Cretaceous, Ohta and Sakai (2003) refers to this Formation as Middle or Late Jurassic. Therefore the Birafu Formation is a key to the understanding of the J/K bivalve faunal turnover.

Some important ecological and taxonomic changes can be recognized for the faunal turnover across the J/K boundary as inferred from Member B1. The most important change is the occurrence of *Linearia* sp. aff. *nankaiana* Tashiro and Kozai. Until now, the first occurrence of the family Tellinidae was considered as Early Hauterivian based on *Linearia subhercynica* from Hannover, Germany and *Linearia subconcentrica* from Paris Basin, France, (Skelton and Benton, 1993). As discussed in this paper, *Linearia* sp. aff. *L. nankaiana* Tashiro and Kozai occurred on Member B1, which is considered as Berriasian. Therefore this is the oldest fossil record of the Tellinidae. In contrast to the labial-palp deposit feeders protobranchians, which appeared in early Paleozoic, the Tellinidae are new deposit-feeders using siphon for feeding.

The occurrence of *Caestocorbula morinoi* (Yehara), is also important because this is probably the first record of the subfamily Caestocorbulinae, as well as being first of the genus *Caestocorbula*. This species is also paleoecologically important, because this may be one of the first heterodont bivalves that developed deep pallial sinus.

Also, there are several genera which are considered to have occurred first here, such as *Isocyprina* spp. and *Linearia* spp.

The first appearance of these modern-type life habits in the bivalves in the Member B1 suggests new environmental conditions prevailing in this age, which facilitated such evolutionary diversification. However, further detailed study is needed to reveal the environmental and biotic changes that occurred around the J/K boundary.



Fig. 6. Lithological section of the Members B1 and B2 and occurrence list of bivalves (modified after Kozai et al., 2004).

8. CONCLUSIONS

1. Oxfordian to Berriasian radiolarian assemblages provide biostratigraphic constraints for the Birafu Formation.

2. Cretaceous forms from the base of Member B1 strongly suggests that this member can be regarded as Berriasian and that the J/K boundary can be placed at base B1.

3. Very different Lower Cretaceous and Upper Jurassic bivalves co-occur in the Berriasian of the Birafu Formation. Hence, the Berriasian is regarded as a transitional zone of Jurassic to Cretaceous bivalve taxa.

4. The bivalve fauna of Member B1 suggests a Lower Cretaceous (Berriasian) shallow marine environment in which brackish forms have been transported, probably by floods.

5. The concurrence of Jurassic and Cretaceous marine taxa suggests that the survival of the Jurassic marine taxa in the Lower Berriasian is a feature related to the Jurassic-Cretaceous turnover and that the Berriasian is a transitional zone in Jurassic - Cretaceous bivalve evolution.

6. First appearance of some modern life habits in bivalves are recorded in the B1 member of the Birafu Formation, such as siphonate deposit-feeding heterodont (Tellinidae) and siphonate suspension-feeding heterodont with deep pallial sinus (Caestocorbulinae). These suggest a widespread new environment near J/K boundary. This new environment facilitated the evolution of new life habit groups in Bivalves.

Species	Kimmeridgian	Tithonian	Berriasian	Valanginian	Hauterivian
<i>Parallerodon koikensis</i> Tamura					
<i>Grammatodon takiensis</i> Tamura					
Cucullaea acticarinata Nagao					
Arcomytilus laitmairensis (De Loriol)					
Aguilerella nagatoensis (Ohta)					
Parvamussium habunokawense (Kimura)					
Ctenoides tosana Kimura					
Pterotrigonia toyamai (Yehara)					
Protocardia tosensis Kimura					
<i>Isocyprina japonica</i> Tashiro					
<i>Linearia aff. nankaiana</i> Tashiro and Kozai					
<i>Miltha japonica</i> Tashiro					
<i>Eomiodon kumamotoensis</i> Tamura					
<i>E. nipponicus</i> Ohta					
<i>Crenotrapezium kitakamiense</i> Hayami					
<i>Corbula grobosa</i> Tamura					
C. imamurae Hase					
Caestcorbla morinoi Tashiro and Kozai					

Fig. 7. Kimmeridgian to Hauterivian range-chart of the Birafu Bivalve fauna, SW Japan.

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TAXONOMIC NOTES OF BIVALVES

The studied bivalve specimens are kept in Naruto University of Education (NU) and radiolarian specimens are deposited in Tokushima University.

Subclass PTERIOMORPHIA Beurlen, 1944 Order ARCOIDA Stoliczka, 1871 Family PARALLELODONTIDAE Dall, 1898 Genus Parallerodon Meek and Worthen, 1866 Parallelodon sp. cf. P. koikensis Tamura, 1959 Pl. 2, fig. 12

Material.- NU39-0085, internal mold of left valve from loc. 32631.

Remarks.- The poorly preserved specimen from the Birafu Formation is similar to Parallelodon koikensis described by Tamura (1959d) from the Nakanosawa Formation, Soma Group of Northeast Japan. It has an elongate outline and a distinct posterior carina.

Occurrence.- Loc. 32631.

Range.- Kimmeridgian to Berriasian.

Genus Grammatodon Meek and Hayden, 1861 Grammatodon takiensis Kimura, 1956 Pl. 2, figs. 4-5

Grammatodon takiensis Kimura, 1956, p. 82, pl. 1, figs. 5,6; Tamura, 1959b, p. 54, pl.6, figs. 1,2; Tamura, 1959d, p. 172, pl. 19, figs. 4-6; Hayami, Sugita and Nagumo, 1960, p. 91, pl. 8, figs. 6,7; Hayami, 1975, p. 29.

Material.- NU39-0038a,b, right valves from Loc.32655 and NU39-0039, left valve from Loc. 32626.

Remarks.- The specimens are identical with Grammatodon takiensis in the subrectangular outline with fairly distinct posterior carina.

Occurrence.- Loc. 32626, 32655. Range.- Kimmeridgian to Berriasian.

> Family CUCULLAEIDAE Stewart, 1930 Genus Cucullaea Lamarck, 1801 Cucullaea sp. aff. C. acuticarinata Nagao Pl. 2, fig. 33

Material.- NUE39-0040, left valve from 32626.

Remarks.- This specimen is characterized by the strongly inflated valve and the sharp posterior carina. These features are similar to Cucullaea acuticarinata Nagao from the Miyako Group in NE Japan and the Hagino Formation of the Nankai Group in Shikoku. However, the umbo of this specimen situates near the center of shell-length and is smaller than that of C. acuticarinata Nagao.

Occurrence.- Loc.32626.

Range.- Berriasian.

Order MYTILOIDA Férussac, 1822 Family MYTILIDAE Rafinesque, 1815 Genus Arcomytilus Agassiz, 1842 Arcomytilus sp. aff. A. laitmairensis (de Loriol) Pl. 2, figs. 21-22

Material.- NUE39-0086, left valve from 32626; NU39-0087, right valve from loc. 32626; NU39-0088, internal mold of right valve from loc.32626.

Remarks.- The surface of our incomplete specimens has a clear ornamentation. The radial ribs are stronger than those of Arcomvtilus laitmairensis (de Loriol). The Birafu specimens are similar to Tamura's specimens from the Soma Group in Northeast Japan and Torinosu Group in Shikoku.

Occurrence.- Loc.32626.

Range.- Kimmeridgian to Berriasian.

Order PTERIOIDA Newell, 1965 Family BAKEVELLIDAE King, 1850

Genus Aguilerella Chavan, 1951 Aguilerella (Yoshimopsis) nagatoensis (Ohta) Pl. 2, figs. 34-35

Gervillia shinanoensis Yabe and Nagao. Kobayashi and

Suzuki 1939, p. 217, pl. 13, figs. 20-22, pl. 14, figs. 16-19. Bakevelloides (Yoshimopsis) nagatoensisOhta, 1974, p. 81, text-figs. 1-4, pl. 1, figs. 1-11.

Balevellia (Yoshimopsis) nagatoensis (Ohta), Hayami 1975, p. 46, pl. 9, fig. 2.

Aguilerella (Yoshimopsis) nagatoensis (Ohta), Tashiro 1992, p. 72, pl. 18, fig. 6; Tanaka et al. 2000, fig. 4, 8-10.

Material.- NU39-0050 and 0051, left valves from loc. 32631; MU39-0052, left valve from loc.32656.

Remarks.- The specimens from the Birafu Formation are poorly preserved, but can be referred to Aguilerella (Yoshi-

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mopsis) *nagatoensis* (Ohta) by their trapezoidal to subtrigonal outline and ligamental characters. Specimens from the Birafu Formation show changes in their morphological character. They include one-to-three-ligament pits. We have 20 specimens but all of the specimens are left valves and are worn out. This fact shows that they are allochthonous.

Occurrence.- Locs. 32626, 32631, 32656. Range. - Berriasian to Hauterivian.

Family PROPEAMUSSIIDAE Abbott, 1954 Genus Parvamussium Sacco,1897 Parvamussium habunokawense (Kimura) Pl. 2, fig. 23

Propeamussium habunokawensis Kimura, 1951, p. 344, pl. 1, figs. 14-15; Variamussium habunokawense (Kimura), Tamura 1959b, p. 60, pl. 6, figs. 20-22; Tamura 1960a, p. 236; Variamussium cf. habunokawense (Kimura), Hayami, Sugita and Nagumo 1960, p. 92, pl. 8, figs. 11-13; Parvamussium habunokawense (Kimura), Hayami 1975, p. 83; Parvamussium cf. habunokawense (Kimura), Tamura and Murakami 1987, pl. 1, figs. 3-5.

Material.- NU39-0062, external mold of left valve from 32631.

Remarks.- The specimen from the Birafu Formation is an imperfect external mold, but the surface ornamentation is clearly preserved. Ribs include 9 strong internal radial ribs and 10 secondary radial ribs. Though the radial ribs seem more irregular than those of *P. habunokawense*, our specimen shares the essential characters of *Parvamussium habunokawense* (Kimura) from the Torinosu Group in Shikoku. In this respect, our specimen is also similar to a specimen from the Kogoshiro Formation in NW Japan.

Occurrence.- Locs. 32631, 32635.

Range. - Kimmeridgian to Berriasian

Family LIMIDAE Rafihesque, 1815 Genus *Ctenoides* Mörch, 1853 *Ctenoides tosanus* (Kurata and Kimura) Pl. 2, figs. 13-14

Lima (*Ctenoides*) *tosana* Kurata and Kimura, *in* Kimura, 1951, p. 349, pl. 1, fig. 22; Tamura 1959b. p. 62, pl. 6, figs. 44-47; Tamura 1959d. p. 177; Tamura 1960a. p. 238.

Ctenoides tosanus (Kurata and Kimura). Hayami, 1975, p. 87.

Material. NU39-0041 and 0042, right valves from loc. 32626; NU39-0043, left valve from loc.32626.

Remarks.- This species is usually small, but the specimens from the Birafu Formation are relatively large, 39.1 mm in height and 22.4 mm in length. The surface ornamentation consists of numerous fine radial ribs.

Occurrence.- Locs. 32626, 32631.

Range. - Kimmeridgian to Berriasian.

Subclass PALAEOHETERODONTA Newell, 1965 Order TRIGONIOIDA Dall, 1889 Family TRIGONIIDAE Lamarck,1819 Genus Pterotrigonia van Hoepen, 1929 *Pterotrigonia toyamai* (Yehara) Pl. 2, figs. 28-29

Torigonia toyamai Yehara, 1923, p. 78, pl. 9, figs. 4,5; Yehara 1927, p. 35, pl. 3, fig. 9; *Linotrigonia toyamai* (Yehara), Kobayashi 1954, p. 77; Kobayashi 1956, p. 6, pl. 1, figs. 10-12; Tamura 1960, p. 234; Hayami 1975, p. 112; Tamura 1984, pl. 2, figs. 18-22; Tamura and Murakami 1986, Pl. 2, figs. 8-13; *Pterotrigonia (Pterotrigonia) toyamai* (Yehara), Tashiro 1992, p. 160, pl. 43, figs. 7,11.

Material. NU39-0064-0067, left valves from loc.32626; NU39-0068, right valve from 32626.

Remarks.- All the specimens from the Birafu Formation are poorly preserved, but surface ornamentation can be observed. They have a few oblique costae on the area near carina, a feature that is characteristic of the Genus *Pterotrigonia*.

Occurrence.-Loc. 32626. Range.- Oxfordian to Berriasian.

> Subclass HETERODONTA Neumayr, 1884 Order Veneroida Adamus and Adamus, 1856 Family LUCINIDAE Fleming, 1828 Genus *Miltha* Adamus and Adamus, 1857 *Miltha japonica* Tashiro Pl. 2, fig. 32

Miltha japonica Tashiro, 1990, p. 13-14, text-fig. 8, pl. 3, figs. 7-9.

Material.- NU39-0063, left valve from loc. 32631.

Remarks.- The specimen from the Birafu Formation is relatively small, but it is similar in surface ornamentation and outline to the *Miltha japonica* Tashiro that was described from the Kesado Formation in Kyushu. The surface is marked with narrow concentric ribs, which are regular and usually much narrower than their interspaces.

Occurrence.- Locs. 32626, 23631.

Range. - Berriasian to Barremian.

Family CARDIIDAE Lamarck, 1809 Genus *Protocardia* von Beyrich, 1845 *Protocardia* sp. cf. *P. tosensis* Kimura Pl. 2, figs. 15-16

Material.- NU39-0069, right valve from loc. 23656; NU39-0070, left valve from loc.32656.

Remarks.- The outline of the specimens from the Birafu Formation resembles that of *Protocardia tosensis* Kimura from the Torinosu Group in Shikoku. But the poor preservation and deformation of these specimens makes accurate comparison difficult.

Occurrence.- Locs. 32626, 32631, 32656. Range. - Kimmeridgian to Berriasian.

Family ARCTICIDAE Newton, 1891 Genus *Isocyprina* Roeder, 1882 *Isocyprina japonica* Tashiro and Kozai Pl. 2, figs. 9-11

Isocyprina aliquantula (Amano), Hayami 1965b, p. 137, pl. 19, figs. 8-9; *Isocyprina japonica* Tashiro and Kozai, 1989, p. 118-119, pl. 2, text-fig. 2A, figs. 3-9.

Material.- NU39-0076, 0077 and 0081, from loc. 23631, internal molds of right valves; NU39-0078 from loc. 32653 and NU39-0079 and 0080 from loc. 32626, internal molds of left valves; NU39-0082 from loc. 32631, external mold of right valve.

Remarks.- The specimens from Birafu Formation are similar to the species *Isocyprina japonica* Tashiro and Kozai. The test of this species is thin, hence detailed hinge structure can not be observed in these specimens.

Occurrence.- Locs. 32626, 32653, 32631. Range.- Berriasian to Aptian.

Family TELLINIDAE de Blainville, 1814 Genus *Linearia* Conrad, 1860 *Linearia* sp. aff. *L. nankaiana* Tashiro and Kozai Pl. 2, figs. 19-20

Material.- NU39-0083, internal mold of right valve from loc. 32631; NU-39-0084, external mold of right valve from loc. 32631.

Remarks.- Our specimen from the Birafu Formation is compared with *Linearia nankaiana* Tashiro and Kozai. The complete specimen (NU39-0083) is characterized by large size (length 30.8 mm, height 15.2 mm), nearly flat outline and smooth surface. The outline of *Linearia nankaiana* Tashiro and Kozai from the Igenoki Formation in Shikoku is weakly inflated and smaller than this specimen. The surface of this species is covered with very fine radial striae with distinct posterior carina. In this respect, the Birafu specimens differ to *Linearia nankaiana* Tashiro and Kozai. Detail hinge stracture can not be observed in these specimens.

Occurrence.- Loc. 32631.

Range. - Berriasian.

Family NEOMIODONTIDAE Casey, 1955 Genus *Eomiodon* Cox, 1935 *Eomiodon nipponicus* Ohta Pl. 2, figs. 26-27

Astarte sakawana Kobayashi and Suzuki, 1939, p. 219, pl. 13, figs. 12-13.

Eomiodon nipponicus Ohta, 1973, p.252, pl.1. figs. 15-22,

pl. 2. figs. 14-19; Hayami 1975, p. 138; Kozai *et al.* 2001, pl. 1, figs. 1-2.

Material.- NU39-0060, left valve from loc.32631; NU39-0061, external mold of right valve from loc.32631.

Remarks.- These specimens are identical with *Eomiodon nipponicus* Ohta from the Yoshimo Formation of Chugoku, in position of umbo, concentric ribs and shape.

Occurrence.- Locs. 32654, 32631, 32656.

Range. - Berriasian to Hauterivian.

Eomiodon kumamotoensis Tamura Pl. 2, figs. 17-18

Eomiodon kumamotoensis Tamura, 1959c, p. 115, pl. 12, figs. 17, 18; Tamura 1960a, p. 240; Tashiro and Katto 1986, pl. 1, figs. 15-17; Tashiro 1987, p. 93.

Astarte ? kumamotoensis (Tamura). Hayami 1975, p. 128. Material.- One complete right valve, NU39-0058, internal mold of right valve from loc.32631; NU39-0059, external mold of right valve from loc.32631.

Remarks.- These specimens are similar to *Eomiodon kumamotoensis* Tamura, its strongly inflated valve, small size and regularly disposed concentric ribs. However, it has more regular concentric ribs.

Occurrence.- Locs. 32631, 32656.

Range. - Kimmeridgian to Berriasian.

Genus Crenotrapezium Hayami, 1958 Crenotrapezium kitakamiense Hayami Pl. 2, figs. 30-31

Crenotrapezium kitakamiense Hayami, 1960, p.17, textfig. 2, figs. 8-10; Hayami 1975, p. 141; Tashiro and Kozai, 1989, p. 126, pl. 2, figs. 11-13.

Material.- NU39-0071, left valves from loc. 32656; NU-39-0072–0074, left valves from loc. 32631; NU39-0075, right valve from loc.32656.

Remarks.- These specimens are identical with *Crenotrapezium kitakamiense* Hayami in having a distinct posterior carina and subtrigonal outline. They are also similar to *C. kitakamiense kurumense* Hayami, but the posterior extremity of *C. kitakamisense kitakamiense* is more elongated.

Occurrence.- Locs. 32631, 32656. Range. - Berriasian to Barremian.

> Order MYOIDA Striczka, 1870 Family Corbulidae Lamarck, 1818 Genus Corbula Bruguière, 1797 Corbula grobosa Tamura Pl. 2, figs. 6-8

Corbula globosa Tamura, 1959c, p.114, pl.12, figs. 1-4; Tamura 1960a, p. 242; Tamura 1960b, p. 290-291, pl. 33, figs. 8-10; Hayami 1975, p. 145; Ohta and Moji 1976, pl. 1. figs. 8-10; Tashiro and Katto 1986, pl. 1, fig. 18.

Material.- NU39-0046-0047, right valve from loc. 32653; NU-0048–0049, internal mold of right valve from loc. 32631.

Remarks.- The specimens are similar in size, shape and surface ornamentation to *Corbula globosa* Tamura. This species is small and strongly inflated. Surface is ornamented with fine concentric ribs. The species differs from *Corbula imamurae* Hase from the Toyonishi Group in its smaller and more inflated valve.

Occurrence.- Locs. 32653, 32631. Range. - Oxfordian to Berriasian.

Corbula imamurae Hase Pl. 2, figs. 24-25

Corbula(?) *imamurae* Hase, 1960, p. 324, pl. 37, figs. 16-21, pl. 39, figs. 2-4; Kozai 1989, p. 43-44, Fig. 4, nos. 23–25.

Material.- NU39-0044, internal mold of left valve from loc.32653; NU39-0045, internal mold of left valve from 32654.

Remarks.- The specimens from the Birafu Formation are

internal mold of left valves, so surface ornamentation is not preserved, however the outline of these specimens are similar to *Corbula imamurae* Hase.

Occurrence.- Locs. 32653, 32654, 32656.

Range. - Berriasian to Hauterivian.

Genus *Caestocorbla* Vincent, 1910 *Caestocorbula morinoi* Tashiro and Kozai Pl. 2, figs. 1-3

Caestocorbula morinoi Tashiro and Kozai, 1991, p.193-194, text-fig. 5, pl. 1, figs. 12-18.

Material.- NU39-0053 and 0054, internal molds of right valve, loc.23653; NU390055 and 0057, internal molds of .right valve, loc.32631; NU39-0056, external mold of right valve, loc.32631.

Remarks.- This species is distinguished from other species of Genus *Caestocorbula* by much narrowly elongated area.

Occurrence.- Locs.32653 and 32631. Range.- Berriasian.

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Plate 1. Upper Jurassic and Lower Cretaceous radiolarians from the Birafu Formation. Scale bars: 100 µm; a: 22; b: 1, 3, 5, 7, 9, 25, 30, 31; c: 2, 4, 6, 8, 10–24, 26–28; d: 29. Fig. 1, Cinguloturris cylindra Kemkin and Rudenko, from 32607R in C. Fig. 2, Archaeodictyomitra tumandae Dumitrica, from 32607R in C. Fig. 3, Ristola aff. cretacea (Baumgartner), from 32607R in C. Fig. 4, Eucyrtidiellum pyramis Aita, from 32606R in C. Fig. 5, Tethysetta boesii (Parona), from 32604R in B2. Fig. 6, Tethysetta pygmaea Dumitrica, from 32607R in C. Fig. 7-8, Pseudodictyomitra carpatica (Lozyniak), from 32604R in B2. Fig. 9, Archaeodictyomitra apiarium (Rüst), from 32604R in B2. Fig. 10, Sethocapsa pseudouterculus Aita, from 32604R in B2. Fig. 11, Sethocapsa kaminogoensis Aita, from 32604R in B2. Fig. 12, Archaeodictyomitra mitra Dumitrica, from 32604R in B2. Fig. 13, Loopus primitivus (Matsuoka and Yao), from 32604R in B2. Fig. 14, Loopus primitivus (Matsuoka and Yao), from 32602R in A3. Fig. 15, Svinitzium pseudopuga Dumitrica, from 32602R in A3. Fig. 16, Svinitzium depressum (Baumgartner), from 32607R in C. Fig. 17, Solenotrima (?) ichikawai Matsuoka and Yao, from 32602R in A3. Fig. 18, Archeodictyomitra minoensis (Mizutani), from 32602R in A3. Fig. 19, Cinguloturris carpatica Dumitrica, from 32602R in A3. Fig. 20, Protunuma japonicus Matsuoka and Yao, from 32602R in A3. Fig. 21, Eucyrtidiellum ptyctum (Riedel and Sanfilippo), from 32602R in A3. Fig. 22, Ristola altissima (Rüst), from 32602R in A3. Fig. 23, Loopus yangi Dumitrica, from 32602R in A3. Fig. 24, Loopus doliolum Dumitrica, from 32602R inA3. Fig. 25, Pantanellium lanceola (Parona), from 32607R in C. Fig. 26, Kilinora spiralis (Matsuoka), from 32601R in A2. Fig. 27, Tricolocapsa conexa Matsuoka, from 32601R in A2. Fig. 28, Stichocapsa robusta Matsuoka, from 32601R in A2. Fig. 29, Eucyrtidiellum nodosum Wakita, from 32601R in A2. Fig. 30, Hsuum maxwelli Pessagno, from 32601R in A2. Fig. 31, Ristola procera (Pessagno), from 32601R in A2.



Plate 2. Berriasian bivalves from the Birafu Formation. Figs. 1-3, Caestocorbula morinoi Tashiro and Kozai. 1: NU39-0054, from 32631, ×2.5, 2: NU39-0053, from 32631, ×2, 3: NU39-0057, from 32631, ×2.5. Figs. 4-5, Grammatodon takiensis Kimura. 4: NU39-0038a, from 32655, ×2. 5: NU39-0038b, from 32655, ×2. 6-8, Corbula globosa Tamura. Fig. 6: NU39-0049, from 32631, ×3. Fig. 7: NU39-0048, from 32631, ×3. Fig. 8: NU39-0047, from 32653, ×4. Figs. 9-11, Isocyprina japonica Tashiro andÅ@Kozai. 9: NU39-0081, from 32631, ×2. 10: NU39-0076, from 32631, ×2.5. 11: NU39-0078, from 32653, ×3. Fig. 12, Parallelodon sp. cf. P. koikensis Tamura, NU39-0085, from 32631, ×4. Figs. 13-14. Ctenoides tosana (Kurata and Kimura). 13: NU39-0041, from 32626, ×1.5. 14: NU39-0042, from 32626, ×1.5. Figs. 15-16, Protocardia sp. cf. P. tosensis Kimura. Fig. 15: NU39-0069, from 32656, ×1, Fig. 16: NU39-0070, from 32656, ×1. Figs. 17-18, Eomiodon kumamotoensis Tamura. 17: NU39-0059, from 32631, ×2. 18: NU39-0058, from 32631, ×2. Figs. 19-20, Linearia sp. aff. L. nankaiana Tashiro and Kozai. 19: NU39-0084, from 32631, ×1. 20: NU39-0083, from 32631, ×1. Figs. 21-22, Arcomytilus sp. cf. A. laitmairensis (de Loriol). 21: NU39-0087, from 32626, ×1.5. 22: NU39-0086, from 32626, ×2. Fig. 23, Parvamussium habunokawense (Kimura), NU39-0062, from 23631, ×2. Figs. 24-25, Corbula imamurae Hase. Fig. 24: NU39-0044, from 32653, ×2. Fig. 25: NU39-0045, from 32654, ×2. Figs. 26-27, Eomiodon nipponicus Ohta. Fig. 26: NU39-0026, from 32631, ×2. Fig. 27: NU39-0061, from 32631, ×2. Figs. 28-29, Pterotrigonia toyamai (Yehara). Fig. 28: NU39-0069, from 32626, ×1.5. Figs. 29: NU39-0065, from 32626, ×1.5. Figs. 30-31, Crenotrapezium kitakamiense Hayami. 30: NU39-0073, from 32631, ×1. 31: NU39-0074, from 32631, ×1. Fig. 32, Miltha japonica Tashiro, NU39-0063, from 32631, ×1. Fig. 33, Cucullaea sp. aff. C. acuticarinata Nagao, NU39-0040, from 32626, ×1.5. Figs. 34-35, Agillerella nagatoensis (Ohta). 34: NU39-0050a, from 32631, ×0.5. 35: NU39-0050b, from 32631, ×0.5.