

LETTER

Structural investigation of Mg local environments in silicate glasses by ultra-high field ²⁵Mg 3QMAS NMR spectroscopy

KEIJI SHIMODA,^{1,*} YASUHIRO TOBU,¹ MORIAKI HATAKEYAMA,¹
TAKAHIRO NEMOTO,² AND KOJI SAITO¹

¹Advanced Technology Research Laboratories, Nippon Steel Corporation, 20-1 Shintomi, Futtsu, Chiba 293-8511, Japan

²JEOL Ltd., 3-1-2 Musashino, Akishima, Tokyo 196-8558, Japan

ABSTRACT

Structural information on divalent cations such as Mg²⁺ should have important implications for magmatic liquids because of their abundance in the Earth's interior; nevertheless, little is confirmed about their coordination environments. We here apply a ²⁵Mg triple-quantum magic-angle spinning (3QMAS) NMR technique at an ultra-high magnetic field (21.8 T) and successfully show the occurrence of multiple Mg sites in MgSiO₃ glass. We find that these sites are distinguished by the degree of polyhedral distortion, not by the coordination number. The present study concludes that the highly distorted MgO₆ species occur in MgSiO₃ glass, in strong contrast with a recent radial distribution study.

Keywords: Mg local environment, MgSiO₃ glass, ²⁵Mg 3QMAS NMR, ultra-high magnetic field

INTRODUCTION

Alkali and alkaline earth cations act as network modifiers in silicate melts and glasses. Most studies on their structures have focused on SiO₄ and AlO₄ tetrahedral networks, whereas the lack of local information about the monovalent or divalent cations has been a long-standing problem in understanding the complete picture of amorphous structures. Some researchers have examined the local environments of the Mg²⁺ ion in MgSiO₃ and CaMgSi₂O₆ glasses, which are still controversial issues. The coordination number, one of the most important factors affecting dynamic properties, of Mg in such glasses was concluded as fourfold coordination in some studies (Tabira 1996; Wilding et al. 2004) and fivefold (Ildefonse et al. 1995; Li et al. 1999) or sixfold in others (Kroeker and Stebbins 2000).

NMR spectroscopy is a robust tool for collecting structural information focused on a specific element in multi-component systems. Unfortunately, ²⁵Mg NMR has been difficult to apply because of its low resonance frequency (18.4 MHz at 7.0 T) and relatively low natural abundance (10.1%). Using a high magnetic field and ²⁵Mg isotopic enrichment drastically alleviates these shortcomings. Nevertheless, ²⁵Mg MAS NMR further suffers from the large second-order quadrupolar broadening. To overcome this, we here applied the ²⁵Mg 3QMAS technique to Mg-containing silicate glasses and clarified the detailed environments surrounding the Mg²⁺ ions in these glasses.

EXPERIMENTAL PROCEDURE

A series of ²⁵Mg-enriched glass samples (MgSiO₃, CaMgSi₂O₆, Ca₂MgSi₂O₇, Mg₃Al₂Si₃O₁₂, Na₂MgSi₂O₆, and K₂MgSi₂O₆ glasses) and diopside (CaMgSi₂O₆)

and åkermanite (Ca₂MgSi₂O₇) crystals were prepared from appropriate mixtures of 99%-enriched ²⁵MgO, Ca(OH)₂, K₂CO₃, Al₂O₃, and SiO₂ reagents. The MgSiO₃, CaMgSi₂O₆, Ca₂MgSi₂O₇, and Mg₃Al₂Si₃O₁₂ glasses were obtained by melting the powder mixtures at 1500–1700 °C and quenching in air or water, but Na₂MgSi₂O₆ and K₂MgSi₂O₆ glass was melted at 1300 °C. XRD indicated that they were amorphous. Diopside and åkermanite were prepared by melting at 1500 °C and subsequent annealing at 1000 °C for 6 hours.

The ²⁵Mg NMR measurements were conducted on a JNM-ECA930 spectrometer (21.8 T) at the National Institute for Materials Science. The working frequency was 56.9 MHz. The 90° pulse lengths were optimized to 1.6–1.8 μs for the glasses. Spin-lattice relaxation times *T*₁ were estimated to be less than 0.5 s by the saturation recovery method. All the MAS spectra were acquired with a spin-echo sequence (a pulse length of 1.8 and 3.6 μs) to avoid acoustic ringing, and a recycle delay of 0.5 s. Longer delays did not influence the peak shape.

The ²⁵Mg 3QMAS spectra were acquired with a z-filter sequence and a delay of 0.5 s. Triple-quantum excitation and subsequent conversion pulses were optimized to 6.0 and 2.5 μs, respectively. Samples set in 4 mm MAS rotors were rotated at a spinning rate of 18 kHz. Chemical shifts were externally referenced to MgSO₄ solution at 0.0 ppm.

A series of FID signals in the 3QMAS measurements were Fourier-transformed and sheared to obtain the separate isotropic (F1) and anisotropic (F2) dimensions. The isotropic chemical shift δ_{CS} and quadrupolar product *P*_Q were derived from the sheared two-dimensional spectra with the following equations (Amoureux et al. 2002);

$$\delta_{CS} = \frac{17}{27}\delta_{F1} + \frac{10}{27}\delta_{F2},$$
$$P_Q = \sqrt{\frac{170 [4S(2S-1)]^2}{81 [4S(S+1)-3]} (\delta_{F1} - \delta_{F2})} \cdot \nu_0 \cdot 10^{-3},$$

where δ_{F1} and δ_{F2} are the chemical shifts in F1 and F2 dimensions, ν₀ is the Zeeman frequency, and *S* is the spin number. The *P*_Q is also related to the quadrupolar coupling constant *C*_Q with the equation of *P*_Q = *C*_Q√(1 + η²)/3, where η is the asymmetry parameter (η = 0 to 1).

RESULTS AND DISCUSSION

Figure 1 shows the MAS spectra for Mg-containing silicate glasses. There are apparent differences between MgSiO₃, CaMgSi₂O₆, Ca₂MgSi₂O₇, Mg₃Al₂Si₃O₁₂ glasses (Figs. 1a–1d)

* E-mail: kshimoda@re.nsc.co.jp

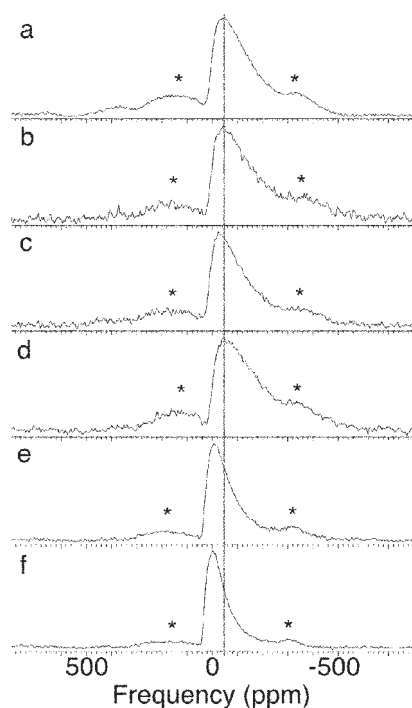


FIGURE 1. ^{25}Mg MAS NMR spectra at 21.8 T for (a) MgSiO_3 , (b) $\text{CaMgSi}_2\text{O}_6$, (c) $\text{Ca}_2\text{MgSi}_2\text{O}_7$, (d) $\text{Mg}_3\text{Al}_2\text{Si}_3\text{O}_{12}$, (e) $\text{Na}_2\text{MgSi}_2\text{O}_6$, and (f) $\text{K}_2\text{MgSi}_2\text{O}_6$ glasses. 22 000–96 000 transients were averaged with relaxation delays of 0.5 s. The asterisks refer to spinning side bands. The vertical line is a guide for the eye.

TABLE 1. Estimates of isotropic chemical shift δ_{CS} and quadrupolar product P_Q from ^{25}Mg 3QMAS spectra at 21.8 T

		δ_{CS} (ppm)*	P_Q (MHz)*
MgSiO_3 glass	site A	10	3.5
	site B	24	3.3
	site C	9	5.8
	site D	1	7.3
$\text{K}_2\text{MgSi}_2\text{O}_6$ glass		37	2.8
diopside (^{51}Mg crystal)		6	2.2
åkermanite (^{51}Mg crystal)		49	2.9

* Errors are within 3 ppm and 0.5 MHz, respectively. The estimates for the crystals were from 3QMAS spectra at 700 MHz (not shown).

and $\text{Na}_2\text{MgSi}_2\text{O}_6$, $\text{K}_2\text{MgSi}_2\text{O}_6$ glass (Figs. 1e and 1f). The observed peak maxima are at -47.1 , -46.3 , -25.7 , -44.6 , -3.4 , and -0.9 ppm, respectively. In this case, the peak maxima are not isotropic chemical shifts due to quadrupolar interactions, and from the MAS data at a single field, the isotropic shifts cannot be obtained for quadrupolar nuclei in glasses. Also, the signal widths (FWHM) of the former group of glasses (140–190 ppm) are significantly larger than the latter (90–110 ppm). These results imply that the Mg local environments in the former glasses are different from those in the latter.

Figures 2a and 2b show the ^{25}Mg 3QMAS spectra of MgSiO_3 and $\text{K}_2\text{MgSi}_2\text{O}_6$ glasses, respectively. The multiple-quantum (MQ) MAS technique is an excellent method for observing high-resolution NMR spectra, free from second-order broadening of half-integer quadrupolar nuclei (Frydman and Harwood 1995). Figure 2a shows four Mg sites in the MgSiO_3 glass structure, in contrast to a single site for $\text{K}_2\text{MgSi}_2\text{O}_6$ glass. The estimated

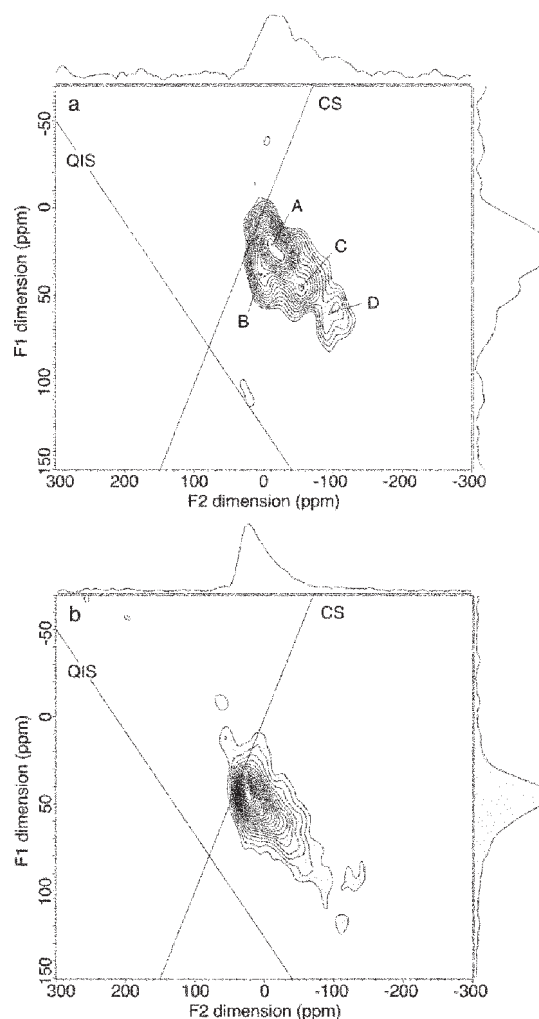


FIGURE 2. ^{25}Mg 3QMAS spectra at 21.8 T for (a) MgSiO_3 and (b) $\text{K}_2\text{MgSi}_2\text{O}_6$ glasses. 52 800 and 14 400 transients were collected for each $16 t_1$ increment with delays of 0.5 s, respectively. Contour levels are equally drawn at 5% to 95% with an increment of ca. 5.3%, and 4% to 96% with an increment of ca. 4.2%, respectively.

isotropic chemical shifts (δ_{CS}) and quadrupolar products (P_Q) for the glasses, diopside (a ^{51}Mg site) and åkermanite (a ^{51}Mg site), are listed in Table 1. The values for diopside and åkermanite agree with those reported by Kroeker and Stebbins (2000). MacKenzie and Meinhold (1994) reported the ^{25}Mg MAS spectrum for the fivefold Mg site in granddierite. Its δ_{CS} was estimated to be 55 ppm, which is higher in frequency than that for the fourfold site in åkermanite. This may imply that the Mg site is largely affected by neighboring cations (BO_3 and AlO_3). The δ_{CS} of an eightfold Mg mineral such as pyrope is also unclear (Dupree and Smith 1988). MacKenzie and Meinhold (1994) mentioned that the δ_{CS} of sixfold Mg-containing silicates ranges from 5 to 14 ppm. Thus, we consider that in MgSiO_3 glass the sites A, C, and probably D are due to sixfold Mg species with different environments. Site B may be from a fivefold environment. In the case of $\text{K}_2\text{MgSi}_2\text{O}_6$ glass, we obtained only a single peak with a broad distribution along the quadrupolar-induced shift

(QIS) axis. The δ_{CS} of 37 ppm is less positive in frequency than that for åkermanite, but within the rough estimate for crystalline $\text{K}_2^{IV}\text{MgSi}_5\text{O}_{12}$ (30–50 ppm) (Kroeker and Stebbins 2000). We here conclude that the Mg^{2+} ions occur in a tetrahedral environment, but might be fivefold as deduced previously (Fiske and Stebbins 1994; George and Stebbins 1998). The linewidth of the F2 projections of the 3QMAS spectra for the both glasses are narrower than the corresponding MAS spectra, although they are identical in principle. This may result from the C_Q -dependence of MQ excitation efficiency. Other sites could exist with larger C_Q (P_Q), but they are not observed here.

Our conclusion that $\text{K}_2\text{MgSi}_2\text{O}_6$ glass has a single MgO_4 site supports the results of X-ray emission (Kawazoe et al. 1981) and MAS NMR (Kroeker and Stebbins 2000). The difference in the coordination state between $\text{K}_2\text{MgSi}_2\text{O}_6$ and MgSiO_3 glasses is explained by the effect of cation field strength. The cation field strength, Z/d^2 (Z is the charge and d is M-O distance) increases in the order of K (0.12), Na (0.18), Ca (0.36), and Mg (0.46–0.53) (Brown et al. 1995). This indicates an increase in bond strength in the order of K-O, Na-O, Ca-O, and Mg-O. Thus, in the $\text{K}_2\text{MgSi}_2\text{O}_6$ glass, Mg^{2+} ions strongly and effectively bind with oxygen atoms over competing K^+ ions to reduce the overall Mg coordination number (Kroeker and Stebbins 2000). If we consider that the MgO_4 tetrahedra partly act as weak network formers in $\text{K}_2\text{MgSi}_2\text{O}_6$ glass, they may influence bulk viscosity. Indeed, $\text{K}_2\text{MgSi}_5\text{O}_{12}$ melt shows a decrease in viscosity with an increase in pressure at constant temperature, whereas $\text{CaMgSi}_2\text{O}_6$ melt shows an increase in viscosity (Scarfe et al. 1979; Kushiro 1984; see also Wolf and McMillan 1995). The viscosity is most likely affected by the degree of network polymerization and cation coordination. Thus, this can be explained by the mechanism that the highly polymerized $(\text{Mg},\text{Si})\text{O}_4$ tetrahedral network involves an increase in the Mg coordination number with pressure, resulting in a decrease in viscosity. This may be supported by an ^{17}O 3QMAS study of K-Mg silicate glasses (Allwardt and Stebbins 2004), in which K^+ was found to bond with bridging oxygen atoms ($^{IV}\text{Mg-O-}^{IV}\text{Si}$ in our definition), suggesting its charge balancing role. On the other hand, the less-polymerized $\text{CaMgSi}_2\text{O}_6$ melt is considered to take pressure-induced polymerization of SiO_4 tetrahedral network to effectively increase viscosity (Dickinson et al. 1990), probably with a small counteraction of the Mg coordination increase.

The ^{25}Mg MAS and 3QMAS spectra suggest that Mg^{2+} is mainly in octahedral coordination in MgSiO_3 , $\text{CaMgSi}_2\text{O}_6$, $\text{Ca}_2\text{MgSi}_2\text{O}_7$, and $\text{Mg}_3\text{Al}_2\text{Si}_3\text{O}_{12}$ glasses. For $\text{CaMgSi}_2\text{O}_6$ glass, the present study agrees with the previous MAS NMR study (Kroeker and Stebbins 2000) rather than the XAS studies (Ildelfonse et al. 1995; Li et al. 1999). The present study indicates that the Mg environments in $\text{Ca}_2\text{MgSi}_2\text{O}_7$ and $\text{Mg}_3\text{Al}_2\text{Si}_3\text{O}_{12}$ glasses are different from the crystalline counterparts, åkermanite and pyrope, where Mg^{2+} is in a fourfold and an eightfold site, respectively. Our result is also in contrast with X-ray and molecular dynamics studies for pyrope glass (Okuno and Marumo 1993; Okuno and Kawamura 1995). The MAS spectrum of $\text{Ca}_2\text{MgSi}_2\text{O}_7$ glass shows slightly higher frequency than those of the other three glasses. This may suggest a larger amount of fivefold Mg species. Moreover, Wilding et al. (2004) recently reported the Mg coordination number of 4.5 for MgSiO_3 glass, based on a

reverse Monte Carlo (RMC) simulation with X-ray and neutron diffraction. Our study, however, contradicts their result. This may stem from the short cut-off distance in their study. They also insisted that the mean coordination number increased from 4.5 to 5.0 with an increase in MgO content from MgSiO_3 to Mg_2SiO_4 glass. Another RMC study for Mg_2SiO_4 glass estimated that MgO_4 , MgO_5 , and MgO_6 polyhedra were present in a ratio of 30%, 50%, and 20%, respectively (Kohara et al. 2004). In combination with the present study, we conclude that the Mg coordination decreases from 6 in MgSiO_3 glass to 5 in Mg_2SiO_4 glass. Insufficient numbers of SiO_4 tetrahedra (nominally no bridging oxygen atoms as in olivine) need some MgO_4 units as additional weak network formers to make a network structure in Mg_2SiO_4 glass.

The present 3QMAS NMR spectroscopy suggests the existence of several Mg environments in the MgSiO_3 glass structure (Fig. 2a). The multiple sites can be associated with the diversity of SiO_4 tetrahedral network connectivity. Several spectroscopic studies suggested that the alkaline earth silicate glasses have broad Q^n distributions (Murdoch et al. 1985; Kubicki et al. 1992; Schneider et al. 2003). Mg^{2+} ions should locate in the interstitial positions among the SiO_4 linkage, resulting in the environmental diversity. Another significance is that these sites (A, C, and D) should be distinguished by the degree of MgO_6 octahedral distortions rather than by the difference in coordination states. Site A has a small P_Q value, indicating relatively regular octahedral sites. On the other hand, the large- P_Q sites C and D are attributed to highly distorted octahedra with large electric field gradients (EFG) (MacKenzie and Meinhold 1994). Such polyhedral distortions are also confirmed by inspecting an asymmetrical Mg-O peak in a radial distribution function (Wilding et al. 2004), and Yin et al. (1983) have insisted that MgSiO_3 glass has four Mg-O bonds at 2.08 Å and two additional bonds at 2.5 Å, based on the quasi-crystalline model. The 3QMAS spectrum clearly shows such distorted sites in MgSiO_3 glass. On the other hand, the 3QMAS spectrum of $\text{K}_2\text{MgSi}_2\text{O}_6$ glass shows a single peak with a long tail along the QIS axis (Fig. 2b). This indicates that the Mg^{2+} ions, among the competing K^+ ions, locate in a regular site coordinated by four oxygen atoms, although additional distorted sites may be hidden in the tail.

In the upper mantle, olivine, pyroxene, and garnet are believed to be abundant mineral components. Thus, primitive magmas are Mg-rich liquids, and the structural contribution from Mg^{2+} ions, as well as SiO_4 tetrahedral connectivity, could be an important factor affecting the dynamic properties of the liquids. The increase in the Mg coordination number will easily occur compared with that of the Si atom at depth (Kubicki and Lasaga 1991; Shimoda and Okuno 2006). At moderate pressures where the Si coordination state remains unchanged, the Mg^{2+} ion may control the bulk viscosity of the liquids, involving the pressure-induced network polymerization discussed above (Dickinson et al. 1990). Our findings predict that the viscosity of MgSiO_3 , $\text{Ca}_2\text{MgSi}_2\text{O}_7$, and $\text{Mg}_3\text{Al}_2\text{Si}_3\text{O}_{12}$ melts (containing MgO_6) increases, but that of $(\text{K},\text{Na})_2\text{O-MgO-}n\text{SiO}_2$ melts (containing MgO_4) decreases at low pressures.

Our results further predict that the high- Z/d^2 divalent cations such as Fe^{2+} (0.44–0.51) and Ni^{2+} (0.48–0.55) should have local environments similar to Mg^{2+} ; they have highly distorted octa-

hedral sites in silicate glasses with no (or less) alkali cations, and the tetrahedral sites in the alkali-containing glasses (Waychunas et al. 1988; Keppler 1992; Farges et al. 2001). Jackson et al. (1993) have reported the Fe^{2+} coordination number of 4 for Fe_2SiO_4 melt, in reasonable agreement with that for Mg_2SiO_4 glass (Kohara et al. 2004).

The present study illustrates the potency of the ^{25}Mg MQMAS NMR technique under an ultra-high magnetic field, which shows the first evidence of multiple Mg sites with several degrees of distortion. We conclude that the Mg^{2+} ions in MgSiO_3 , $\text{CaMgSi}_2\text{O}_6$, $\text{Ca}_2\text{MgSi}_2\text{O}_7$, and $\text{Mg}_3\text{Al}_2\text{Si}_3\text{O}_{12}$ glasses are mainly in MgO_6 octahedral sites, in strong contrast with some previous studies. The ^{25}Mg MQMAS technique offers a great opportunity for understanding Mg local environments and their roles in organic and inorganic complexes (Grant et al. 2000; Hatakeyama et al. 2005).

ACKNOWLEDGMENTS

We express our gratitude to Y. Uozumi of IMS and T. Shimizu of NIMS for offering an opportunity to conduct the ultra-high field NMR experiments. We thank T. Nishiura and M. Imafuku for making and checking the ^{25}Mg -enriched samples. This work is fully supported by Special Coordination Funds for Promoting Science and Technology in Japan.

REFERENCES CITED

- Allwardt, J.R. and Stebbins, J.F. (2004) Ca-Mg and K-Mg mixing around non-bridging O atoms in silicate glasses: An investigation using ^{17}O MAS and 3QMAS NMR. *American Mineralogist*, 89, 777–784.
- Amoureux, J.-P., Huguenaud, C., Engelke, F., and Taulelle, F. (2002) Unified representation of MQMAS and STMAS NMR of half-integer quadrupolar nuclei. *Chemical Physics Letters*, 356, 497–504.
- Brown Jr., G.E., Farges, F., and Calas, G. (1995) X-ray scattering and X-ray spectroscopy studies of silicate melts. In J.F. Stebbins, P.F. McMillan, and D.B. Dingwell, Eds., *Structure, Dynamics and Properties of Silicate Melts*, 32, p. 317–410. *Reviews in Mineralogy, Mineralogical Society of America, Chantilly, Virginia*.
- Dickinson, J.E., Scarfe, C.M., and McMillan, P. (1990) Physical properties and structure of $\text{K}_2\text{Si}_4\text{O}_9$ melt quenched from pressures up to 2.4 GPa. *Journal of Geophysical Research*, 95, 15675–15681.
- Dupree, R. and Smith, M.E. (1988) Solid-state magnesium-25 n.m.r. spectroscopy. *Journal of Chemical Society, Chemical Communications*, 1483–1485.
- Farges, F., Brown Jr., G.E., Petit, P.-E., and Munoz, M. (2001) Transition elements in water-bearing silicate glasses/melts. Part I. A high-resolution and anharmonic analysis of Ni coordination environments in crystals, glasses, and melts. *Geochimica et Cosmochimica Acta*, 65, 1665–1678.
- Fiske, P.S. and Stebbins, J.F. (1994) The structural role of Mg in silicate liquids: A high-temperature ^{25}Mg , ^{23}Na , and ^{29}Si NMR study. *American Mineralogist*, 79, 848–861.
- Frydman, L. and Harwood, J.S. (1995) Isotropic spectra of half-integer quadrupolar spins from bidimensional magic-angle spinning NMR. *Journal of the American Chemical Society*, 117, 5367–5368.
- George, A.M. and Stebbins, J.F. (1998) Structure and dynamics of magnesium in silicate melts: A high-temperature ^{25}Mg NMR study. *American Mineralogist*, 83, 1022–1029.
- Grant, C.V., Frydman, V., and Frydman, L. (2000) Solid-state ^{25}Mg NMR of a magnesium(II) adenosine 5'-triphosphate complex. *Journal of the American Chemical Society*, 122, 11743–11744.
- Hatakeyama, M., Nemoto, T., Kanehashi, K., and Saito, K. (2005) Natural abundance solid-state ^{25}Mg MQMAS NMR studies on inorganic solids at a high magnetic field of 16.4 T. *Chemistry Letters*, 34, 864–865.
- Ildefonse, P., Calas, G., Flank, A.M., and Lagarde, P. (1995) Low Z elements (Mg, Al, and Si) K-edge X-ray absorption spectroscopy in minerals and disordered systems. *Nuclear Instruments and Methods in Physics Research B*, 97, 172–175.
- Jackson, W.E., Mustre de Leon, J., Brown Jr., G.E., Waychunas, G.A., Conradson, S.D., and Combes, J.-M. (1993) High-temperature XAS study of Fe_2SiO_4 liquid: reduced coordination of ferrous iron. *Science*, 262, 229–233.
- Kawazoe, H., Kokumai, H., Tanazawa, T., and Gohshi, Y. (1981) Coordination of Mg^{2+} in oxide glasses determined by X-ray emission spectroscopy. *Journal of Physics and Chemistry of Solids*, 42, 579–581.
- Keppler, H. (1992) Crystal field spectra and geochemistry of transition metal ions in silicate melts and glasses. *American Mineralogist*, 77, 62–75.
- Kohara, S., Suzuya, K., Takeuchi, K., Loong, C.-K., Grimsditch, M., Weber, J.K.R., Tangeman, J.A., and Key, T.S. (2004) Glass formation at the limit of insufficient network formers. *Science*, 303, 1649–1652.
- Kroeker, S. and Stebbins, J.F. (2000) Magnesium coordination environments in glasses and minerals: New insight from high-field magnesium-25 MAS NMR. *American Mineralogist*, 85, 1459–1464.
- Kubicki, J.D. and Lasaga, A.C. (1991) Molecular dynamics simulation of pressure and temperature effects on MgSiO_3 and Mg_2SiO_4 melts and glasses. *Physics and Chemistry of Minerals*, 17, 661–673.
- Kubicki, J.D., Hemley, R.J., and Hofmeister, A.M. (1992) Raman and infrared study of pressure-induced structural changes in MgSiO_3 , $\text{CaMgSi}_2\text{O}_6$, and CaSiO_3 glasses. *American Mineralogist*, 77, 258–269.
- Kushiro, I. (1984) Structures and some physical properties of silicate melts of geological interest. In I. Sunagawa, Ed., *Materials Science of the Earth's Interior*, p. 39–60. TERRAPUB, Tokyo.
- Li, D., Peng, M., and Murata, T. (1999) Coordination and local structure of magnesium in silicate minerals and glasses: Mg K-edge XANES study. *Canadian Mineralogist*, 37, 199–206.
- MacKenzie, K.J.D. and Meinhold, R.H. (1994) ^{25}Mg nuclear magnetic resonance spectroscopy of minerals and related inorganics: A survey study. *American Mineralogist*, 79, 250–260.
- Murdoch, J.B., Stebbins, J.F., and Carmichael, I.S.E. (1985) High-resolution ^{29}Si NMR study of silicate and aluminosilicate glasses: the effect of network-modifying cations. *American Mineralogist*, 70, 332–343.
- Okuno, M. and Kawamura, K. (1995) Molecular dynamics calculations for $\text{Mg}_3\text{Al}_2\text{Si}_3\text{O}_{12}$ (pyrope) and $\text{Ca}_3\text{Al}_2\text{Si}_3\text{O}_{12}$ (grossular) glass structures. *Journal of Non-Crystalline Solids*, 191, 249–259.
- Okuno, M. and Marumo, F. (1993) The structure analyses of pyrope ($\text{Mg}_3\text{Al}_2\text{Si}_3\text{O}_{12}$) and grossular ($\text{Ca}_3\text{Al}_2\text{Si}_3\text{O}_{12}$) glasses by X-ray diffraction method. *Mineralogical Journal*, 16, 407–415.
- Scarfe, C.M., Mysen, B.O., and Virgo, D. (1979) Changes in viscosity and density of sodium disilicate, sodium metasilicate, and diopside composition with pressure. *Carnegie Institute of Washington Year Book*, 78, 547–551.
- Schneider, J., Mastelaro, V.R., Zanutto, E.D., Shakhmatkin, B.A., Vedishcheva, N.M., Wright, A.C., and Panepucci, H. (2003) Q^0 distribution in stoichiometric silicate glasses: thermodynamic calculations and ^{29}Si high resolution NMR measurements. *Journal of Non-Crystalline Solids*, 325, 164–178.
- Shimoda, K. and Okuno, M. (2006) Molecular dynamics study of CaSiO_3 -Mg- SiO_2 glasses under high pressure. *Journal of Physics: Condensed Matter*, 18, 6531–6544.
- Tabira, Y. (1996) Local structure around oxygen atoms in $\text{CaMgSi}_2\text{O}_6$ glass by O K-edge EXELFS. *Materials Science and Engineering B*, 41, 63–66.
- Waychunas, G.A., Brown Jr., G.E., Ponader, C.W., and Jackson, W.E. (1988) Evidence from X-ray absorption for network-forming Fe^{2+} in molten alkali silicates. *Nature*, 332, 251–253.
- Wilding, M.C., Benmore, C.J., Tangeman, J.A., and Sampath, S. (2004) Evidence of different structures in magnesium silicate liquids: coordination changes in forsterite- to enstatite-composition glasses. *Chemical Geology*, 213, 281–291.
- Wolf, G.H. and McMillan, P.F. (1995) Pressure effects on silicate melt structure and properties. In J.F. Stebbins, P.F. McMillan, and D.B. Dingwell, Eds., *Structure, Dynamics and Properties of Silicate Melts*, 32, p. 505–561. *Reviews in Mineralogy, Mineralogical Society of America, Chantilly, Virginia*.
- Yin, C.D., Okuno, M., Morikawa, H., and Marumo, F. (1983) Structure analysis of MgSiO_3 glass. *Journal of Non-Crystalline Solids*, 55, 131–141.

MANUSCRIPT RECEIVED DECEMBER 10, 2006

MANUSCRIPT ACCEPTED JANUARY 19, 2007

MANUSCRIPT HANDLED BY BRYAN CHAKOUMAKOS